



Coastal permafrost erosion & the future Arctic Ocean's CO₂ uptake

David Nielsen, Fatemeh Chegini, Joeran Maerz, Armineh Barkhordarian, Sebastian Brune, Paul Overduin, Patrick Pieper, Mikhail Dobrynin, Victor Brovkin, Johanna Baehr, Tatiana Ilyina

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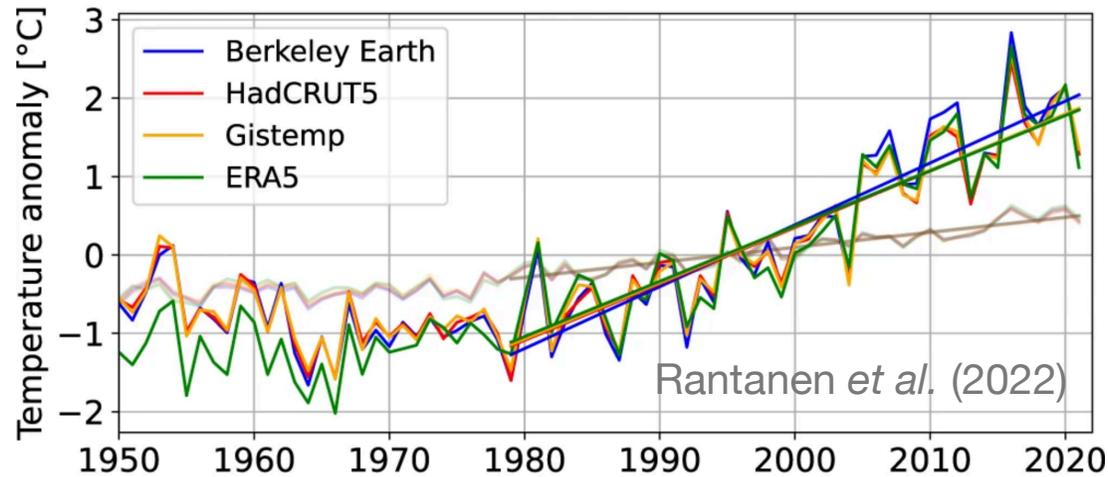


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- Abrupt permafrost thaw contributes to warming via the **permafrost-climate feedback**
(Turetsky et al. 2019, 2020)



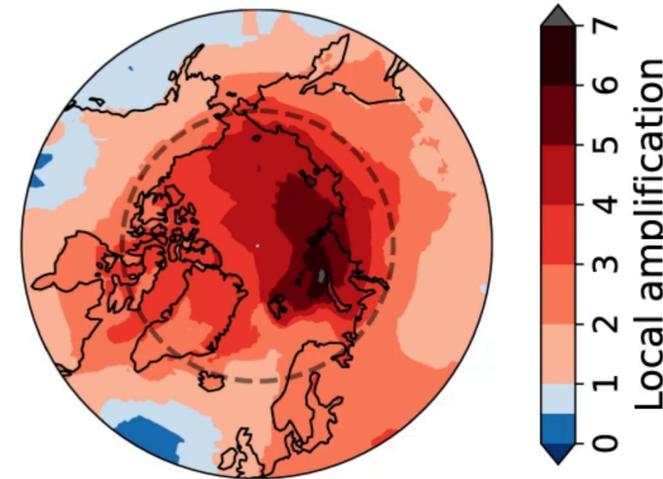
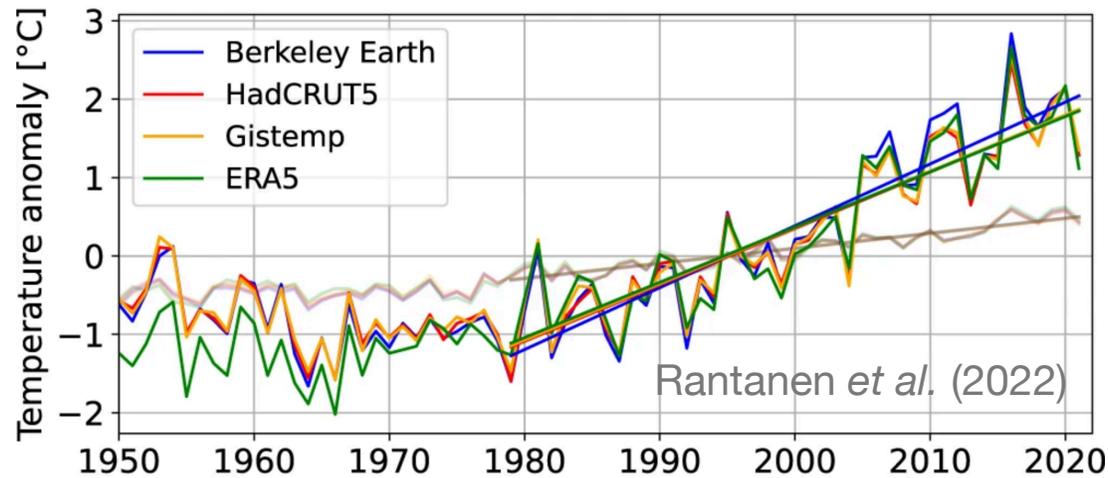
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Rantanen et al. (2022), Chylek et al. (2022)

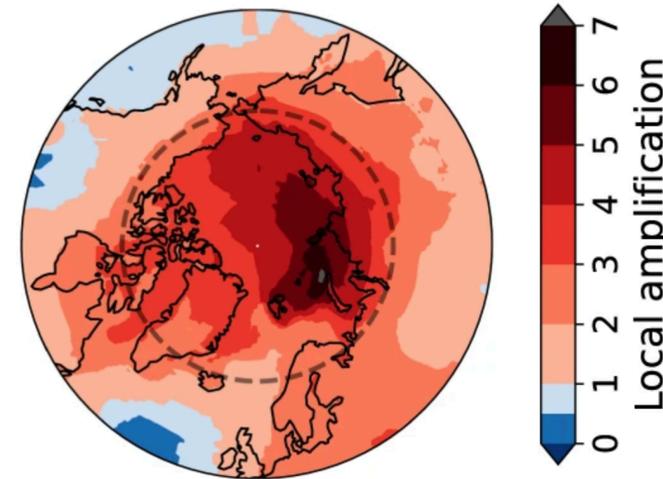
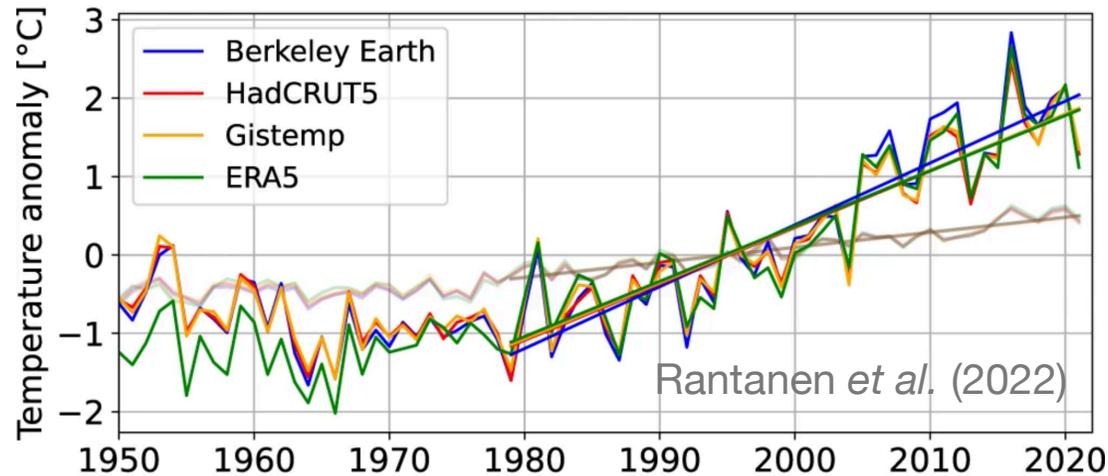
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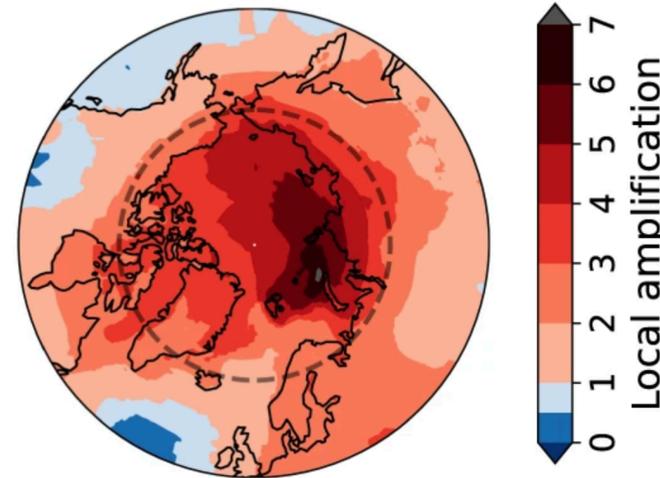
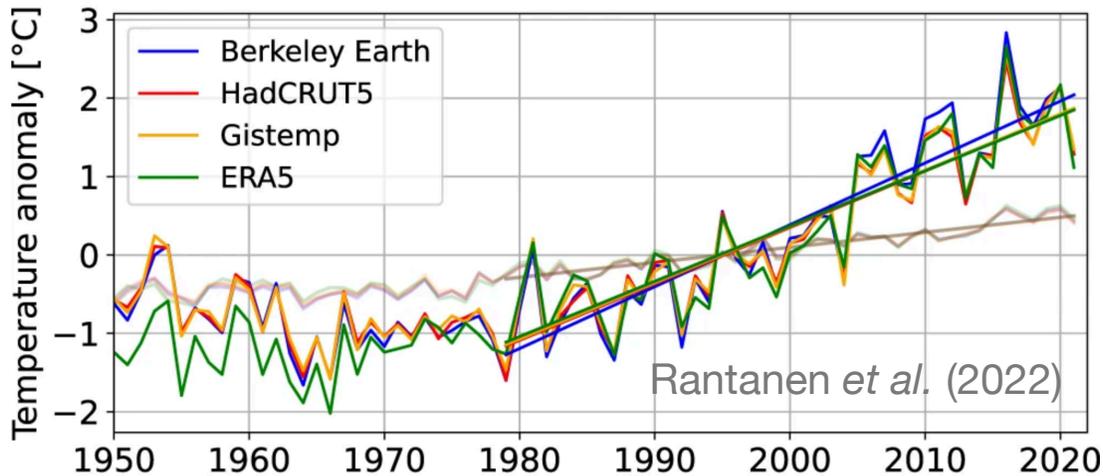
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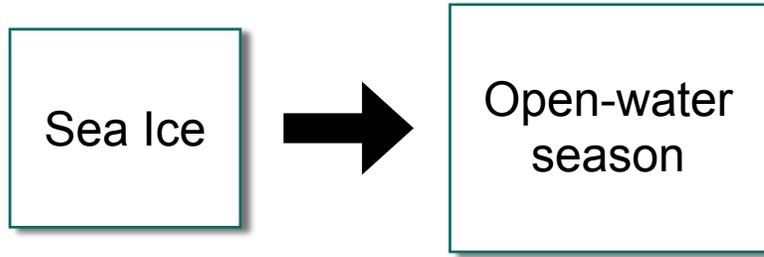
How will Arctic coastal permafrost erosion change in the future?

Drivers of coastal permafrost erosion

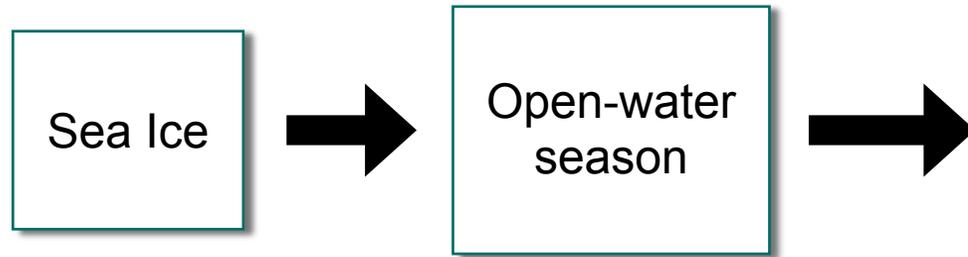


Sea Ice

Drivers of coastal permafrost erosion



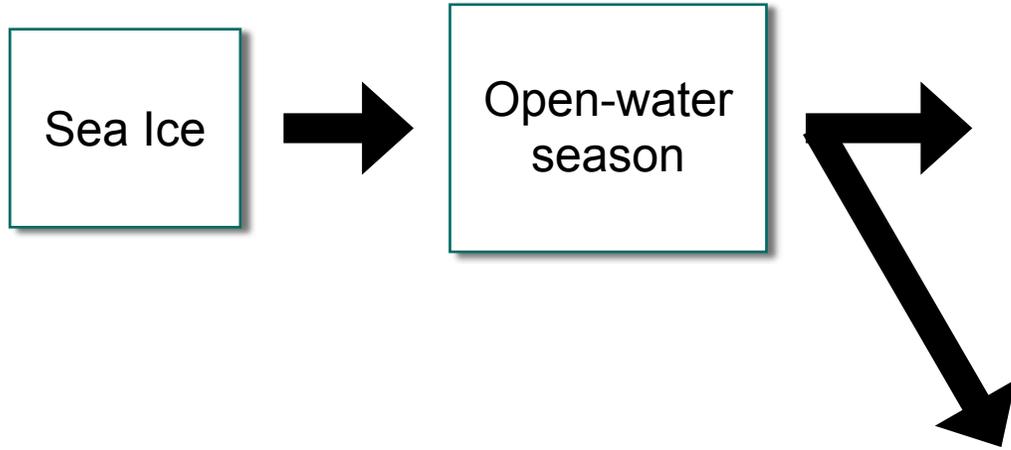
Drivers of coastal permafrost erosion



Thermo-denudation



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Thermo-abrasion



Adapted from Günther et al. (2015) in “*Observing Muostakh disappear*”

