





# Dust prediction models

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# Questions will be welcome!



#### Introduction

#### What do we need to forecast dust storms?

- 1. Satellites, surface observations, NWP models and dust models.
- 2. Good knowledge of the dust climatology in the region.
- 3. Good knowledge of observation limitations.
- Good knowledge of the dust model limitations.

#### **Outlook**

#### 1. Dust cycle and associated processes

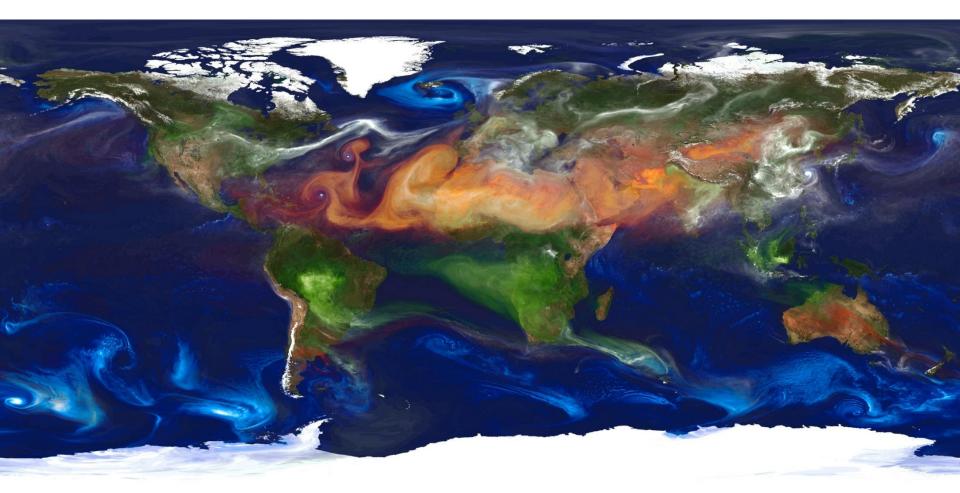
- The atmospheric dust cycle
- Dust global climatology
- Types of dust storms and model forecasting skills

#### 2. Dust forecasting models

- Dust emission schemes and dust sources
- Dust transport
- Dust deposition and sedimentation

#### 3. Modeling the dust cycle at BSC: From R&D to operational

# **Dust impacts and its extension**



**Organic Carbon + Elemental carbon** 

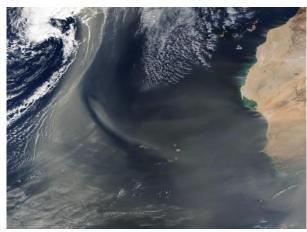
**Dust** 

Sulfate

Sea salt



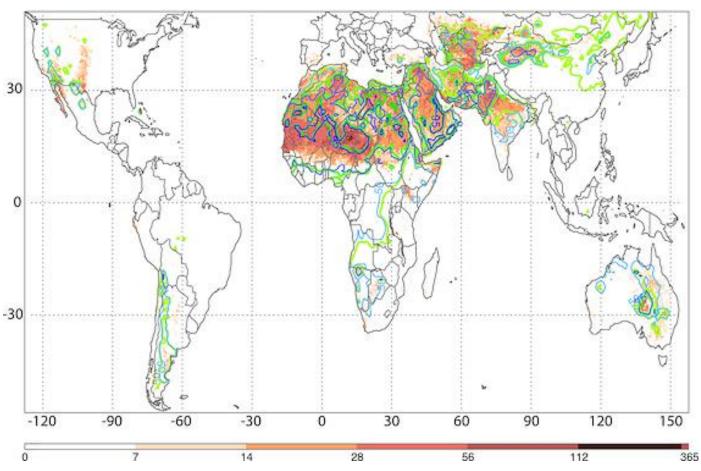
MODIS true colour composite image for March 2005 depicting a dust storm initiated at the Bodélé Depression (Chad Basin)



MODIS True color Western Africa –
Altantic Ocean

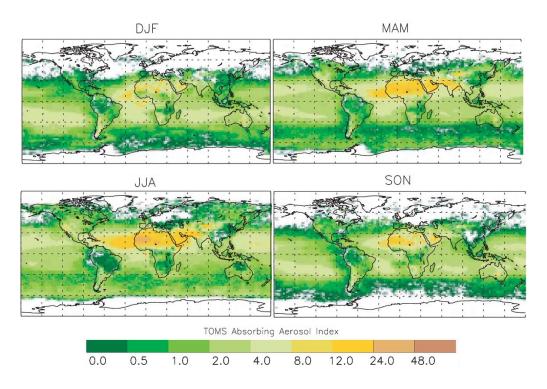
**Dust transport** is a global phenomenon. However, **dust emission** is a threshold phenomenon, sporadic and spatially heterogeneous, that is locally controlled on small spatial and temporal scales.

#### Dust global distribution



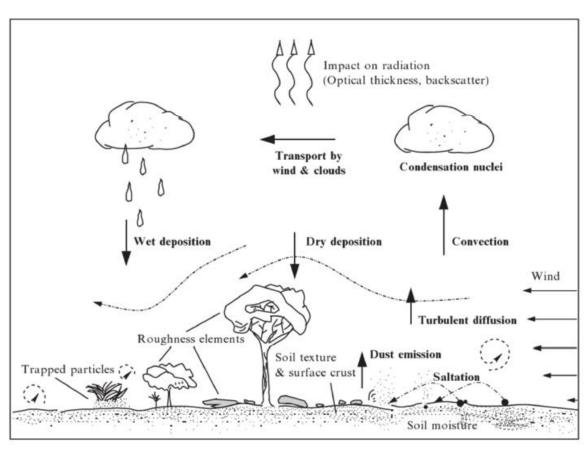
Global-scale attribution of anthropogenic and natural dust sources and their emission rates based on MODIS Deep Blue aerosol products by Ginoux et al. (2012)

Temporal changes in the dust distribution: SEASONAL and DECADAL CHANGES



- Seasonal dust distribution changes well characterized. Follows seasonal changing weather regimes (mainly) and vegetation changes (in semi-arid areas)
- Interannual/decadal changes are controlled by climate and surface modification (land use, desertification). Decadal changes are not well captures by models

The atmospheric dust cycle and involves a variety of processes:



- Dust emission from dry unvegetable surfaces (dust sources)
- Mid- and long-range transport
- Sedimentation, wet and dry deposition

Extracted from Shao (2008)

#### **Dust Impacts**

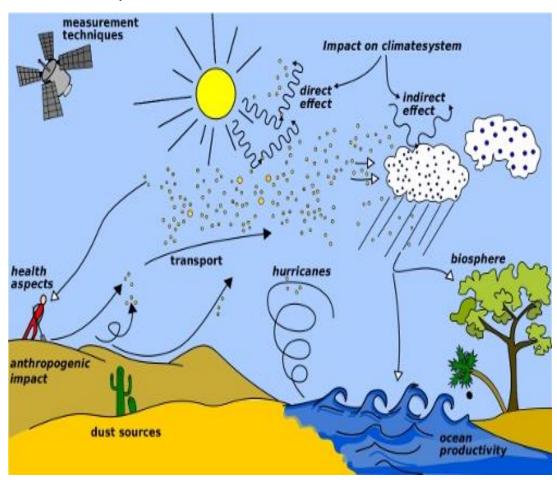


Image from WMO website (http://www.wmo.int/pages/prog/arep/wwrp/new/hurricanes.html)

