Impacts of the increased horizontal resolution on the predictability of the tropical Pacific variability

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Background

Seasonal to Decadal climate prediction (S2DCP) is a field of research attracting growing interest beyond the scientific community due to its strong potential to guide decision-making in many sectors (e.g. energy production, agriculture) in the face of the pressing dangers of climate change (*Hermanson et al., 2022*).

However, one major limitation common to current S2D systems is the little skill that they present over the continents, which appears to be connected to an incorrect representation of the teleconnection mechanisms that, mediated via the atmosphere, connect the ocean with the neighbouring continents.



Atmospheric forcing (assimilation scheme) Applied to the ocean-sea ice surface Variables: t10, u10, v10, precip., q10, qlw, qsw **Ocean surface restoring (ORAS5)** Applied to the surface with fixed strength Variables: temperature and salinity

Sub-surface 3D-nudging (EN4v4.2)

Applied to the sub-surface with varying strength in depth and weak nudging in the equatorial band Variables: temperature and salinity



There are several indications that the current generation of models at standard resolution misrepresents those key teleconnections, and that **higher resolution** versions might improve them, decreasing common biases of global models and improving the skill to predict certain regions at seasonal scales, e.g. in tropical sea surface temperature (*Prodhomme et al. 2016*).

We want to assess how the predictive skills change when the horizontal resolution is **increased** and how these changes are related to the simulation of physical processes. The comparison of **two different Climate models** allow to better understand how the development of model-specific biases can impact predictive skills.

FIG. 1. Schematic summarizing our protocol to produce in-houses ocean reconstruction, used as initial conditions for the prediction systems. Courtesy Alba Santos Espeso.

		Atmospheric initial cond.	Ocean initial cond.	
CNRM-CM6.1	SR	ERA-interim	Glorys 2v4	
	HR	ERA5	Glorys 12v1	
EC-Earth3.3	SR	ERA5 interpolated to the	In-house reconstruction using NEMO with the corresponding grid	
	HR	corresponding IFS grid		

Experimental approach

We used two different models, EC-Earth3 (Döscher el al., 2021) and CNRM-CM6.1 (Voldoire et al., 2019), in their Atmosphere-Ocean configuration. We compare their configuration at two different resolutions, standard (SR) and high (HR) resolution. We performed two sets of seasonal retrospective hindcasts with each resolution configuration of the two models.

		Atmospheric component	Oceanic-sea ice component		hindcast systems	
				Hindcasts	EC-Earth3	CNRM-CM6.1
CNRM-CM6.1	IVIODEI	ARPEGE V6	NEIVIO3.6 - GELATO V6	every year	1990-2015	1993-2014
	Standard resolution	~130 km (TL127)	~100 km (eORCA1)	20-member prediction started 1 May 1994 20-member prediction started	Twice a year	Twice a year
	High resolution	~50 km (TL359)	~25 km (eORCA025)	20-member prediction started 1 May 2014	May and November	May and November
EC-Earth3.3	Model	IFS cy36r4	NEMO3.6-LIM3	1 May 1993	20 members	30 members
					8 forecasted months	6 forecasted months
	Standard resolution	~80 km (TL255)	~100 km (ORCA1)		TOTAL: 693 eq. years	TOTAL: 660 eq. years
	High resolution	~40 km (TL511)	~25 km (ORCA025)	1993 Observations		
TAB 2 Model compo	nents and corresponding grid	and resolution for the two models u	sod CNRM_CM6.1 and EC_Earth2	EIG 2 Schematic of a prodiction system simulation structure		





Next steps

- Analysis of the physical processes in the tropical band that can explain this increase in skill:
 - Surface winds
 - Thermocline dynamics
 - Bjerknes Feedback
 - Equatorial Atlantic teleconnections
 - Indian Ocean teleconnections
- Analysis of the differences in the mean bias and their change with resolution for each model
- Investigate the initialised November forecast system
- the EC-Earth3 forecast Extension of system over a longer hindcast period

FIG.3. Maps of Anomaly Correlation Coefficient (ACC) for the first forecasted season (June-July-August) of sea surface temperatures (SST) for (top) EC-Earth3 and (bottom) CNRM-CM6.1 forecast system initialised in May. The left column corresponds to the HR version of each model, the middle one to the SR version and the right column shows the difference in correlation between HR and SR. In this right column, red indicates an improvement in the skill with increased resolution, blue indicates a degradation of the skill. The reference dataset is the ESA SST CCI dataset. The ACC has been computed over the 1993-2014 period for each individual grid point. All the models and reference dataset have been interpolated to a regular grid of 1º of resolution before computing the ACC. Dashes indicate that the values are statistically significant at the 95% level.



FIG.4. ACC skill scores for the (left) Indian Ocean dipole (IOD) and (right) Niño3.4 SST indexes a function of forecasted month for (black) EC-Earth3-SR, (blue) EC-Earth3-HR, (dark green) CNRM-CM6.1-SR and (light green) CNRM-CM6.1-HR forecast systems initialised in May. The reference datasets are the ESA SST CCI dataset and ERA5 reanalysis.

Both models show an overall increase in their predictive skill of the tropical Pacific band sea surface temperature (SST) (Fig.3) and air surface temperature (not shown) with increased horizontal resolution.

The increase in the skill is confirmed when assessing integrated value over specific regions, such as the Indian Ocean and equatorial regions associated with the El Niño-Southern Oscillation (ENSO), such as Niño3.4 (Fig. 4).

(1960-2021) and for a longer forecasted length (up to 3 years) to evaluate how the increased resolution impacts the skill in the tropical Pacific: prediction of ENSO beyond the Spring predictability barrier, prediction of double-dip La Niña (*Wu et al., 2021*), etc.

References

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