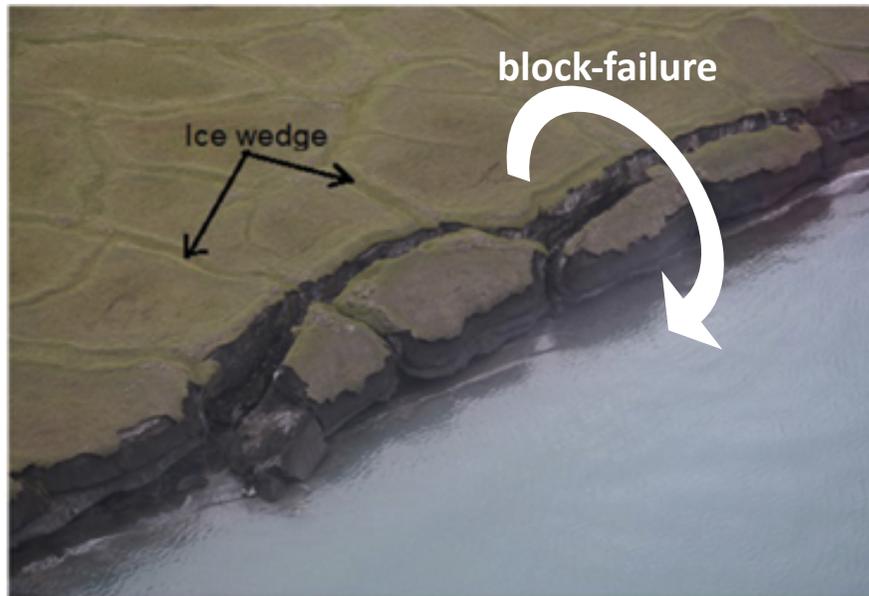
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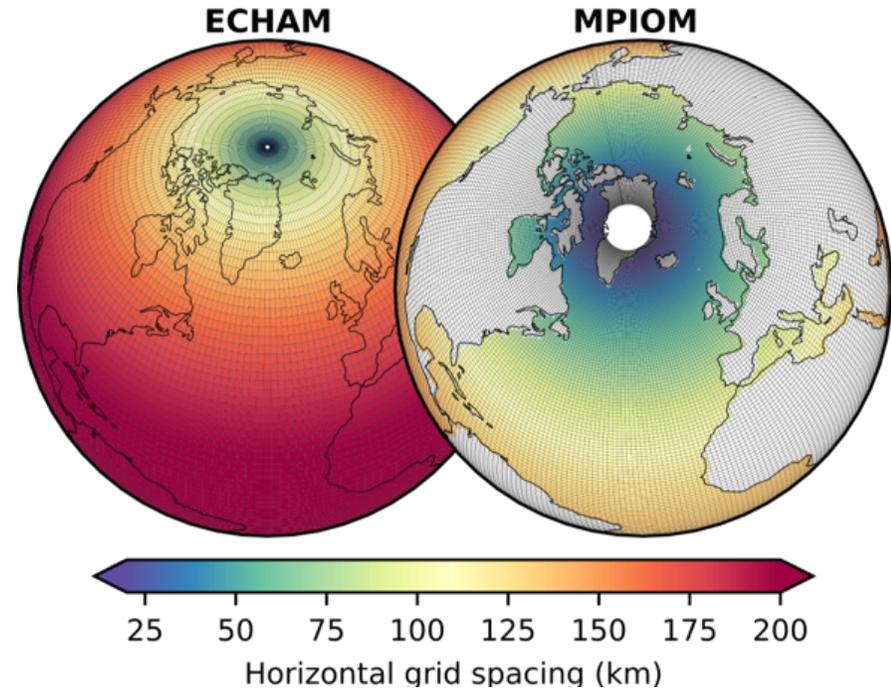
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There is a scale gap

between Arctic coastal erosion and Earth system models (ESMs)



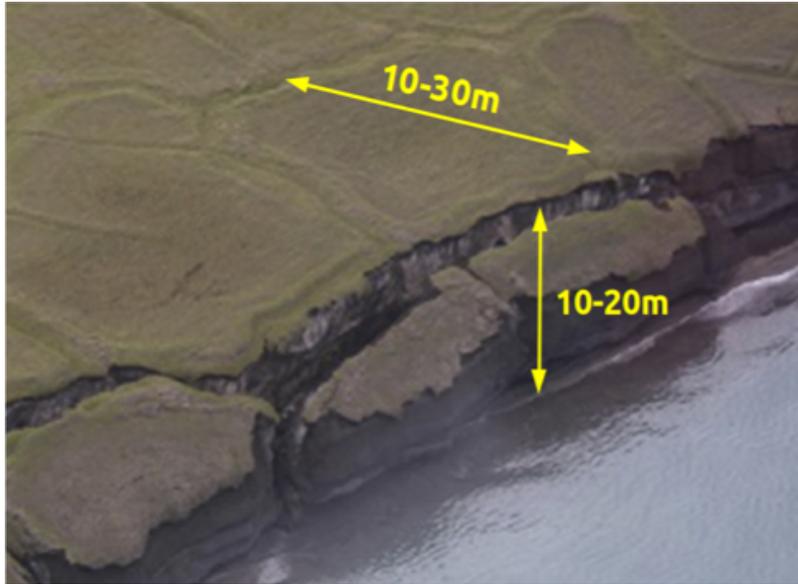
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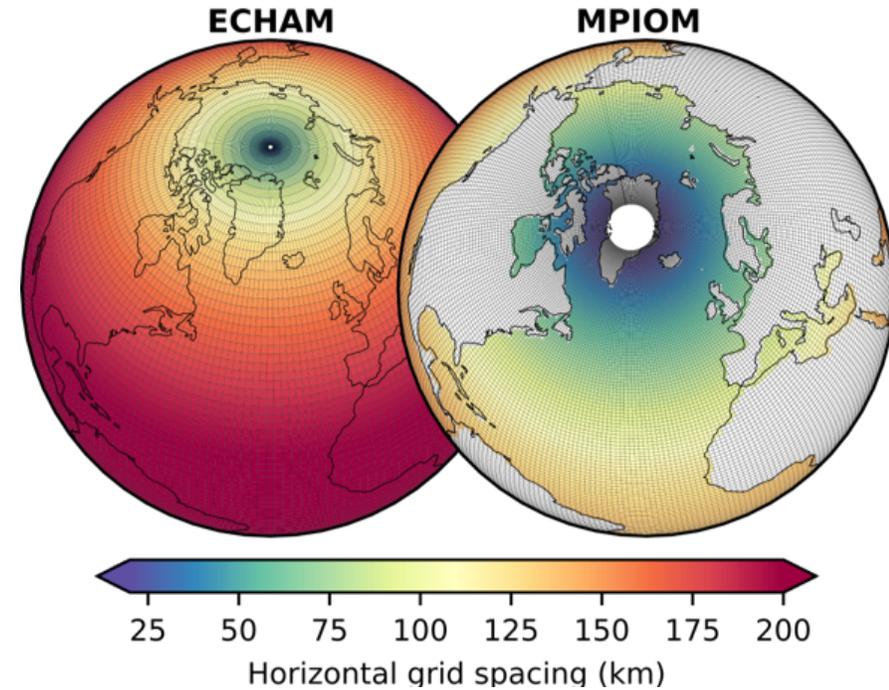
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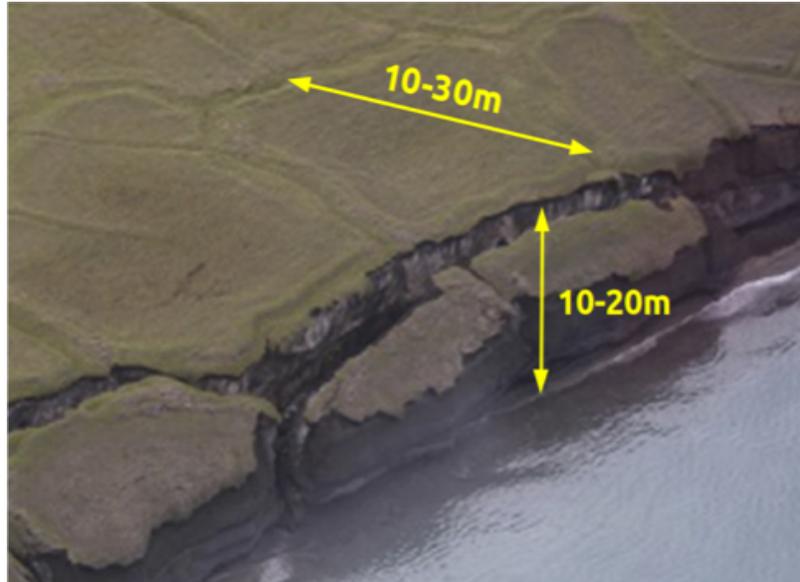


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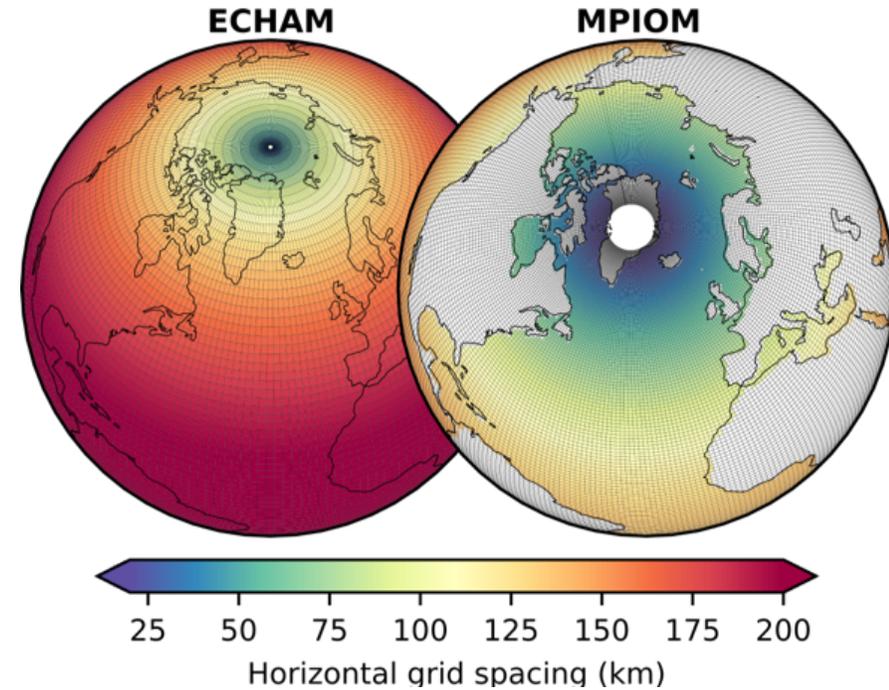
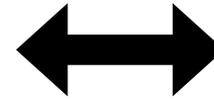
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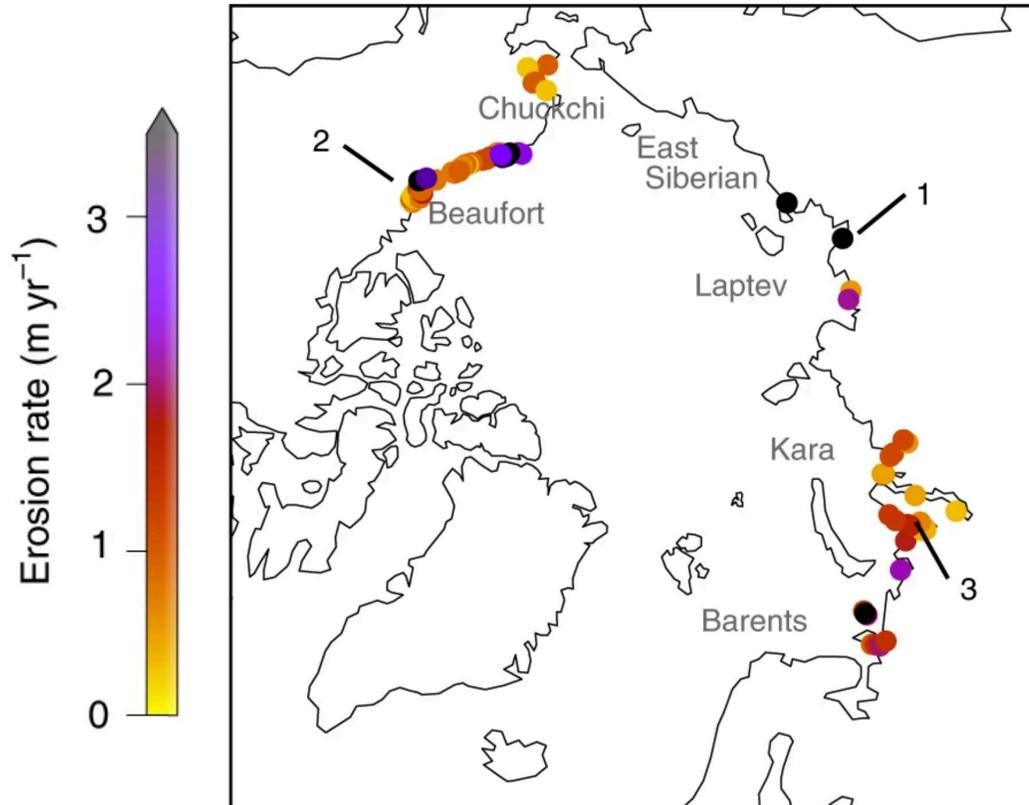
- Existing erosion models are **not compatible with ESMs** used for climate projections (e.g. CMIP6)
- Future projections of **permafrost-carbon loss** could be largely underestimated
- **New modelling frameworks are needed to close the scale gap**

Turetsky et al. (2019, 2020), Fritz et al. (2017)

A simple, semi-empirical model for Arctic coastal erosion



Observations from the Arctic Coastal Dynamics (ACD) database (Lantuit et al. 2012)



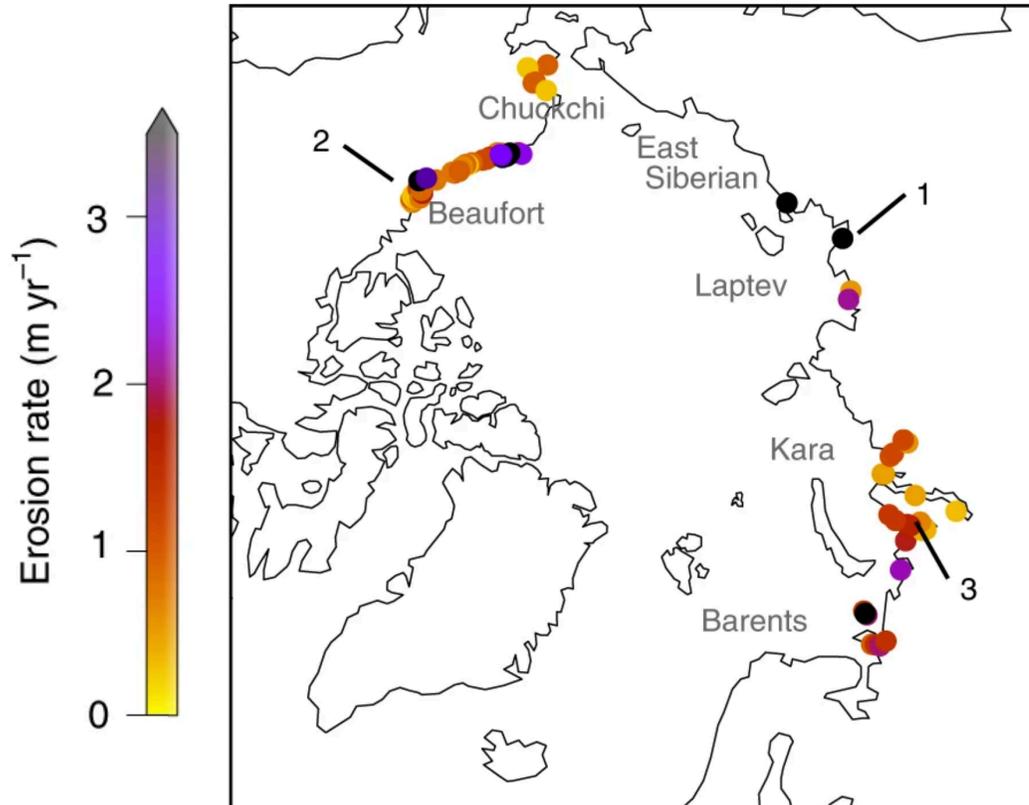
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$$E(x, t) = \overline{E}(t) + \Delta E(x, t)$$