# Ecosystems, meteorology and climate

- Marine productivity
- Coral mortality
- Hurricanes formation

#### Air Quality and Human Health

- Respiratory disease (asthma)
- Eye infections
- Meningitis in Africa
- Valley Fever in the Americas

#### **Aviation and Ground Transportation**

Low visibility (i.e. air disasters)

Agriculture and fishering

**Energy and industry** 

#### Types of dust storms:

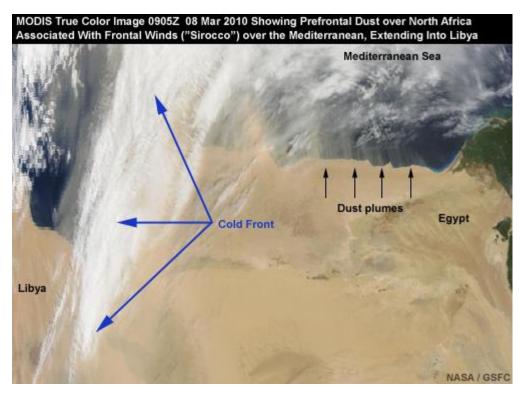
**Synoptic dust storms** (large scale weather systems)

- Prefrontal winds
- Postprontal winds
- Large-scale Trade winds
- ...

#### Mesoscale dust storms

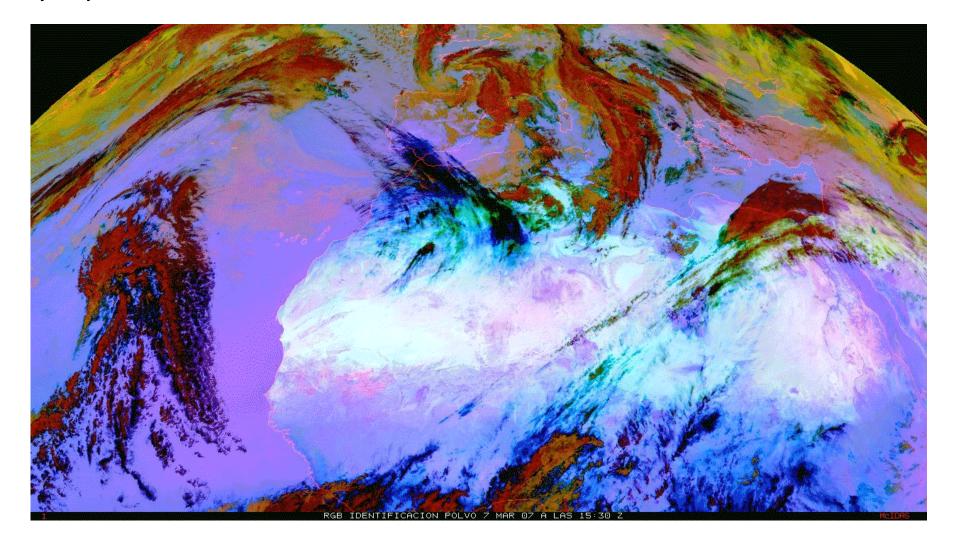
- Downslope winds
- Gap flow
- Convection (dust devils and Haboobs)
- Inversion downburst storms
- ...

Synoptic dust storms: Pre-frontal

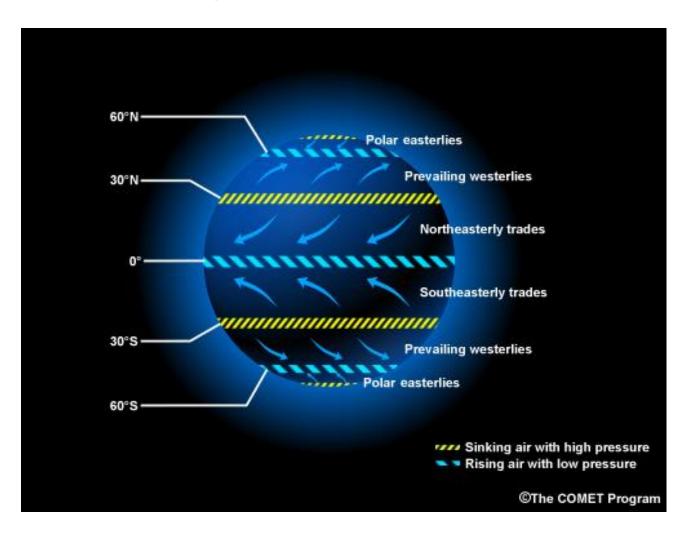




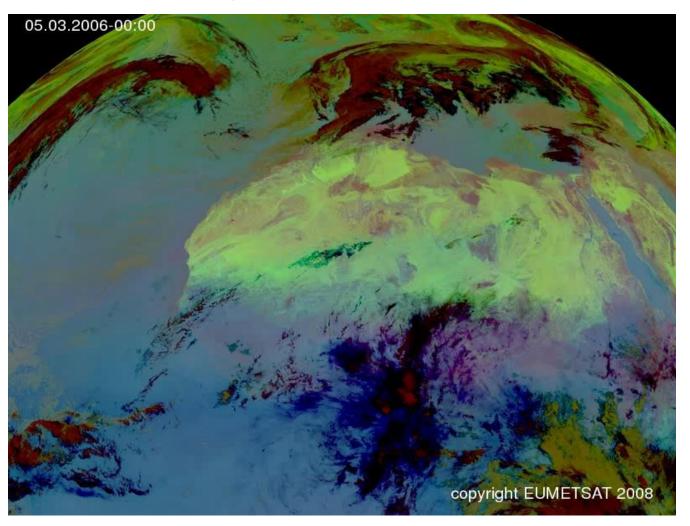
Synoptic dust storms: Post-frontal



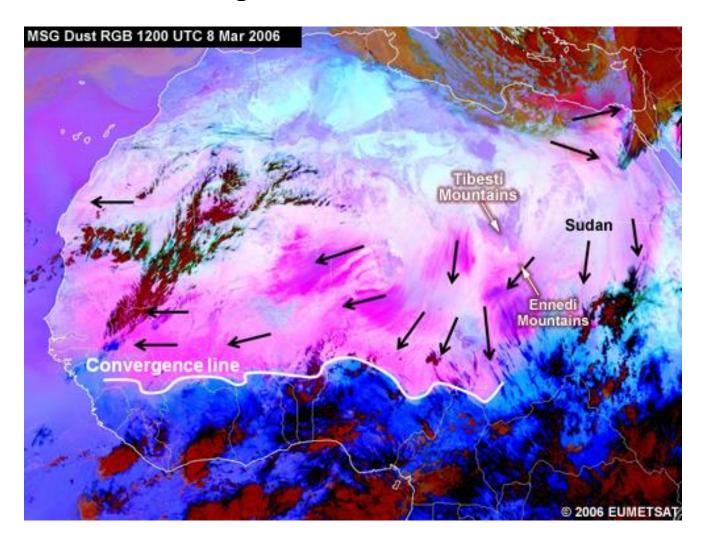
Synoptic dust storms: Large-scale trade winds