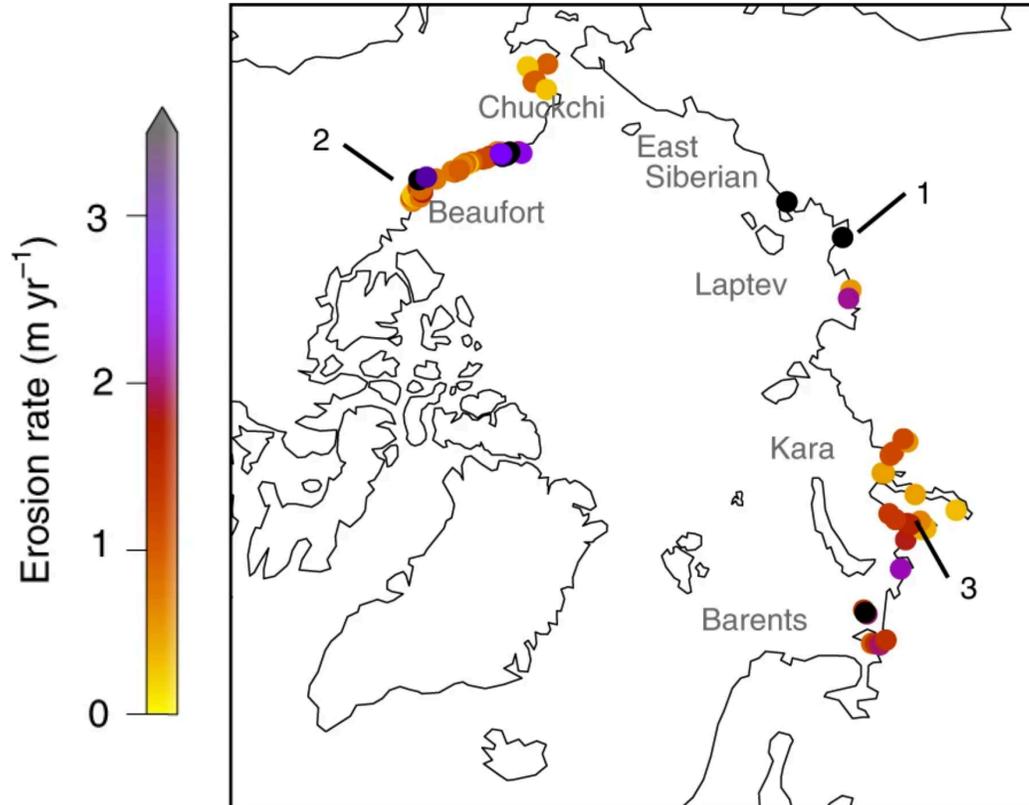
Temporal

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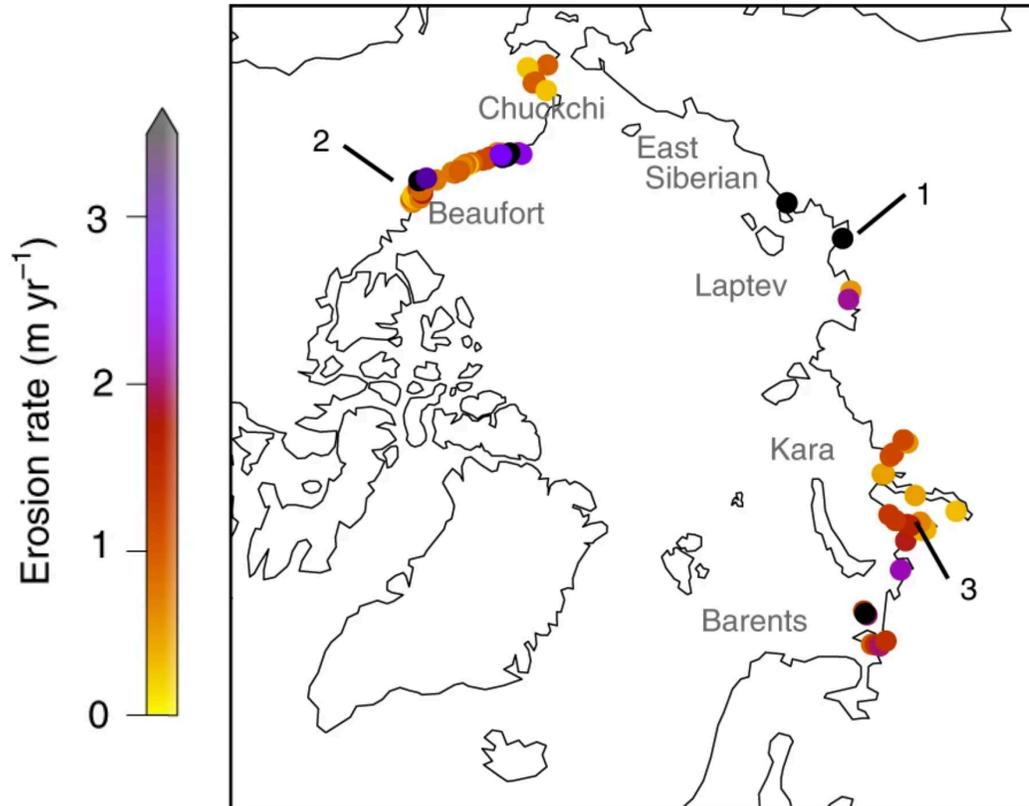
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- Significant wave heights
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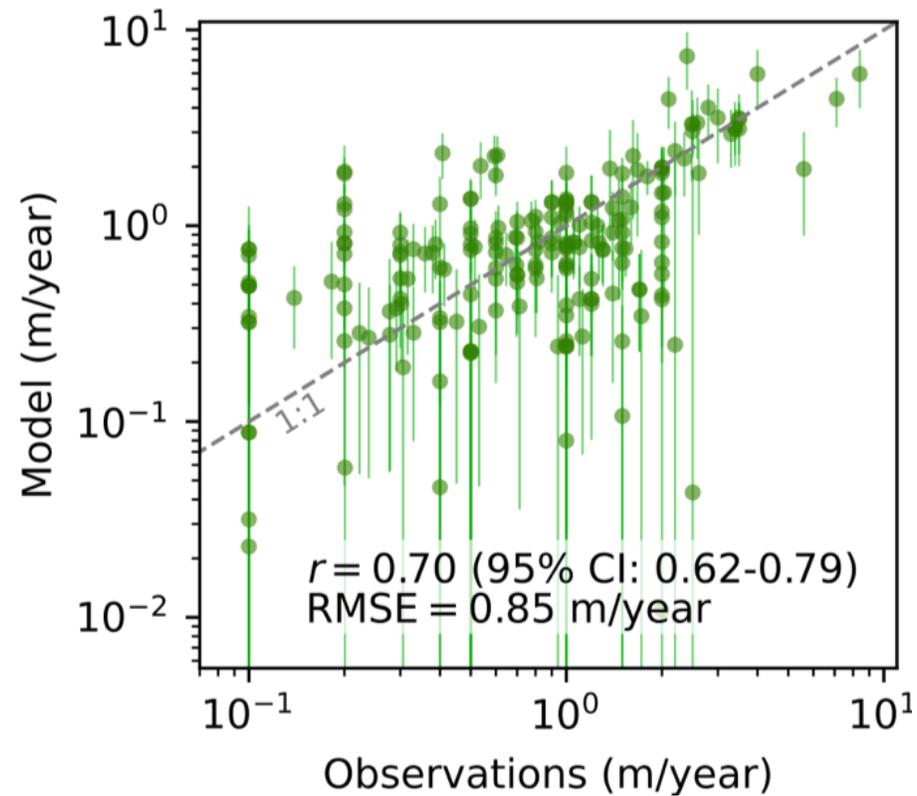
Local daily-accumulated yearly:

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- Positive surface air temperature
- Ground-ice content

Pan-Arctic spatial variability



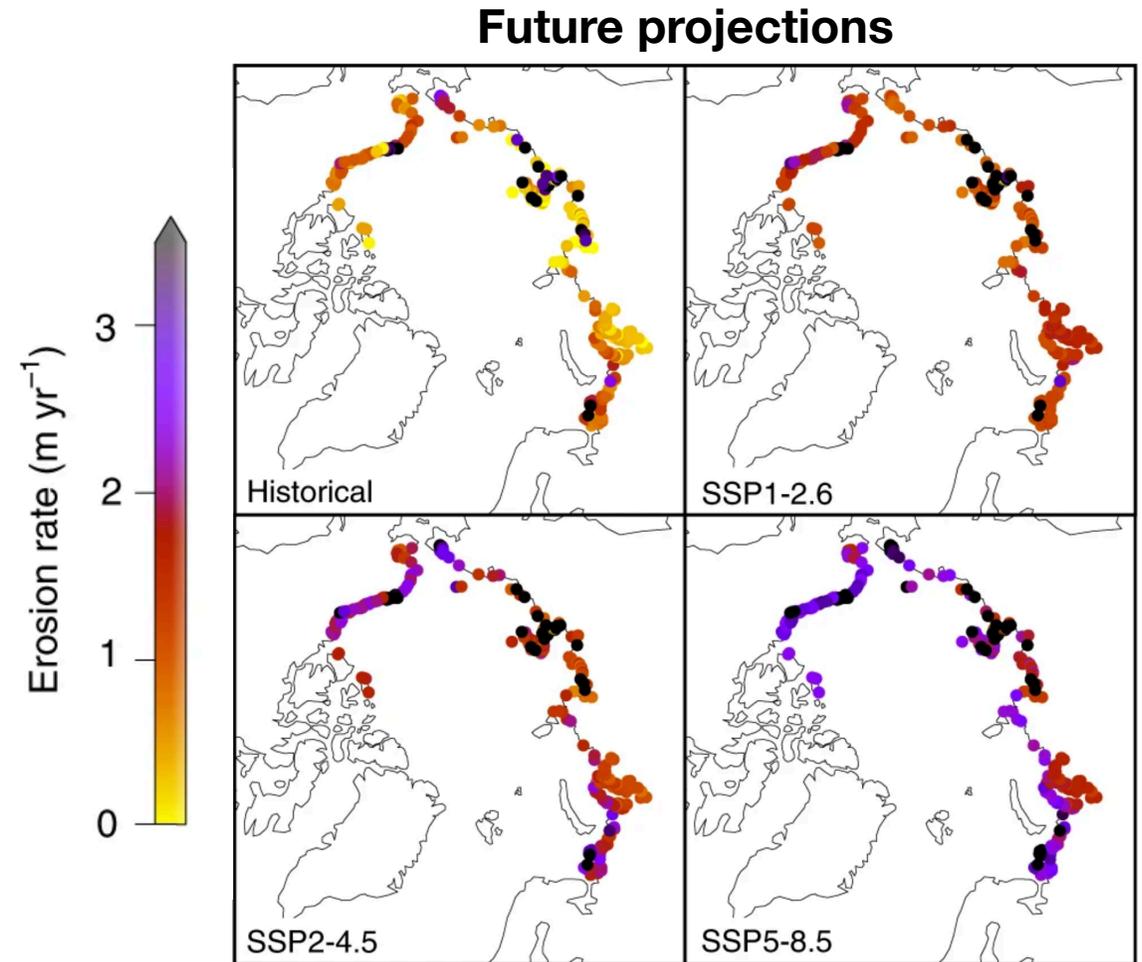
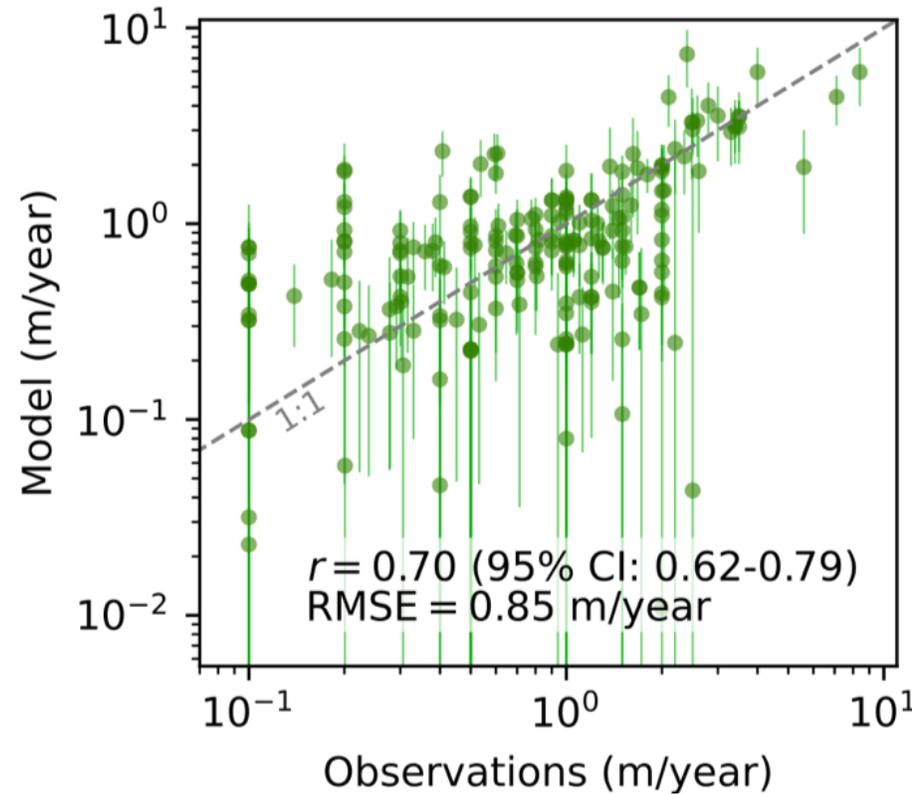
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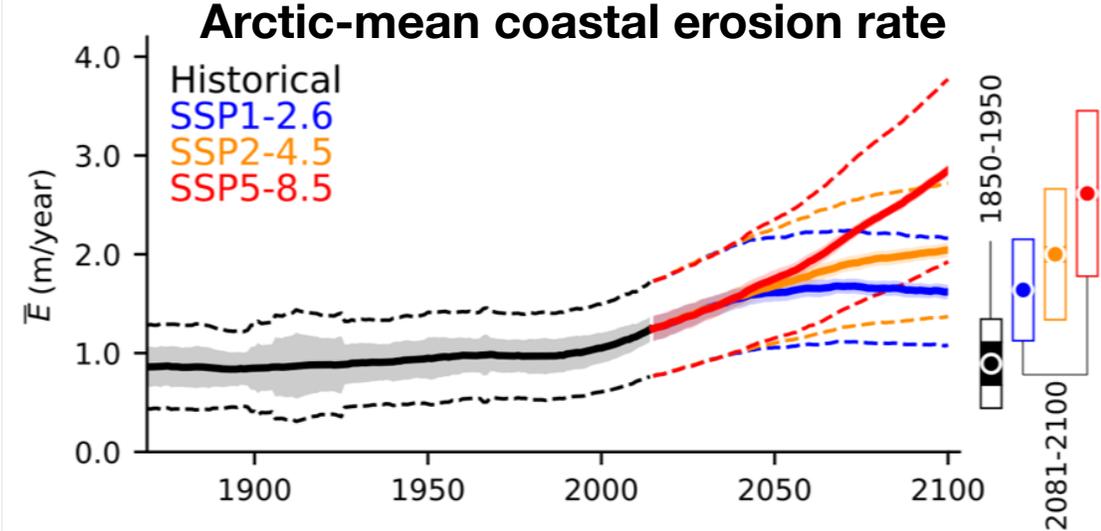


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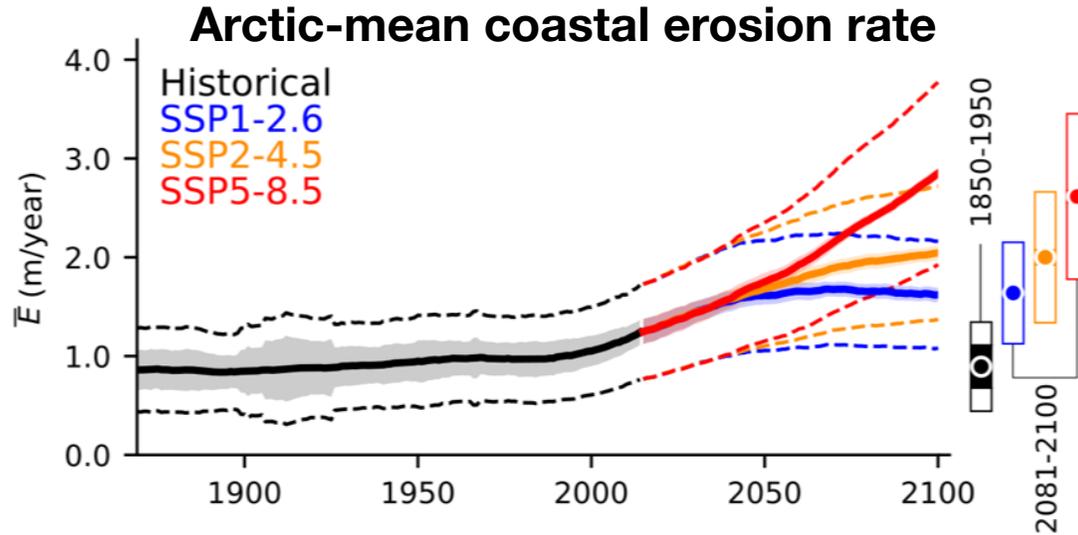


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Historical and 21st-century projections



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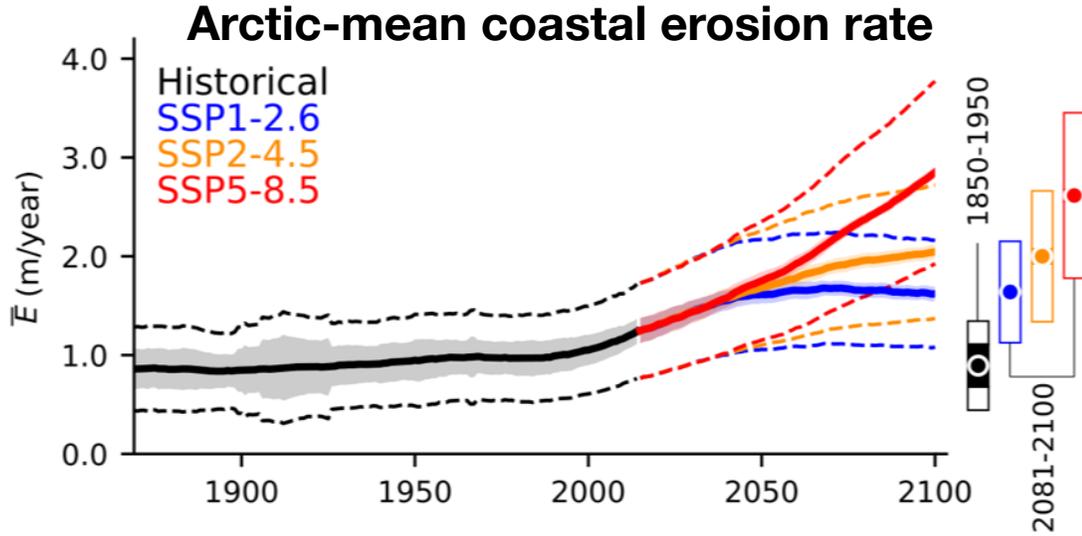


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x2.1 x2.9

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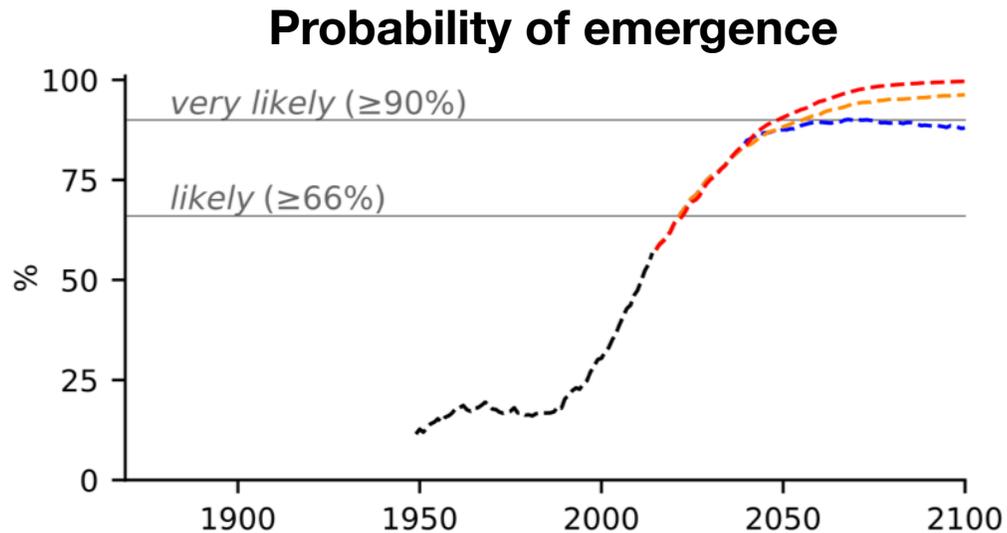
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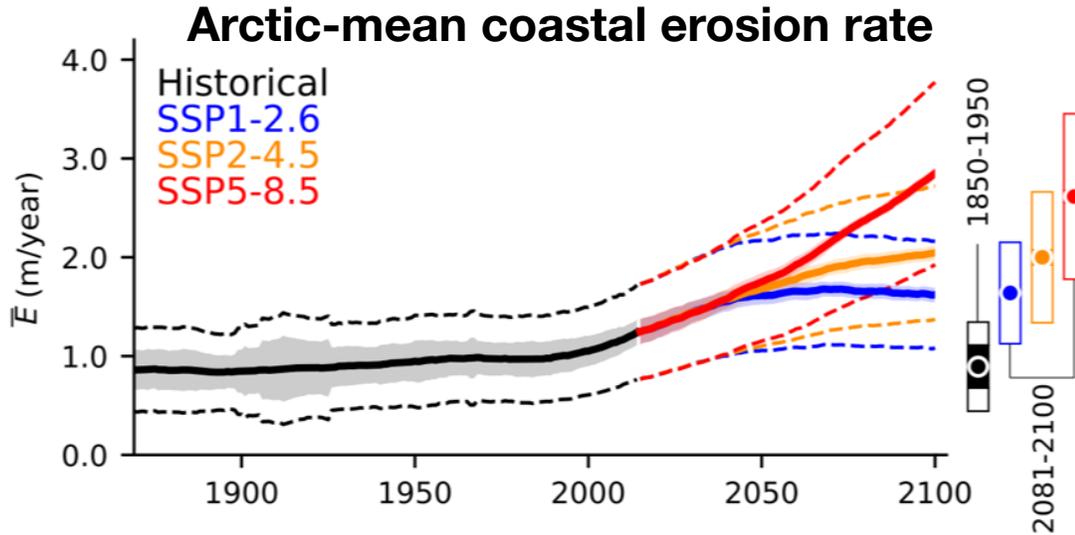
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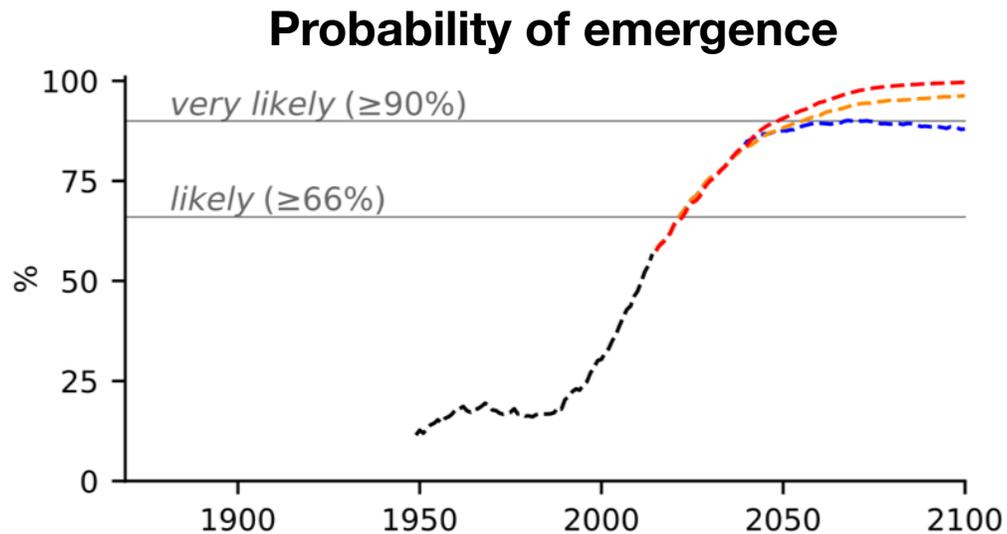
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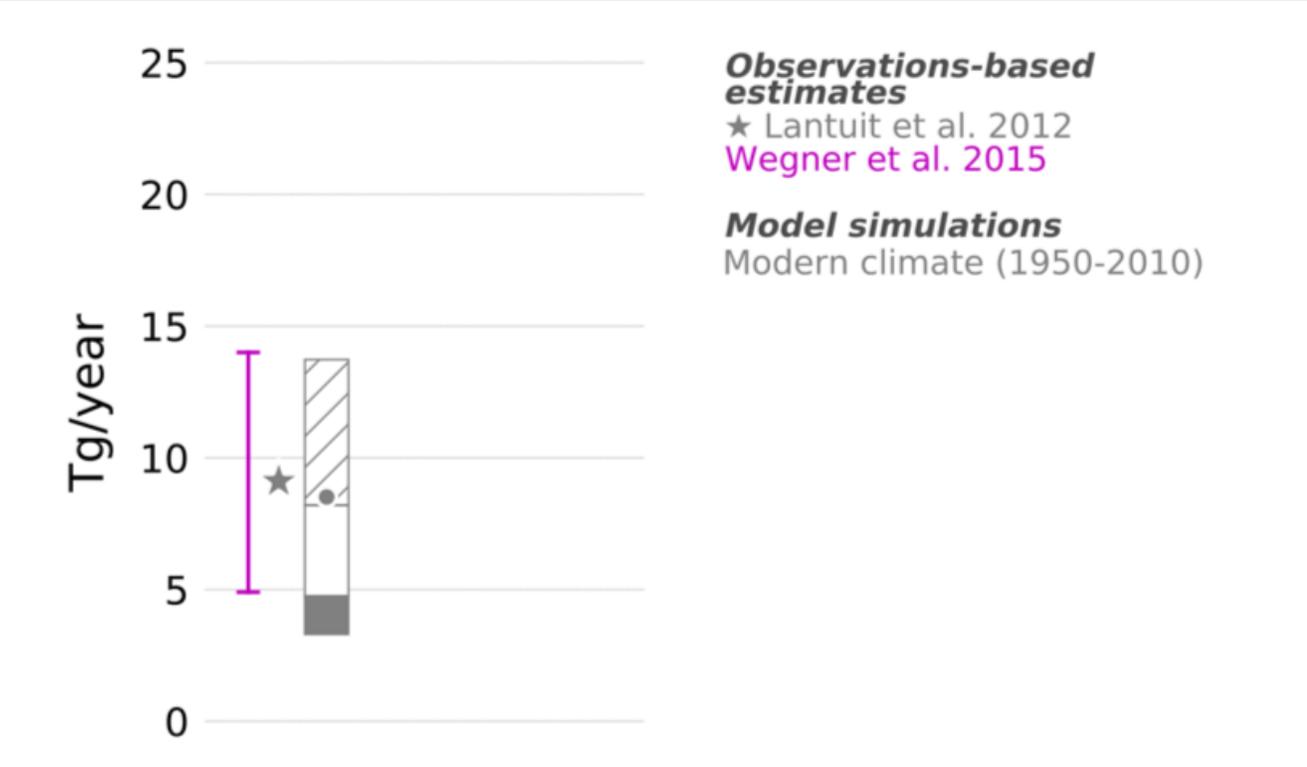
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- Sensitivity to global-mean surface air temperature (GMST) about 0.4 (m/year)/°C (SSP2-4.5)