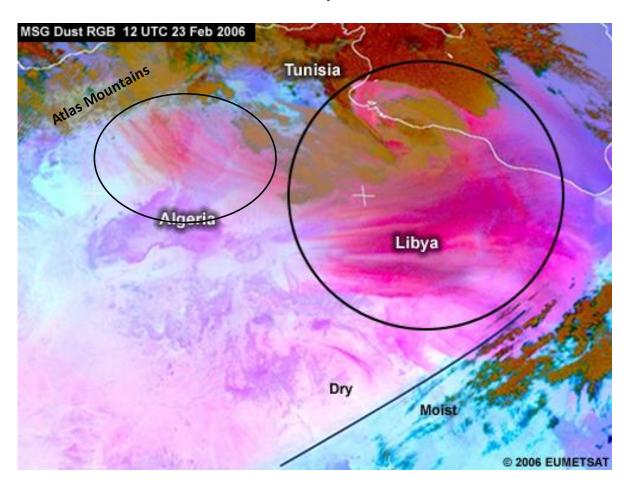
Synoptic dust storms: Large-scale trade winds



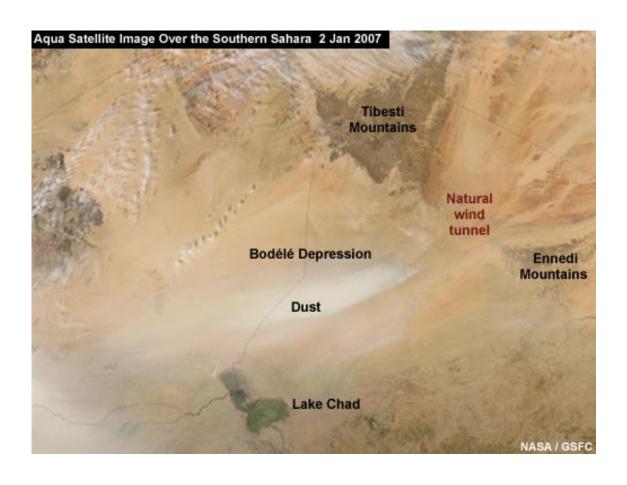
Synoptic dust storms: Large-scale trade winds



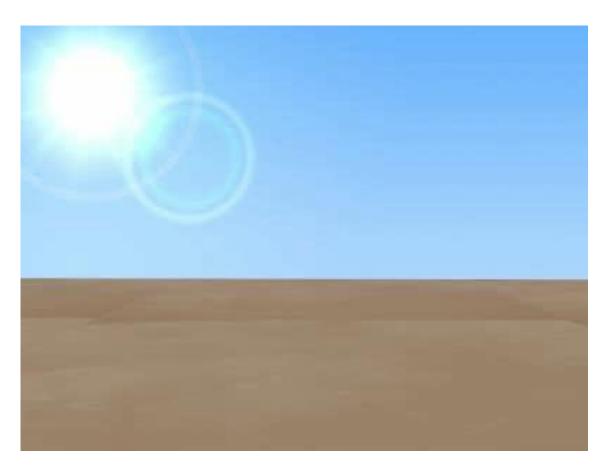
Mesoscale dust storms: Downslope winds



Mesoscale dust storms: Gap flow



Mesoscale dust storms: Dust devils (convection)



Movie from the COMET program at http://meted.ucar.edu/ of the University Corporation for Atmospheric Research (UCAR)

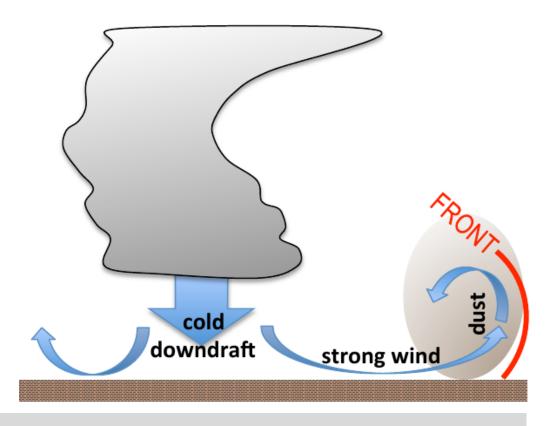
Mesoscale dust storms: Haboobs



Movie from the COMET program at http://meted.ucar.edu/of the University Corporation for Atmospheric Research (UCAR)

Mesoscale dust storms: Haboobs

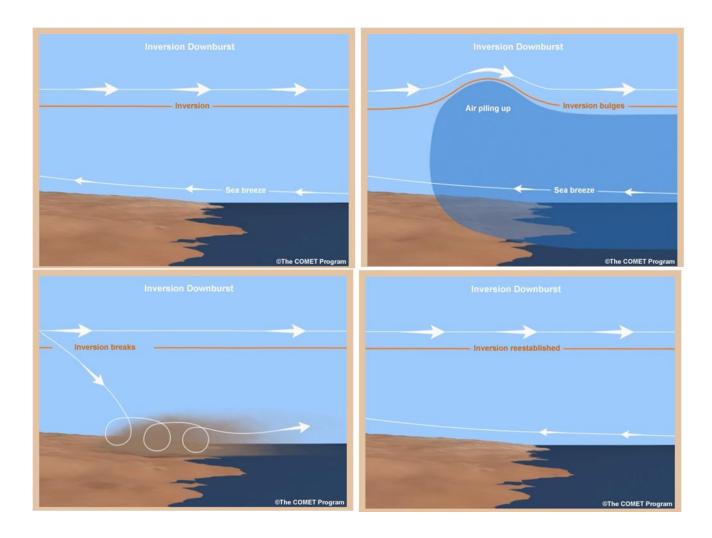
Intensive cold
downbursts from
convective cells
produced high velocity
surface wind, creating
cold front which was
lifting, mixing and
pushing dust



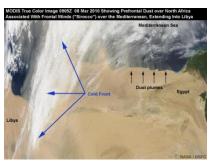
**Expected:** high wind speed, drop in temperature, rise in humidity, rise in pressure, reduction of visibility.

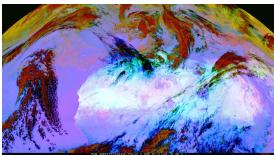


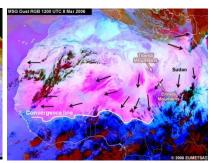
#### Mesoscale dust storms: Inversion downbursts



Synoptic dust storms (large scale weather systems) Well captured by models.







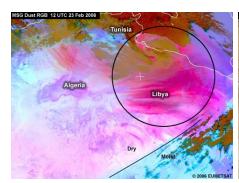
Pre-frontal winds

Post-frontal winds

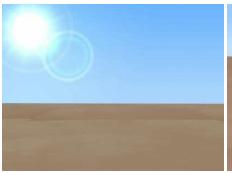
Large-scale trade winds

Mesoscale dust storms Poorly captured by models.

Some types improve in regional models.