Permafrost organic carbon loss



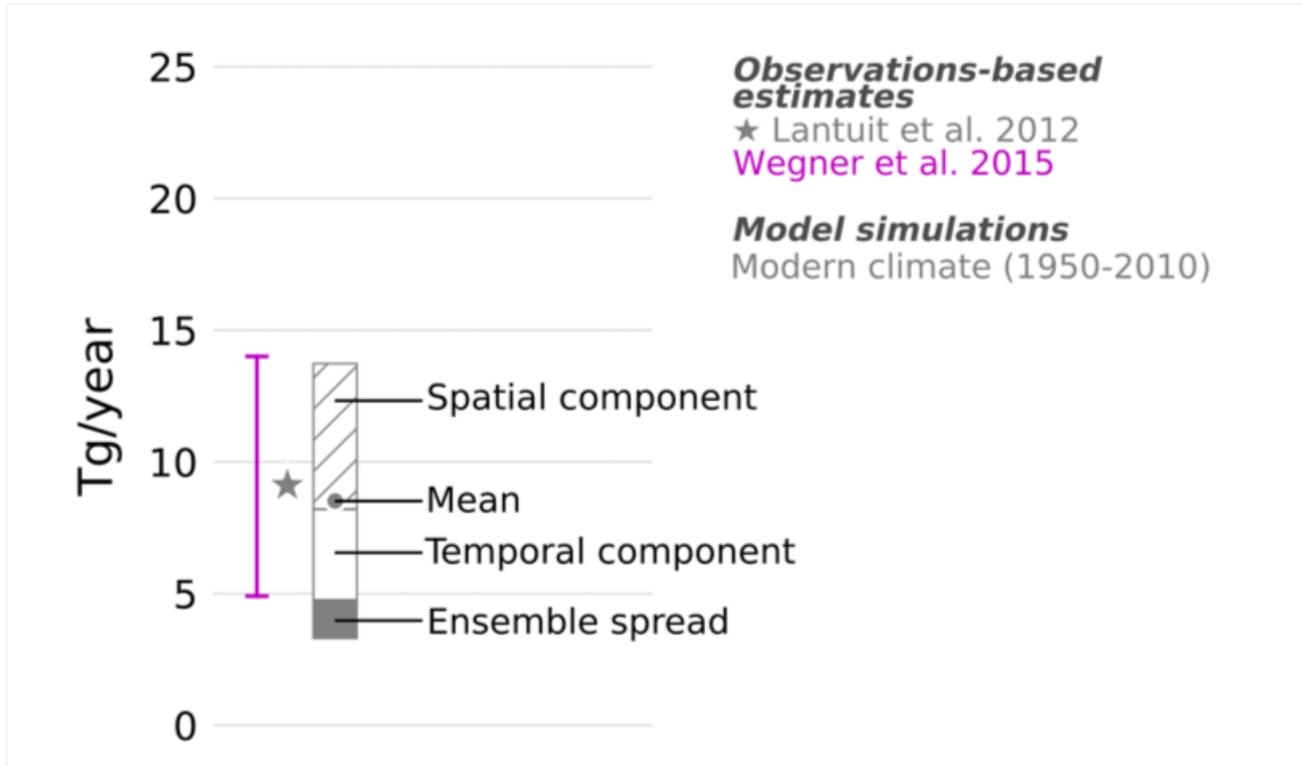
Yearly loss due to coastal erosion



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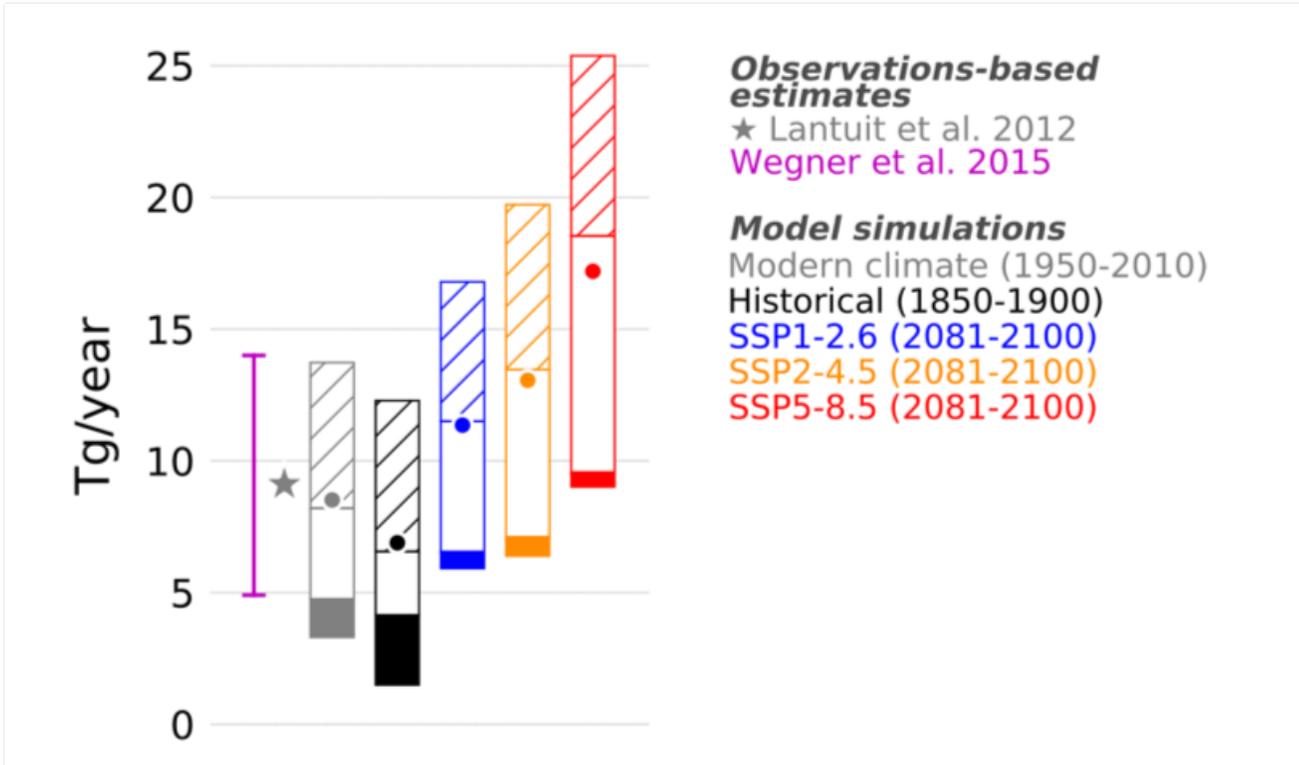


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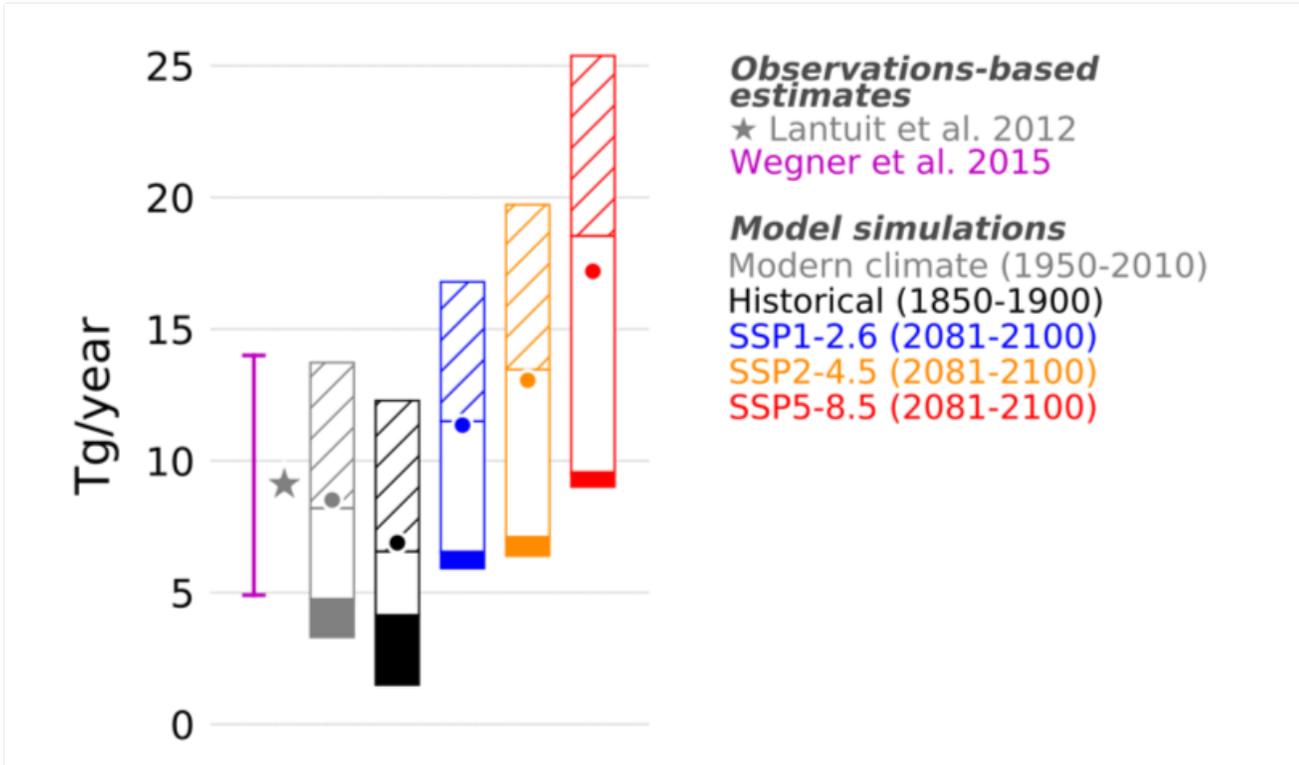


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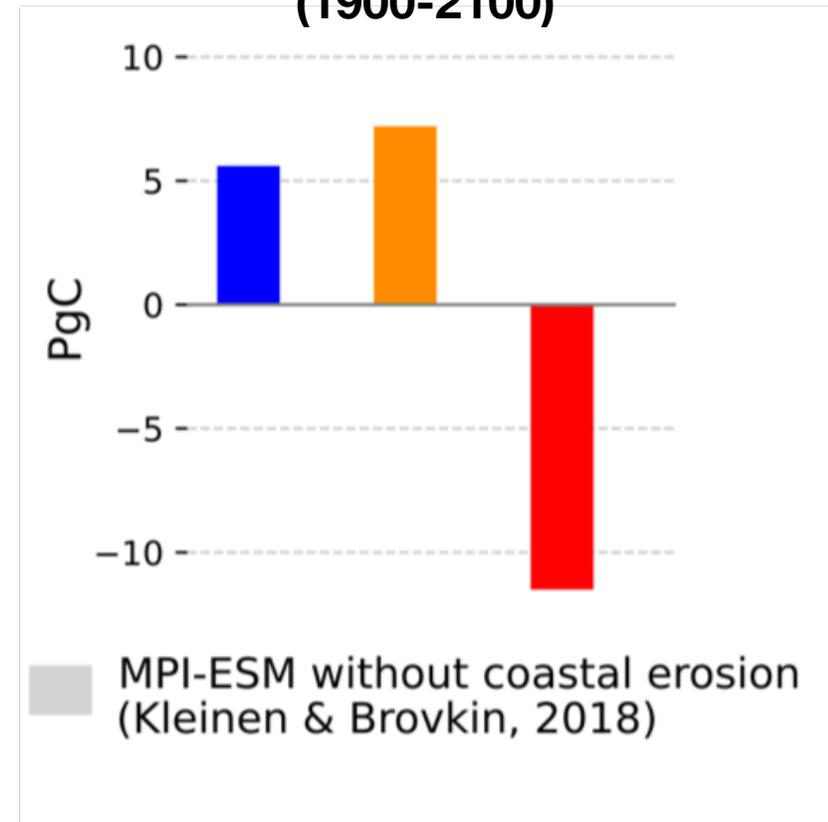
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Net cumulative change of permafrost carbon (1900-2100)

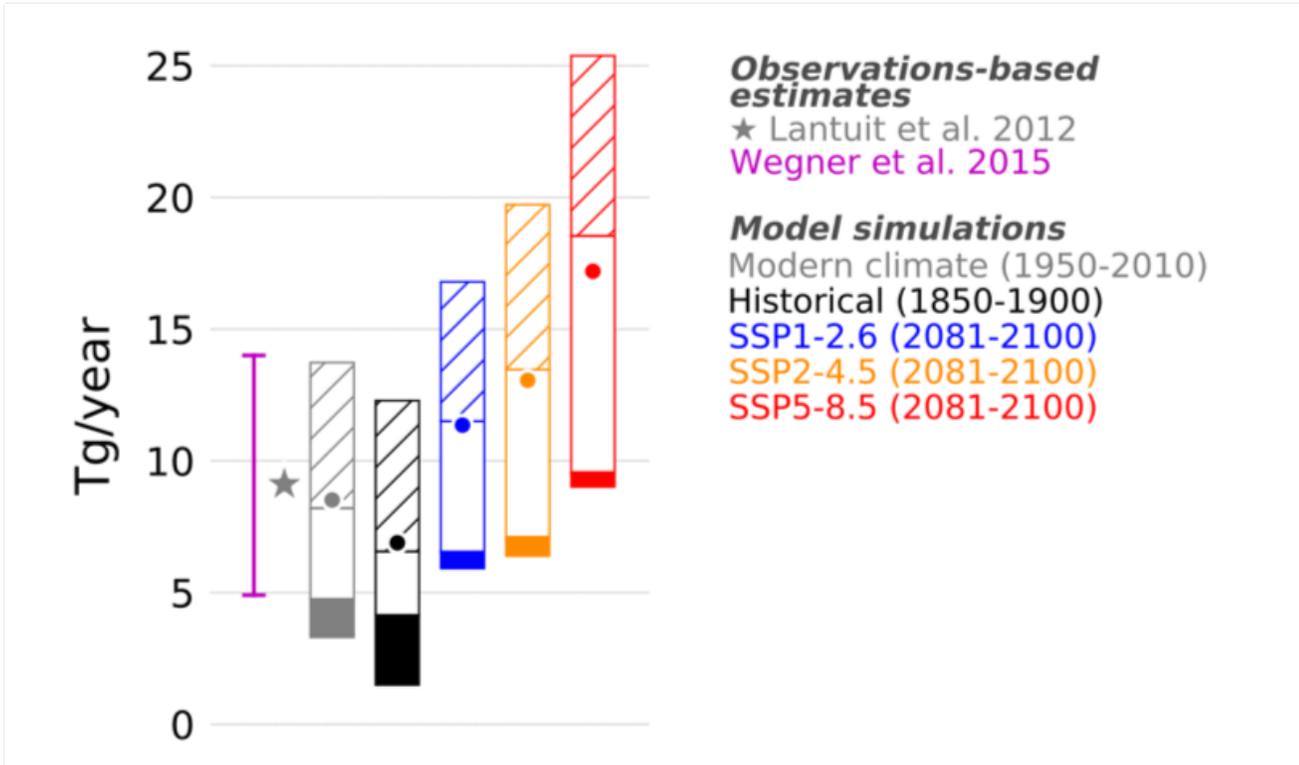


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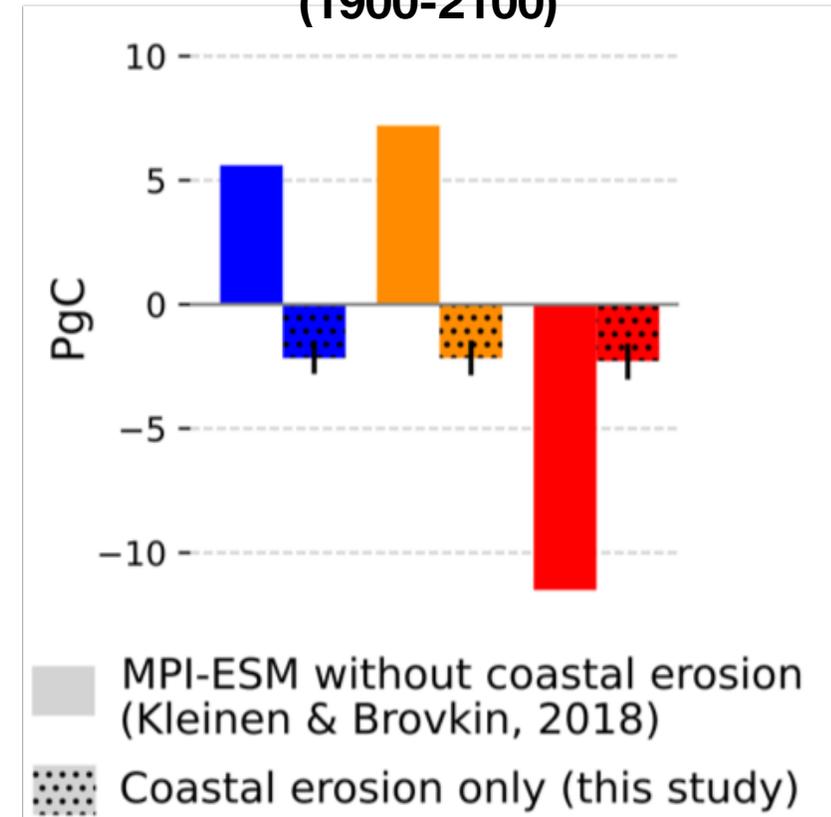
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- Coastal erosion accelerates permafrost carbon loss in all scenarios





Organic Matter

- OM remineralization **releases CO₂** to the ocean and atmosphere

Tanski *et al.* (2019, 2020)



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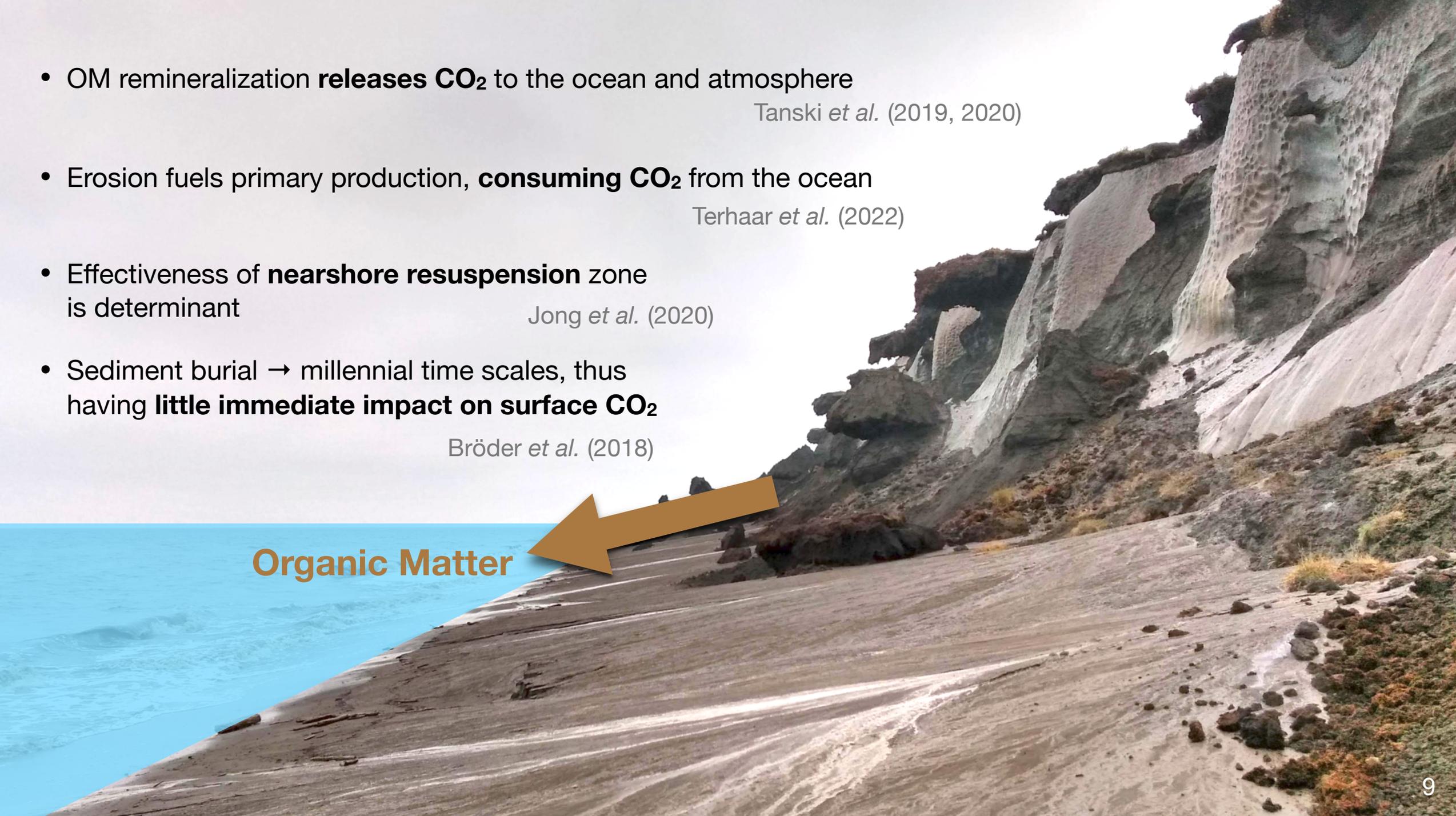
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Organic Matter



What is the impact of increasing permafrost erosion on the Arctic Ocean's CO₂ uptake from the atmosphere?

Permafrost Organic Matter (OM) in MPI-ESM



Limitations

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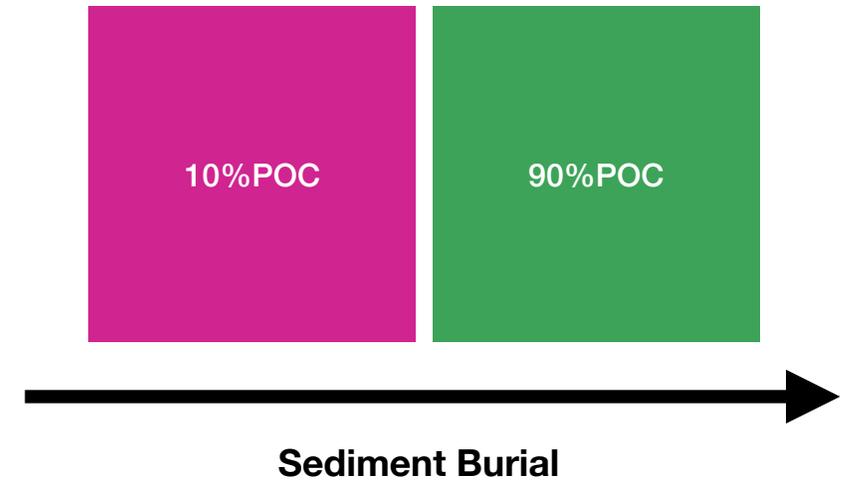


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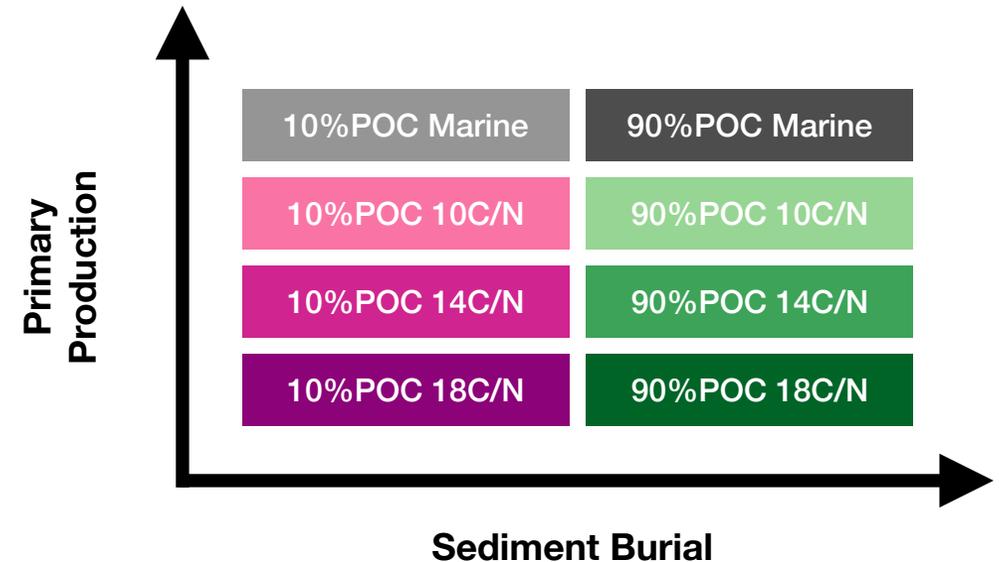


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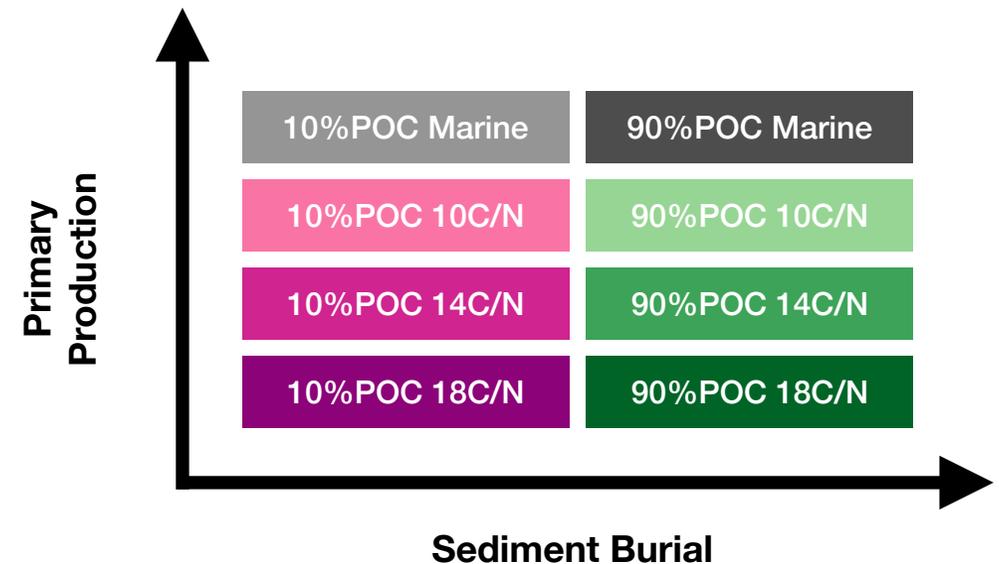


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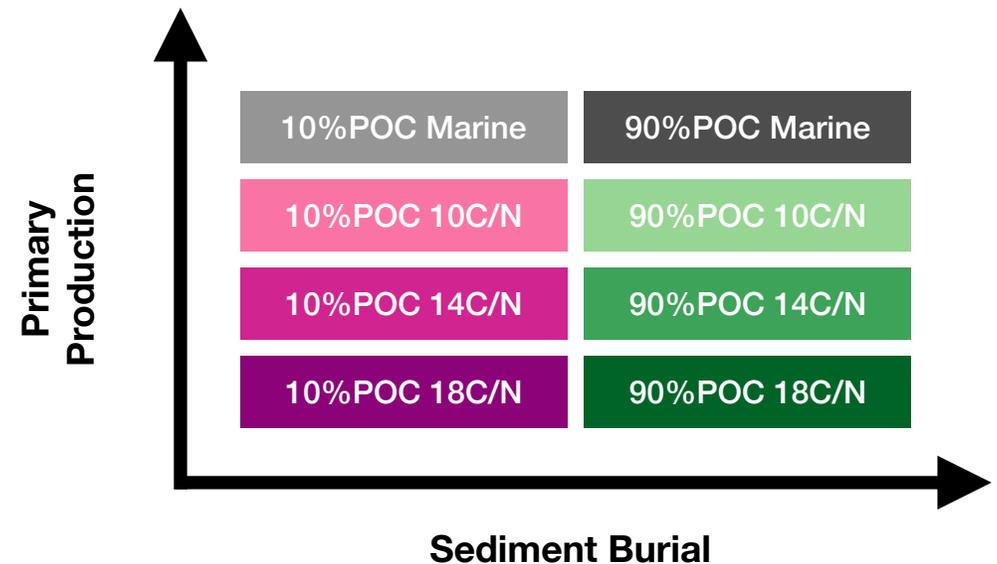


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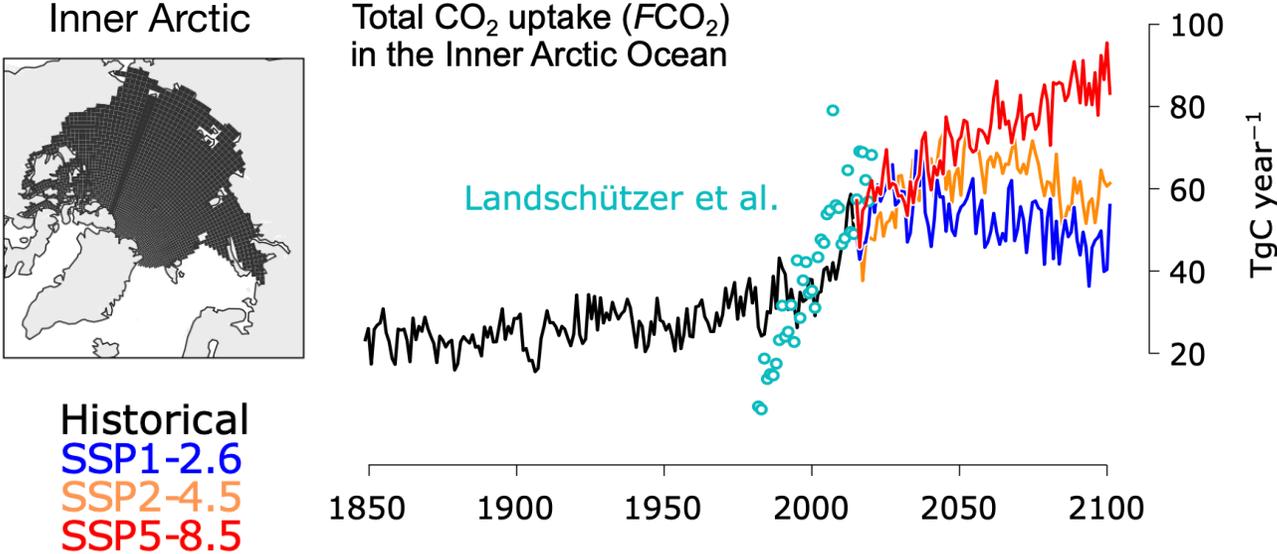
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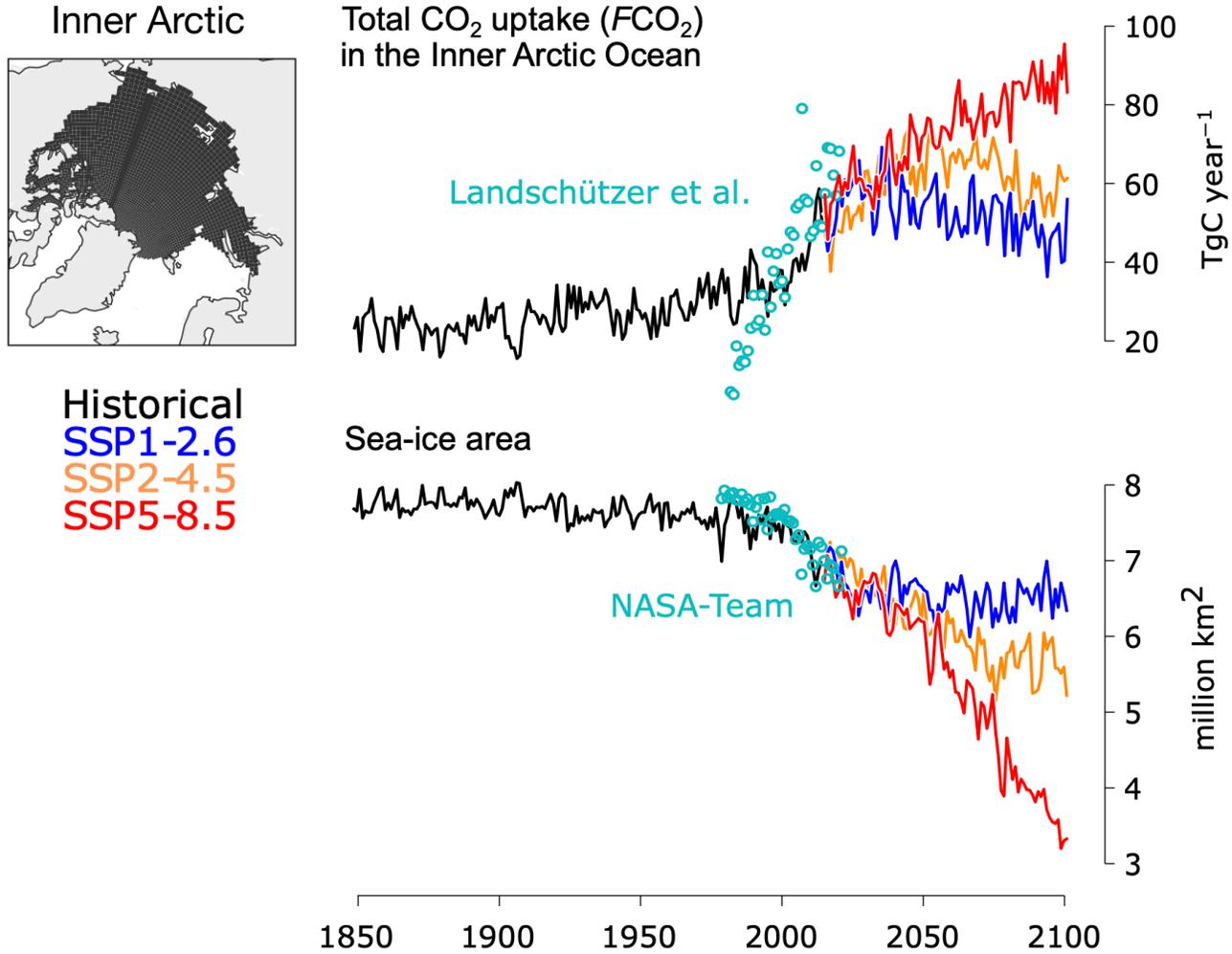
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- 350 years of spin-up with historical erosion rates until stable surface Arctic Ocean
- Historical and Future scenarios until 2100 (CMIP6)



Arctic Ocean's CO₂ uptake



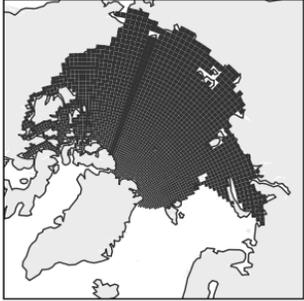
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Arctic Ocean's CO₂ uptake decreases due to erosion