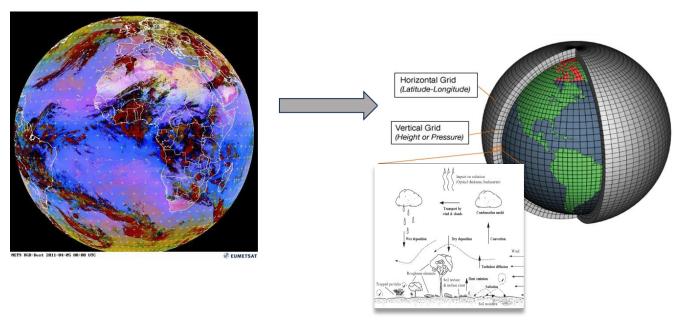
Downslope winds

Gap flow

Dust devils

Haboobs

Dust models are a **mathematical representation** of atmospheric dust cycle.



- ✓ To complement dust-related observations, filling the temporal and spatial gaps of the measurements.
- ✓ To help us to understand the dust processes and their interaction with climate and ecosystems.
- ✓ To predict the impact of dust on surface level concentrations used as **SHORT-TERM FORECASTING TOOLS** (3-5 days ahead)

Dust forecasting models do **not** take account dust **resuspension** 



Atmos. Chem. Phys., 14, 11753–11773, 2014 www.atmos-chem-phys.net/14/11753/2014/ doi:10.5194/acp-14-11753-2014 © Author(s) 2014. CC Attribution 3.0 License.

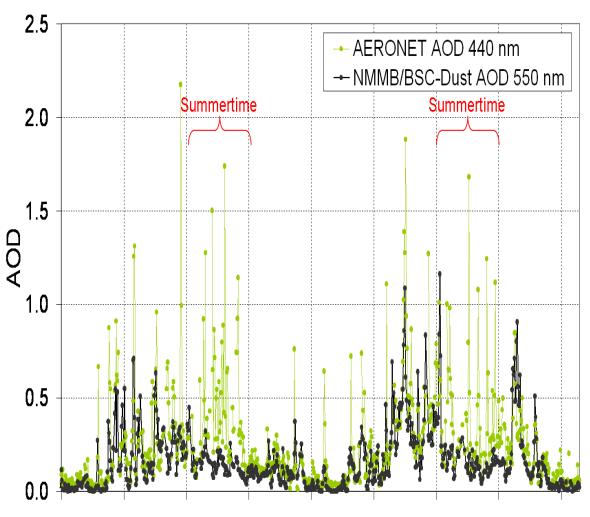




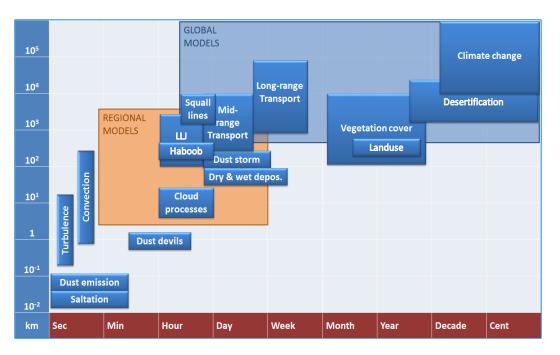
#### Aerosol characterization at the Saharan AERONET site Tamanrasset

C. Guirado<sup>1,2</sup>, E. Cuevas<sup>2</sup>, V. E. Cachorro<sup>1</sup>, C. Toledano<sup>1</sup>, S. Alonso-Pérez<sup>2,3,4</sup>, J. J. Bustos<sup>2</sup>, S. Basart<sup>5</sup>, P. M. Romero<sup>2</sup>, C. Camino<sup>2</sup>, M. Mimouni<sup>6</sup>, L. Zeudmi<sup>6</sup>, P. Goloub<sup>7</sup>, J. M. Baldasano<sup>5,8</sup>, and A. M. de Frutos<sup>1</sup>





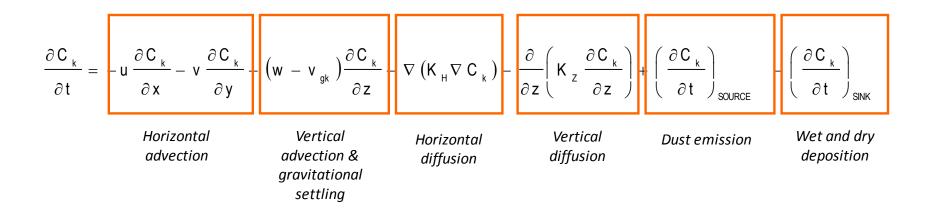
Jan-07 Mar-07 Jun-07 Sep-07 Dec-07 Mar-08 Jun-08 Sep-08 Dec-08

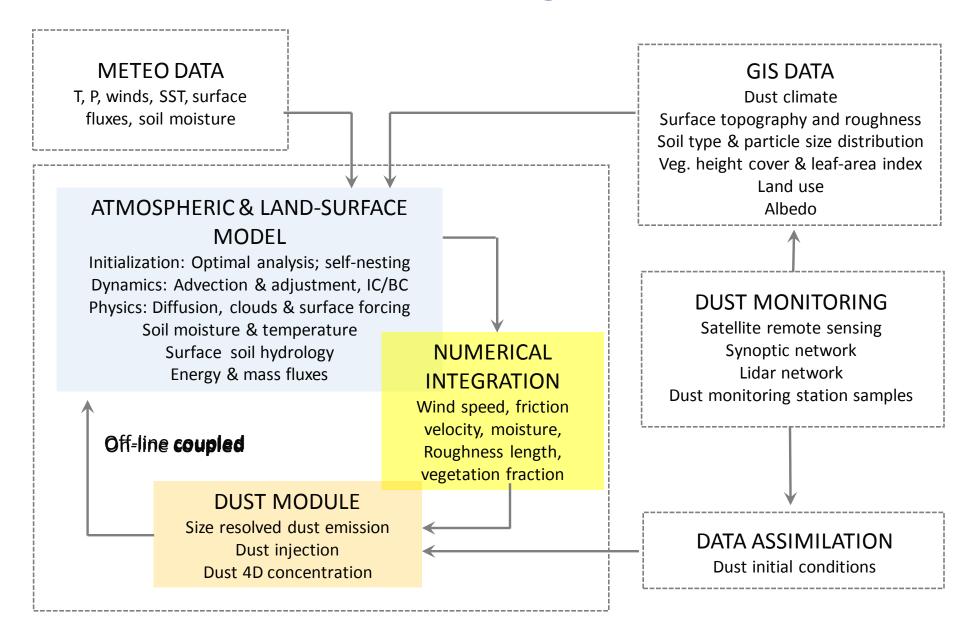


- Dust processes span over five orders of magnitude in space and time. **Dust transport** is a global phenomenon. However, **dust emission** is a threshold phenomenon, sporadic and spatially heterogeneous, that is locally controlled on small spatial and temporal scales.
- To correctly describe and quantify the dust cycle, one needs to understand equally well local-scale processes such as saltation and entrainment of individual dust particles as well as large-scale phenomena such as mid- and long-range transport.

Accurate representation of dust sources and sinks is critical for providing realistic magnitudes and patterns of atmospheric dust fields.