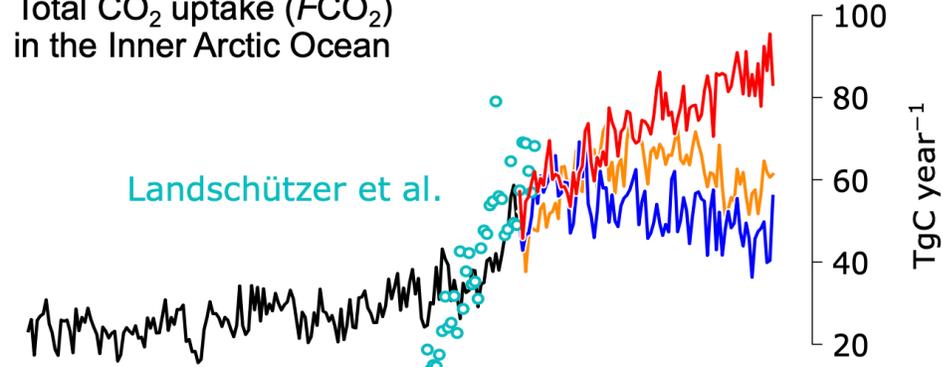


Inner Arctic



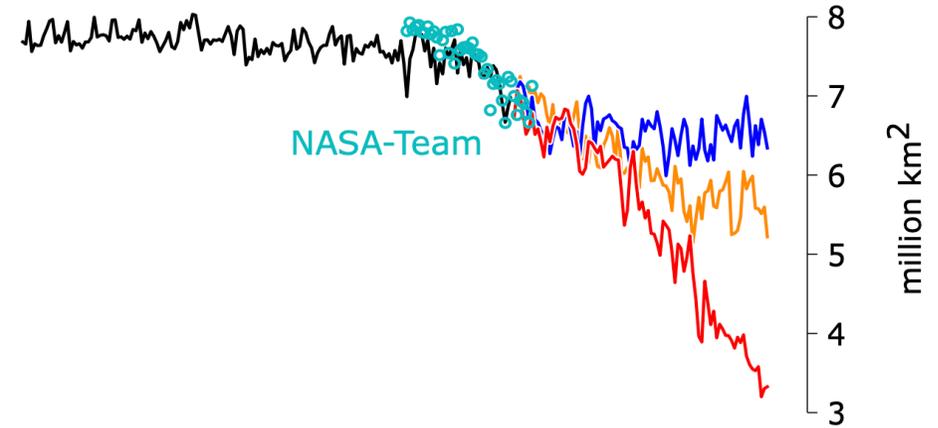
Historical
SSP1-2.6
SSP2-4.5
SSP5-8.5

Total CO₂ uptake (FCO_2)
in the Inner Arctic Ocean



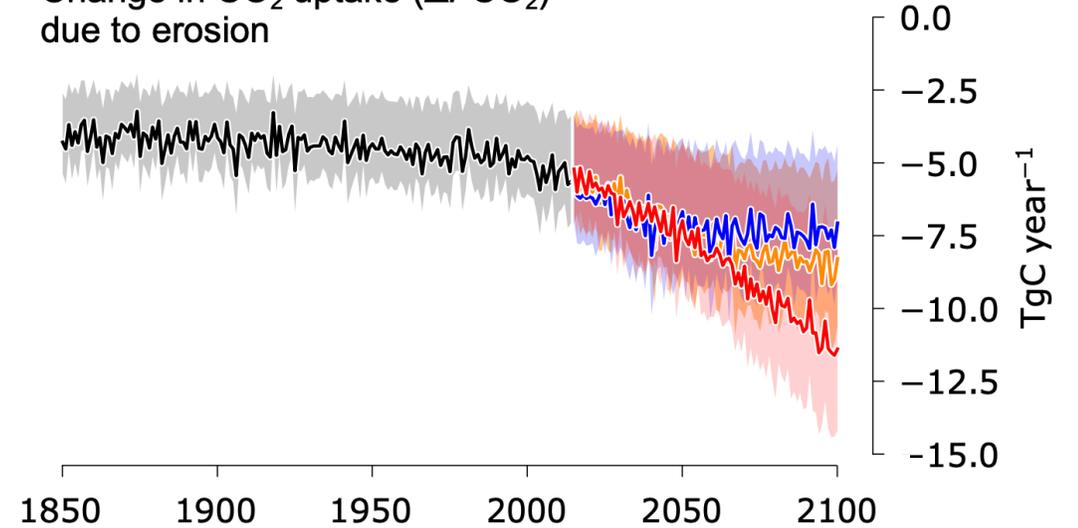
Landschützer et al.

Sea-ice area



NASA-Team

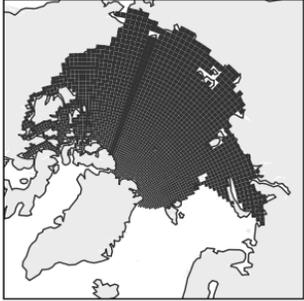
Change in CO₂ uptake (ΔFCO_2)
due to erosion



Arctic Ocean's CO₂ uptake decreases due to erosion

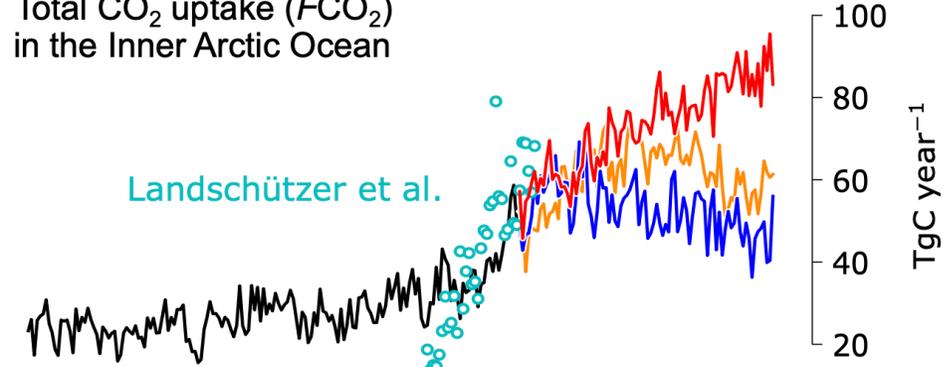


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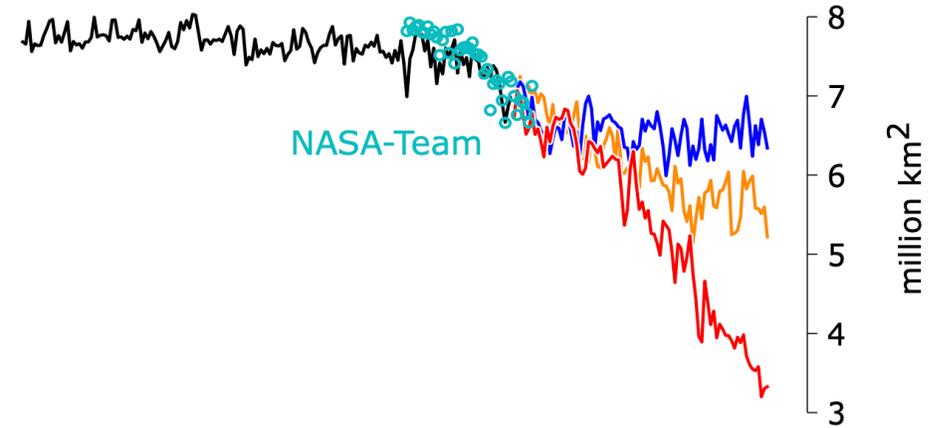
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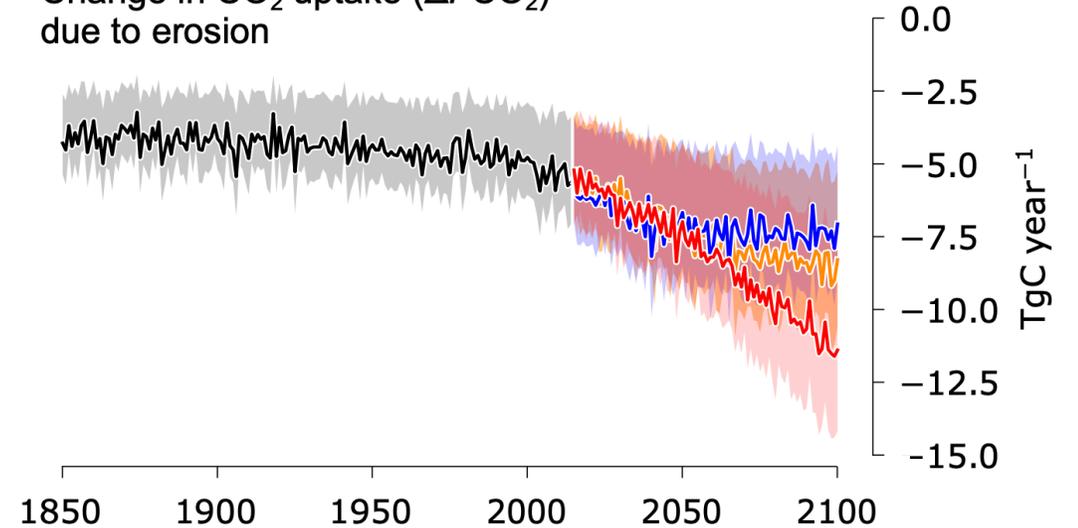
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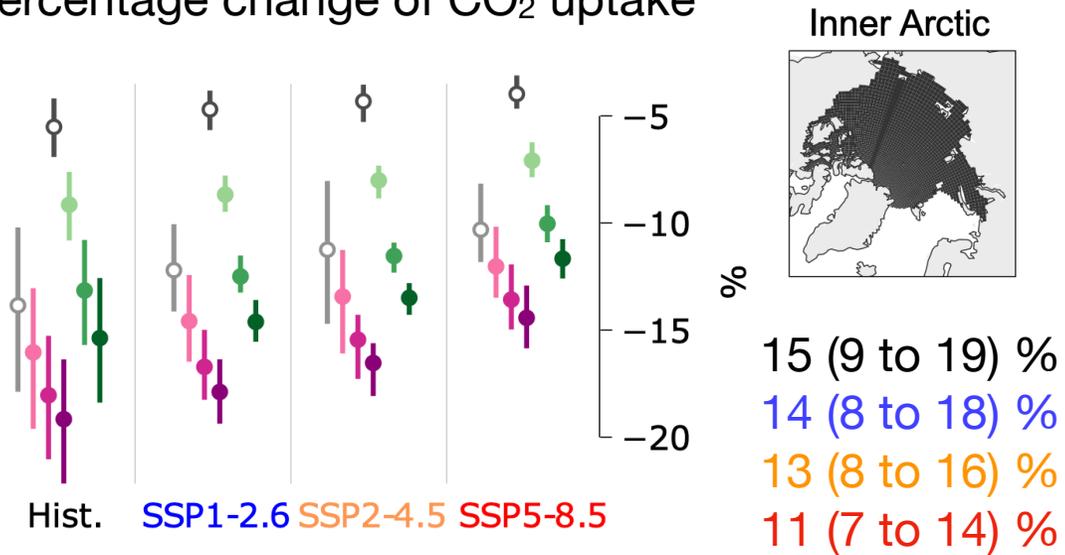


- Reduction of CO₂ uptake by about 5 to 14 TgC/year
- Model uncertainties are as large as scenario uncertainties

Uncertainties originate in the OM representation



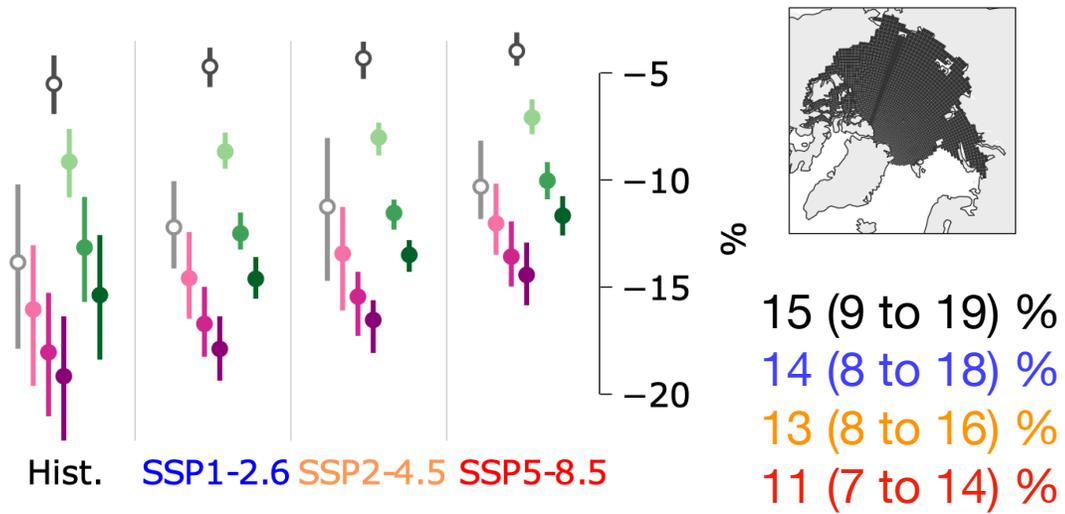
Percentage change of CO₂ uptake



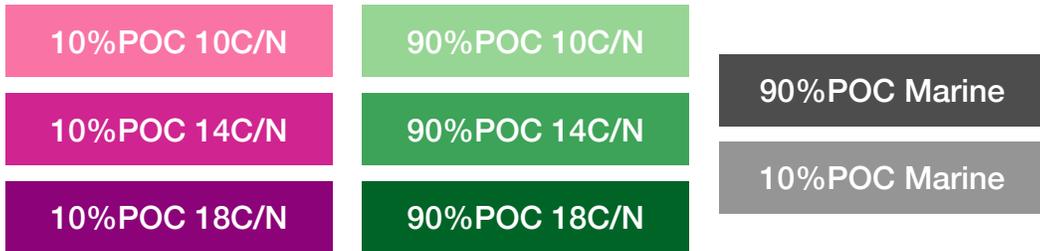
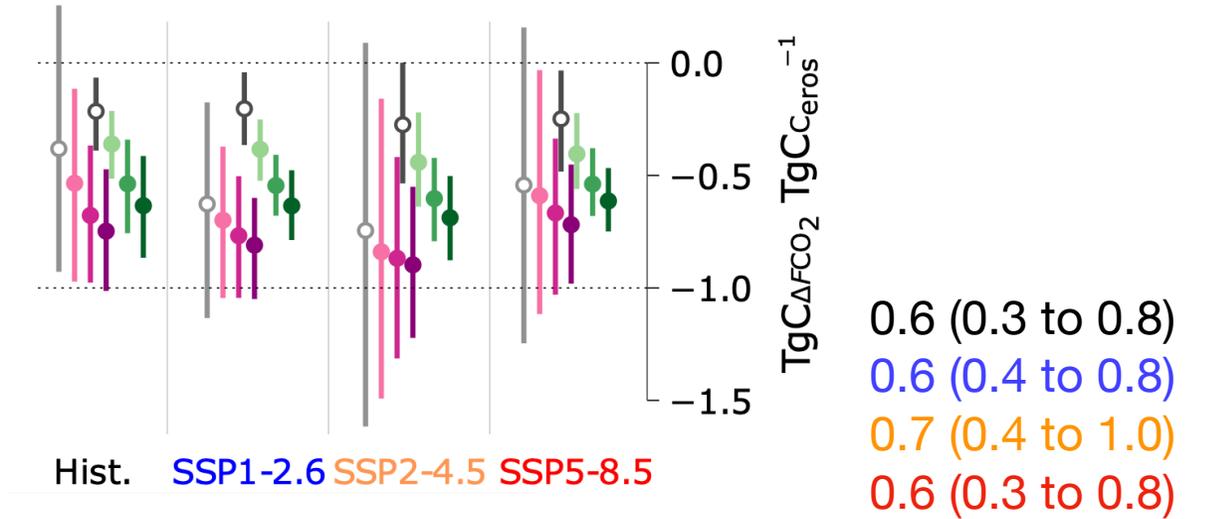
Uncertainties originate in the OM representation