**Dust models** simulate the atmospheric dust cycle and involves a variety of processes:





#### **Desert dust soil types**









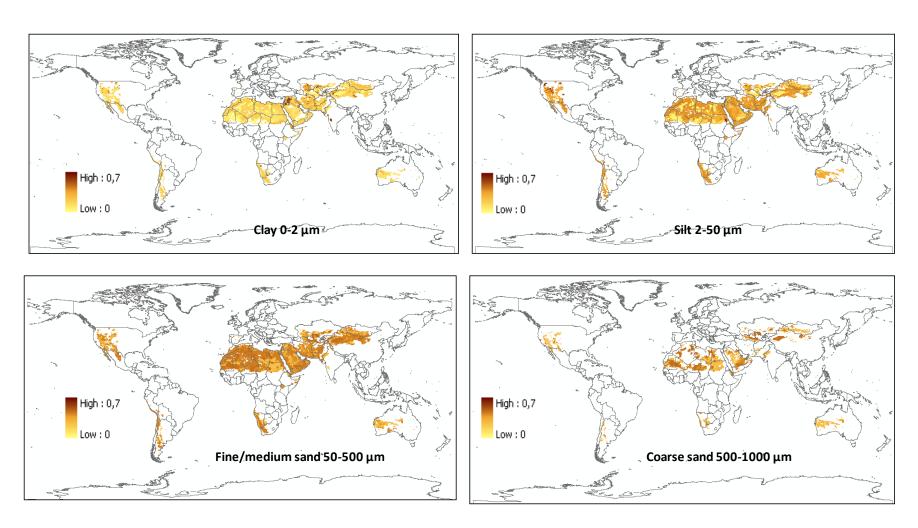




Main landscapes of the North Africa (Photos from Callot et al. 2000):

- A) Central part of Saharan Atlas. In the background, mountains, and in front, an overgrazed plain;
- B) Northern part of Saharan Atlas. Esparto grass steppe degraded by a strong anthropic action. The sandy soil disappears, denuding the sandstone substratum;
- C) The Great Hamada south-west of El-Abiodh-Sidi-Cheikh;
- D) Daïa in the Mechfar, at Hassi Cheikh well;
- E) North-east of the Great Western Erg: coarse sand interdune corridor with deflation cauldron and palaeolake deposits;
- F) North-east of the Great Western Erg: great coarse sand dome dunes, covered by fine sand active dunes.

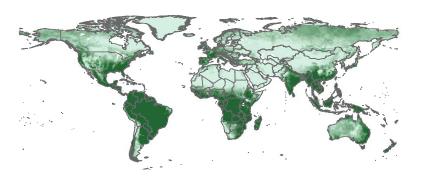
#### Soil size distribution derived from soil texture



Four top soil texture classes according STASGO-FAO 1km database are converted to 4 parent soil size categories following Tegen et al. [2002].

## Vegetation, roughness, soil moisture

Vegetation fraction (MODIS)



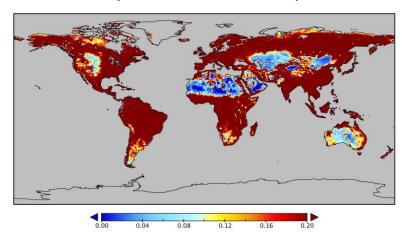
Soil moisture (model based)

90
60
30
0
-30
-60
-60

135

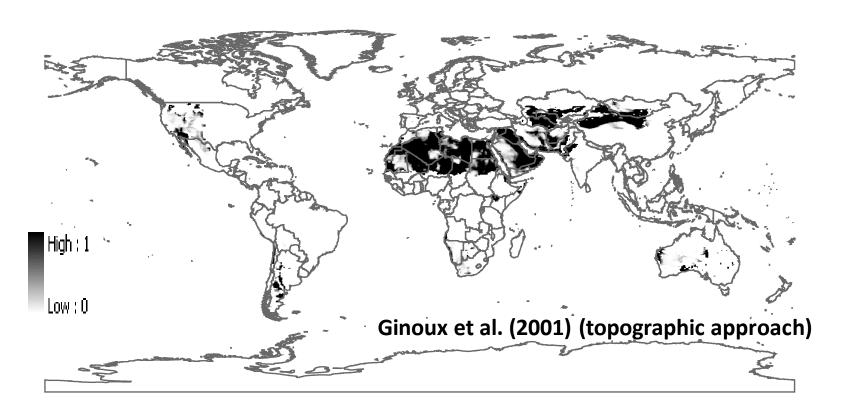
Roughness length (ASCAT + PARASOL)

-135





# Source mapping: why?



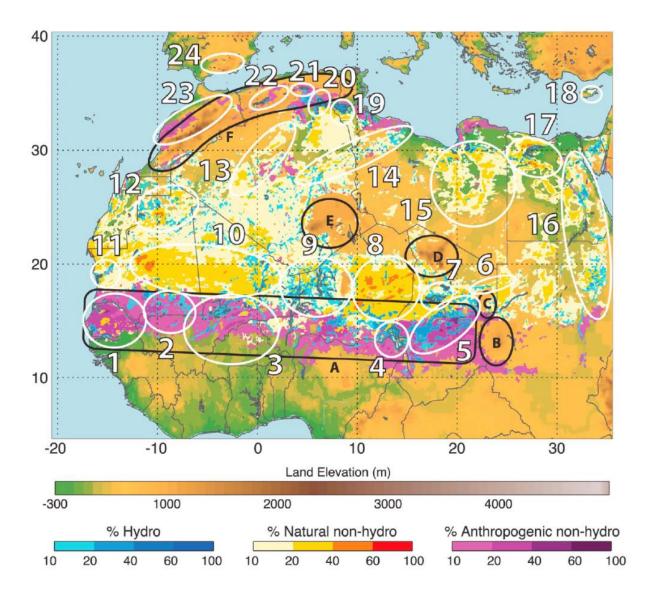
$$S = \left(\frac{z_{\text{max}} - z_{i}}{z_{\text{max}} - z_{\text{min}}}\right)^{5}$$

S: probability to have accumulated sediments in the grid cell i of altitude zi

best fit with the sources identified by Prospero et al. 2000

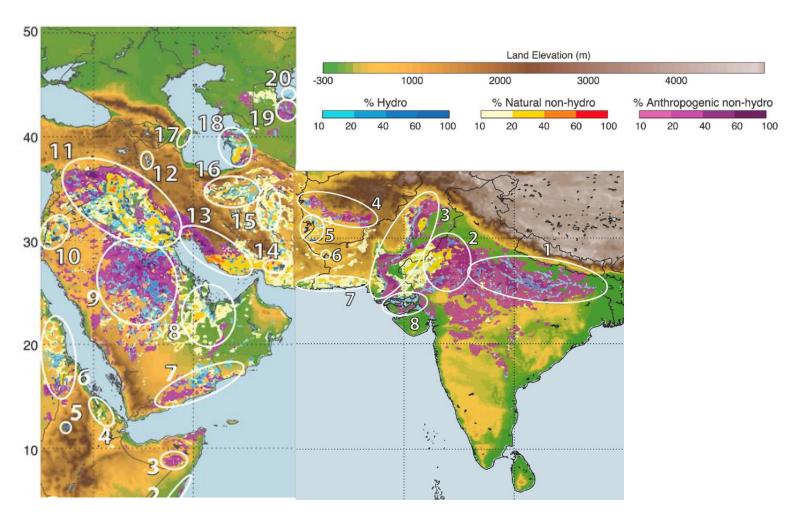


# Natural and anthropogenic dust sources



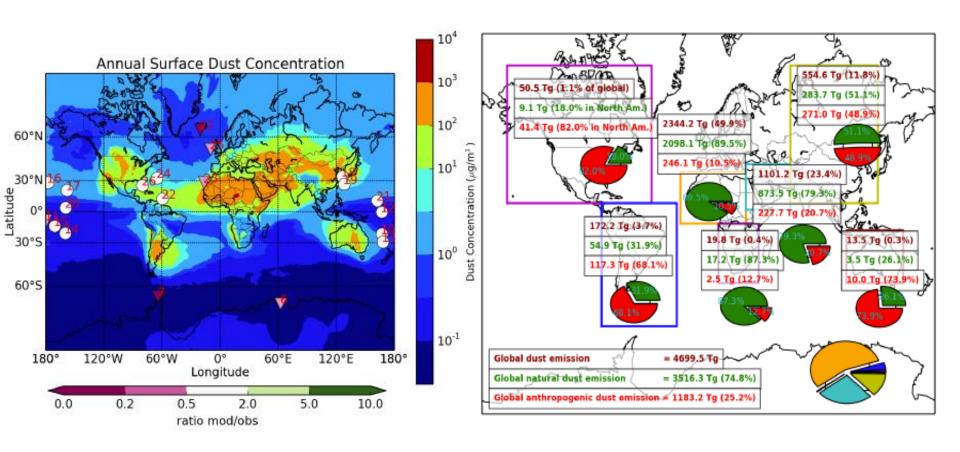


# Natural and anthropogenic dust sources



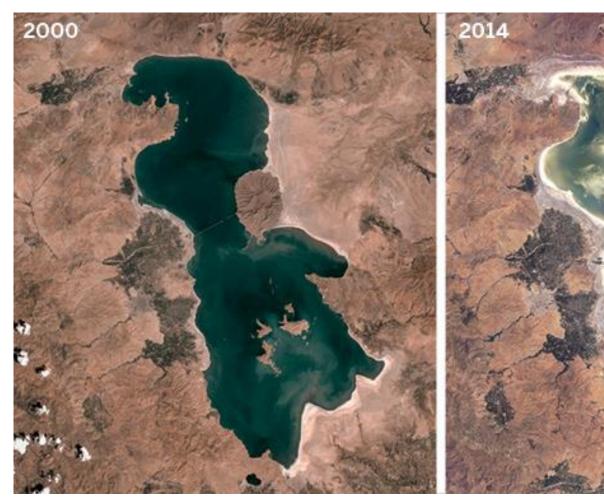


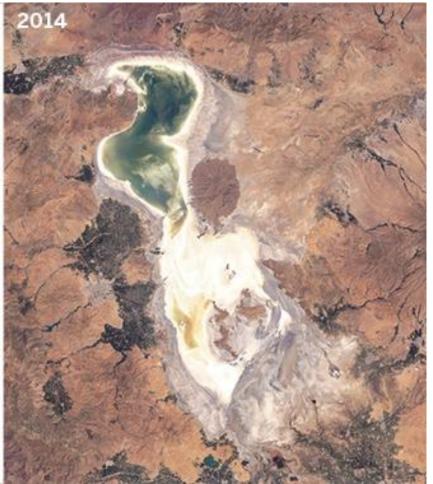
# Current quantification natural vs. anthropogenic





## Major challenge for modeling

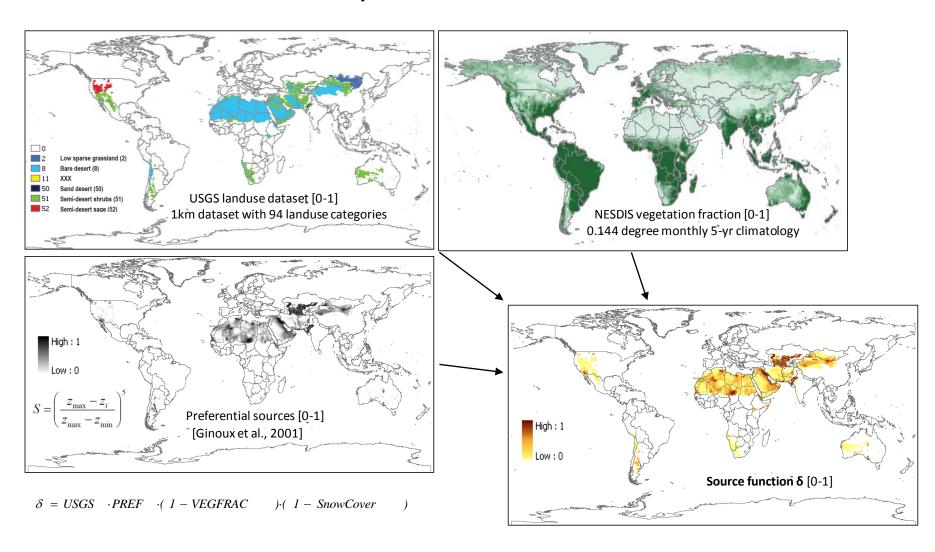






#### **Dust sources functions**

#### **Dust source function: the NMMb/BSC-Dust model**



#### **Dust emission mechanisms**

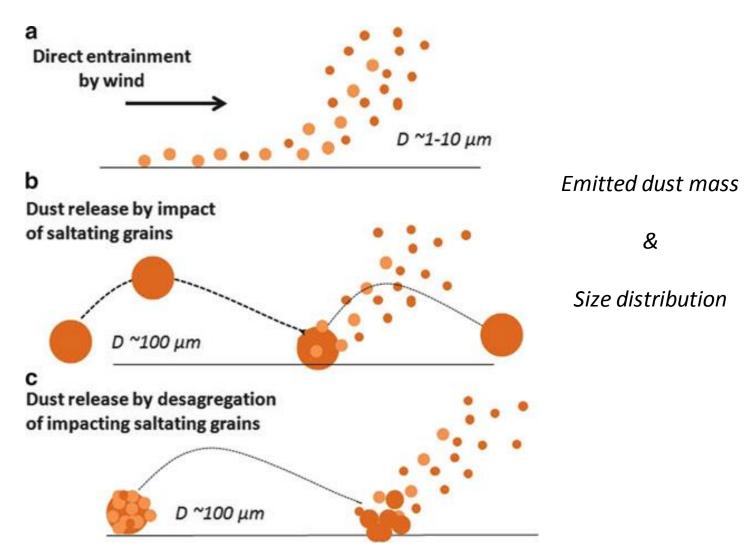
- Complex physical process involving entrainment of soil particles by the surface winds.



- Creep or rolling motion of the largest particles (> 500 um)
- Saltation or horizontal motion of large soil grains (sand) (50-500um)
- Suspension of dust
   (after sandblasting or saltation bombardment)
   (0.1-50 um)

Movie from the COMET program at http://meted.ucar.edu/of the University Corporation for Atmospheric Research (UCAR)

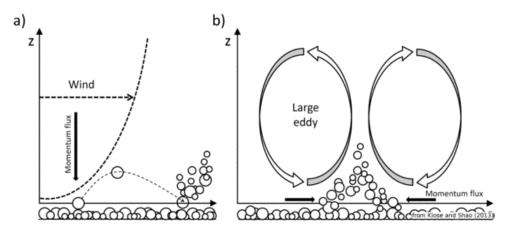
#### **Dust emission mechanisms**





#### Dust storm generation requires:

- High wind
- Wind shear and turbulence
- Unstable boundary layer



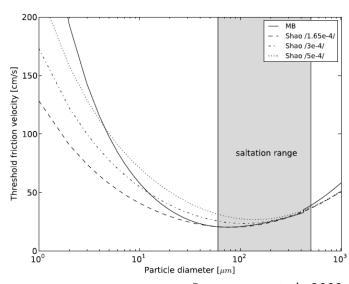
(a) Conventional view of dust emission via saltation bombardment; (b) Illustration of convective turbulent dust emission.

Extracted from Shao (2008)

**Friction velocity** is as the velocity above which soil particles begin to move in saltation flux.

- Depends on soil grain size, soil moisture and roughness among others

Threshold friction velocity vs particle radius  $\rightarrow$ 



Darmenova et al., 2009

#### Simple schemes

Formulation of vertical dust flux (F)

$$F = c \cdot f \cdot P(u_*^n, u_{*_{th}}) \qquad \text{if} \quad u_* > u_{*_t}$$

c: dimensional scale dependent constant proportinality

f: relative surface area of each soil particle fraction (which includes de source function,  $\delta$ )

 $u_*$ : friction velocity

 $u_{*t}$ : threshold friction velocity

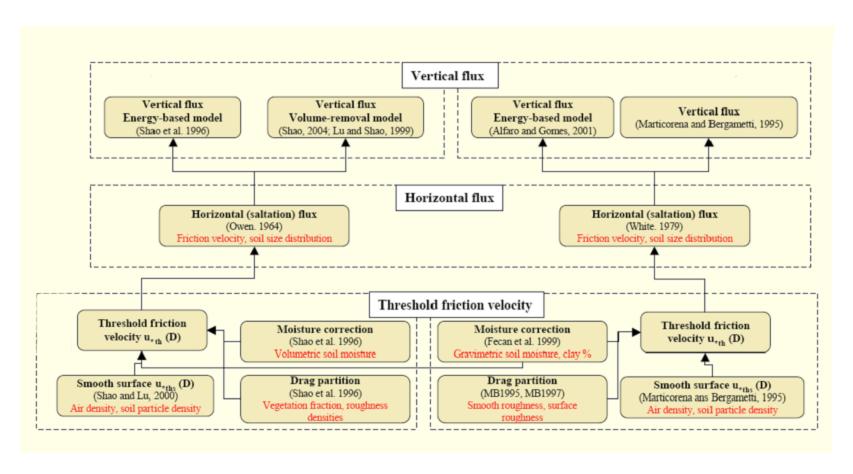
P: polinomial of degree n

Study	Scheme
Uno et al. (2001) CFORS	$F = cu_{10}^2(u_{10} - u_{10t})$
Liu and Westphal (2001) COAMPS	$F = f u_{10}^{2} (u_{10} - u_{10t})$
Liu and Westphal (2001) COAMPS	$F = fcu_*^4$

#### Simple schemes

#### Limitations

- Oversimplified physical representation of dust emission.
- Normalization constant C is not known
- Erodible fraction is prescribed for predefined dust sources
- Threshold friction velocity is usually a fixed value (no dependence on the land surface properties)
- Assuming constant threshold friction velocity will introduce bias in the modelling of the timing and intensity of dust events.
- The prescribed constant is model dependent and can result in large discrepancies in calculated dust loadings between different models.



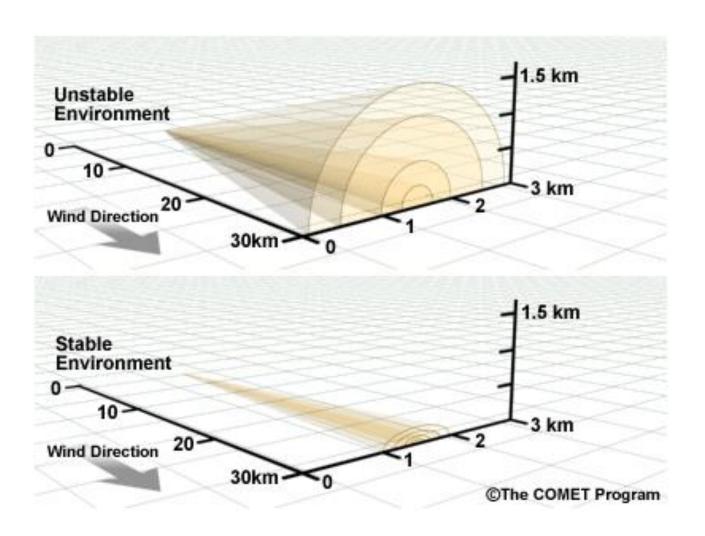
Physically-based **dust emission schemes** employ different parameterizations of the related physical processes, as well as require different input data.

# **Dispersion**

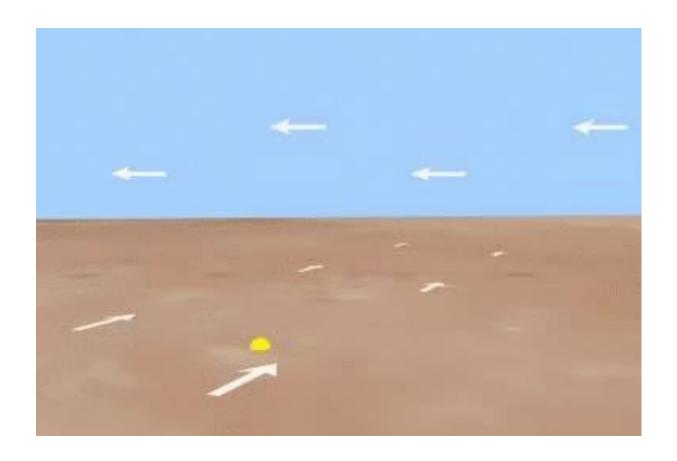


Movie from the COMET program at http://meted.ucar.edu/ of the University Corporation for Atmospheric Research (UCAR)

### **Dispersion**



### **Advection and diffusion**



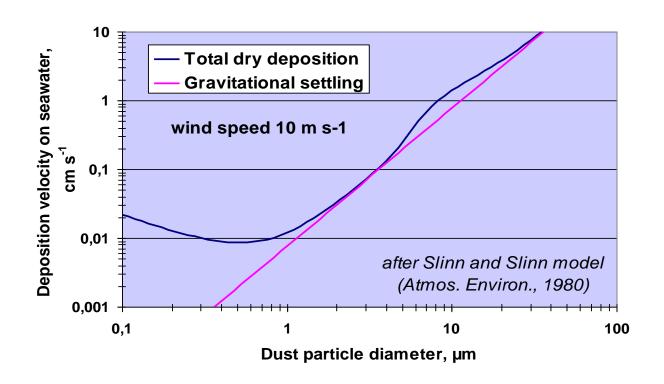
Movie from the COMET program at http://meted.ucar.edu/ of the University Corporation for Atmospheric Research (UCAR)

### Sedimentation and dry deposition



Movie from the COMET program at http://meted.ucar.edu/ of the University Corporation for Atmospheric Research (UCAR)

### Sedimentation and dry deposition



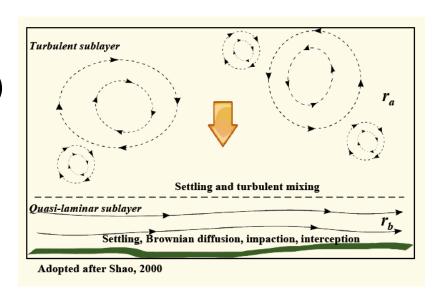
- **Dry deposition** depends on the variety of factors such as meteorological conditions near the surface, physicochemical properties of mineral dust and the nature of the surface itself.
- Sedimentation (or gravitational settling) is the settling of particles fall down due to gravity → Very large particles will settle out quickly

### **Dry deposition**

**Dry deposition velocity** is represented as 3 resistances in series parallel to a second pathway - gravitational settling velocity:

$$v_d = \frac{1}{r_a + r_b + r_c} \longrightarrow F_d = -C \cdot v_d$$

- Aerodynamic resistance to transfer  $(r_a)$
- Quasi-laminar surface layer resistance  $(r_h)$
- Resistance to surface uptake  $(r_c)$



### Wet deposition

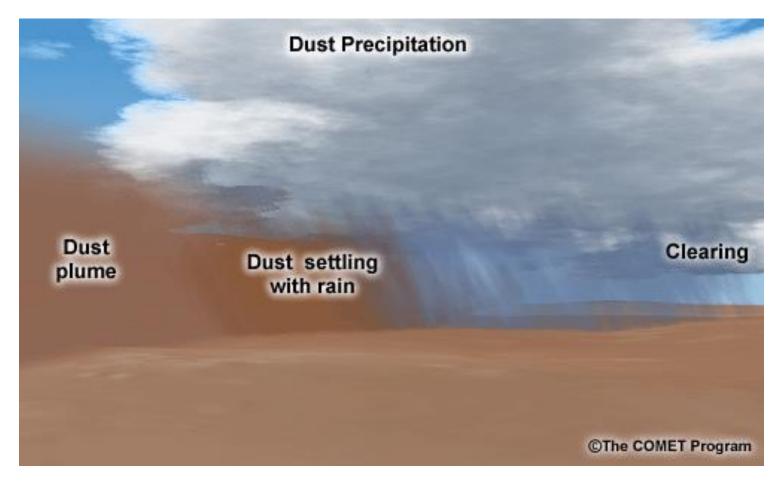
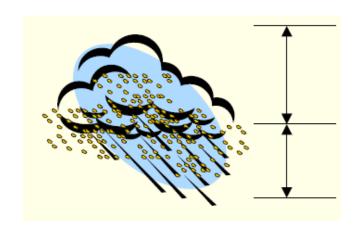


Image from the COMET program at http://meted.ucar.edu/of the University Corporation for Atmospheric Research (UCAR)

### Wet deposition



#### In-cloud scavenging:

- **nucleation scavenging** by activation and growth of particles to cloud droplets
- **collection** of a non-activated fraction of particles by coagulation with cloud and rain droplets

#### **Below-cloud scavenging:**

Collection by falling raindrops of particles under their collision.

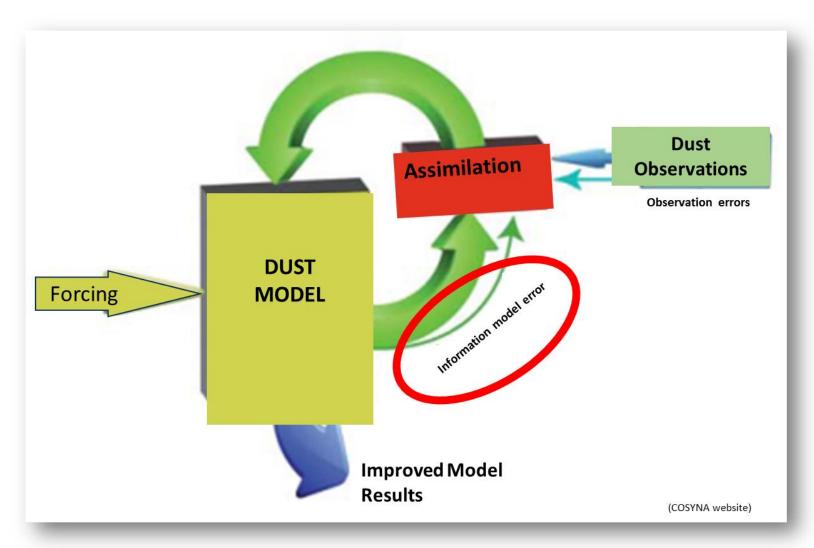
Decrease rate of the aerosol concentration due to **wet scavenging** in a layer with uniform concentration can be described by a first-order equation:

$$\frac{\partial C}{\partial t} = -\lambda C$$

#### The **scavenging coefficient (C)** depends on:

- the particle size and solubility
- the collectors size distribution and fall speeds
- precipitation rate and phase (rain or snow).

#### **Data Assimilation**



Obtaining the 'best' estimate of current atmospheric dust conditions (analysis)
Creating datasets describing the recent history of dust in the atmosphere (reanalysis)

### **Dust forecasting models**

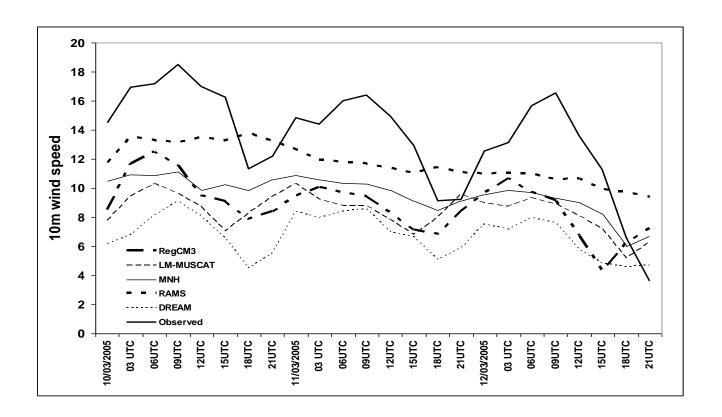
#### Main differences between dust models

- 1. Meteorological driver
- 2. Meteorological input files IBC
- 3. Emission scheme
- 4. Geographic-information database (source mask)
- 5. Land-surface scheme
- 6. Dry deposition scheme
- 7. Wet depositioon scheme
- 8. Spatio-temporal resolution
- 9. Data assimilation
- 10. ....

### **Dust forecasting models**

**Experimental campaigns: BODEX 2005 (Todd et al. 2008, JGR)** 

First regional model intercomparison in the Bodélé hot spot



Strong differences between models!!!! → Meteorology and emission scheme





















# Thank you

Acknowlegde to Carlos Pérez García-Pando, Emilio Cuevas, Slodoban Nickovic, Francesco Benincasa, Enza DiTomaso, Oriol Jorba, Kim Serradell, Enric Terradellas as well as AERONET, MODIS, U.K. Met Office MSG, MSG Eumetsat and EOSDIS World Viewer principal investigators and scientists for establishing and maintaining data used in the present contribution. Also special thank to all researchers, data providers and collaborators of the WMO SDS-WAS NA-ME-E Regional Node.

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