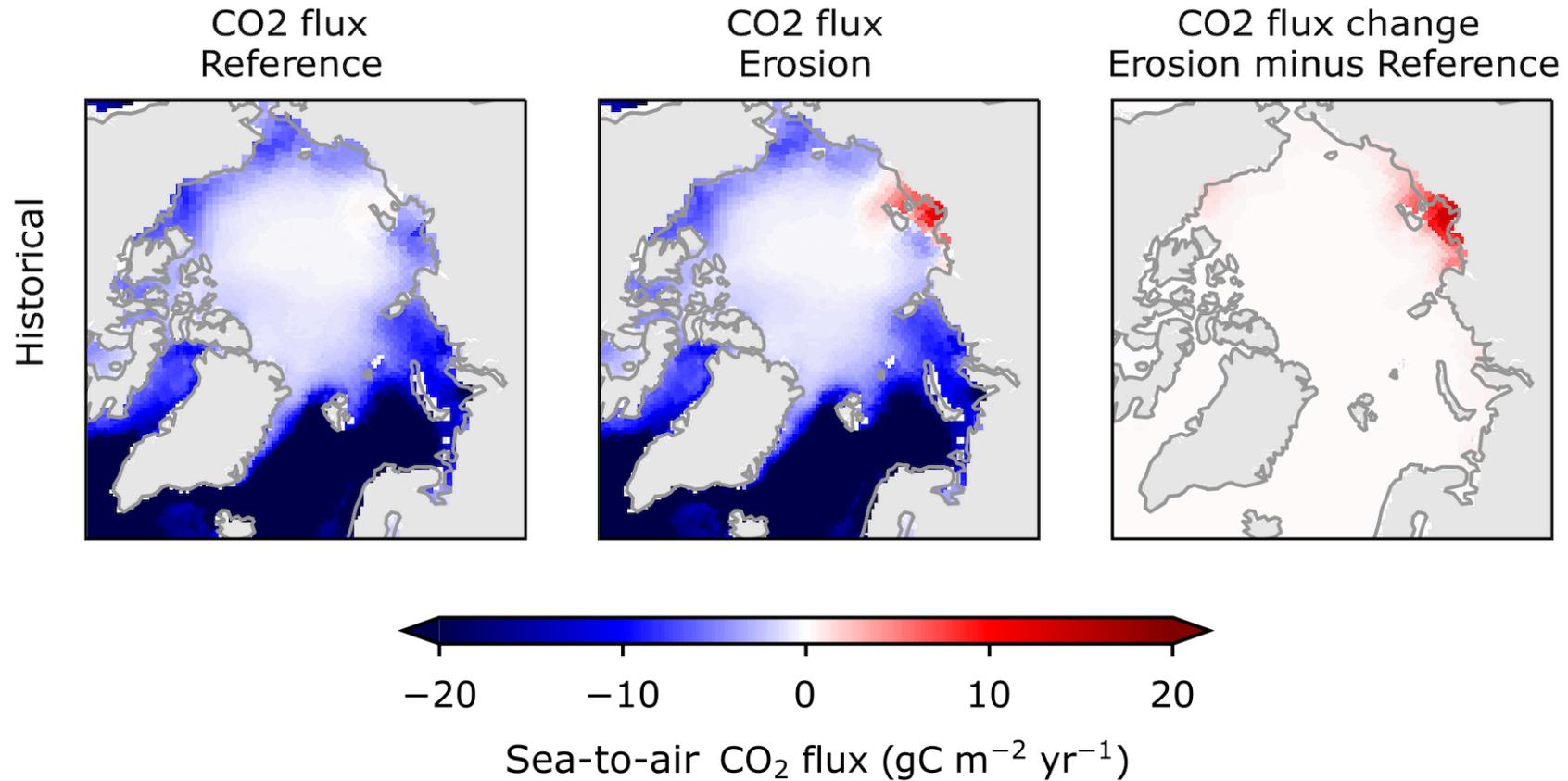
Percentage change of CO₂ uptake



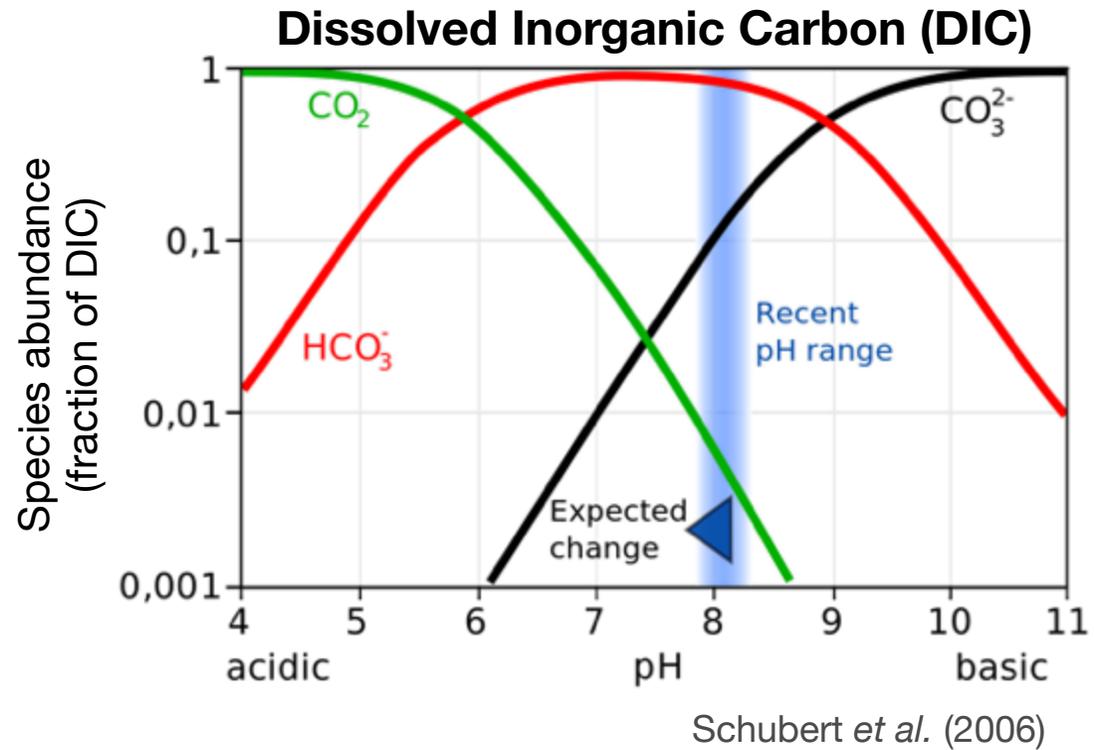
Sensitivity of CO₂ uptake change to erosion



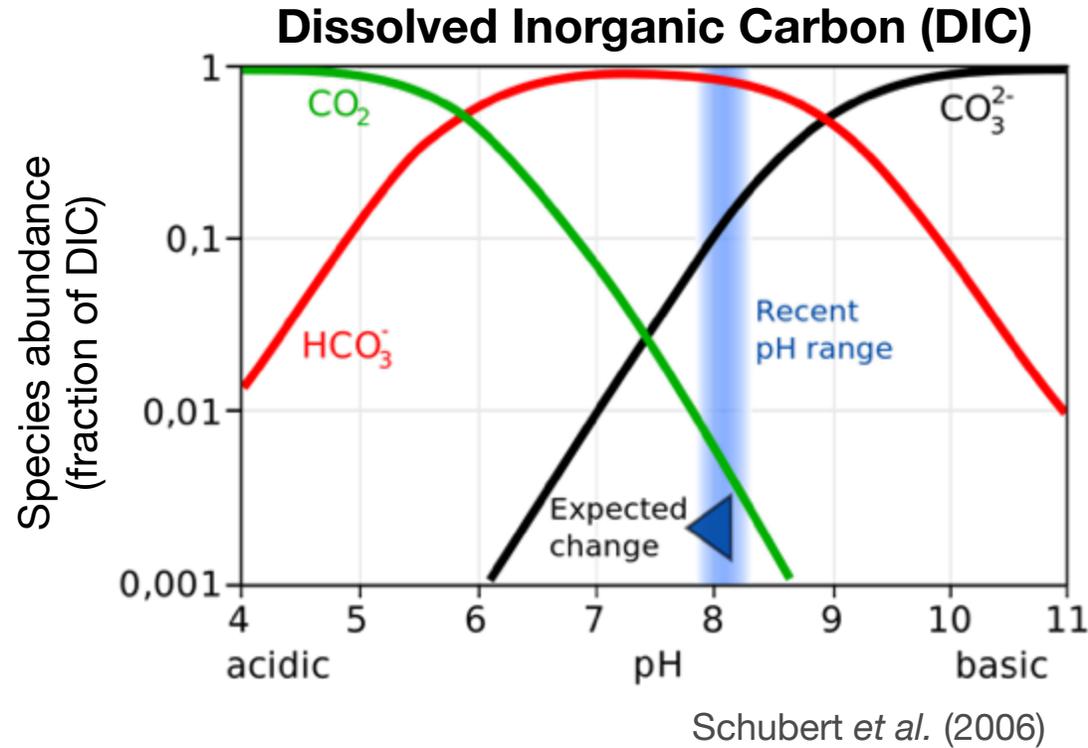
Erosion inverts the air-sea CO₂ flux direction on the Siberian shelf



Mechanisms of pCO₂ change in the Ocean



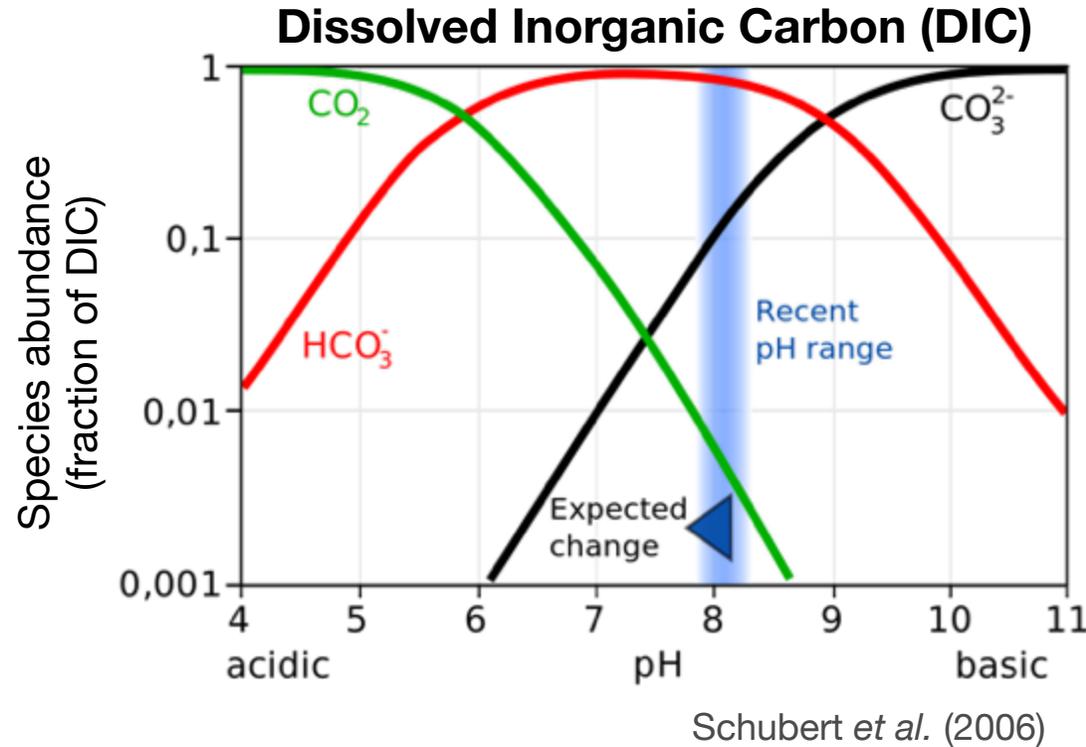
Mechanisms of pCO₂ change in the Ocean



↑DIC → ↑pCO₂

↓Alkalinity → ↑pCO₂

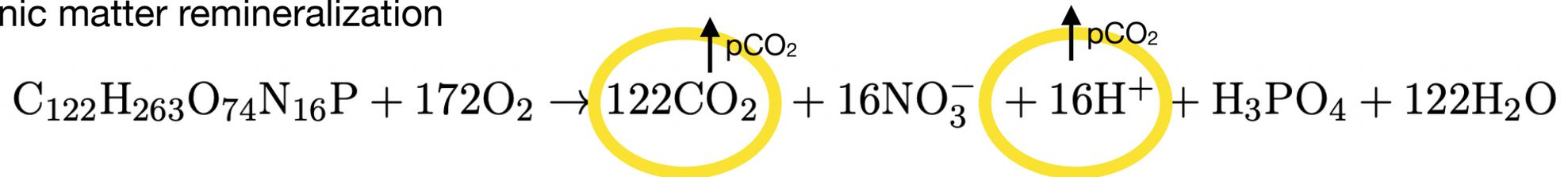
Mechanisms of pCO₂ change in the Ocean



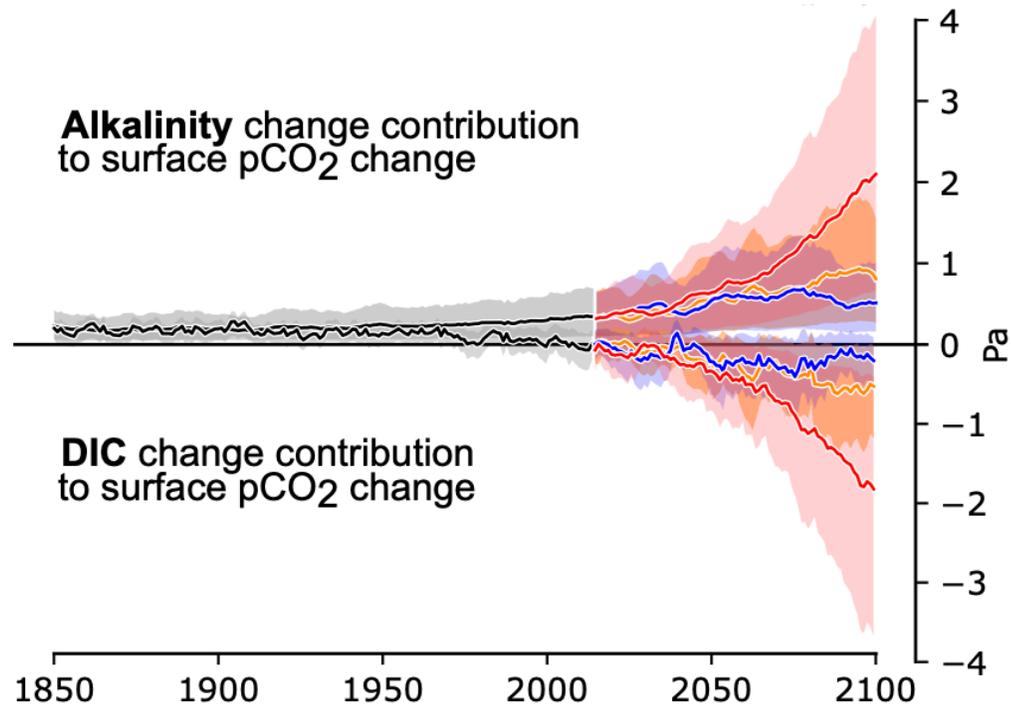
↑DIC → ↑pCO₂

↓Alkalinity → ↑pCO₂

Organic matter remineralization

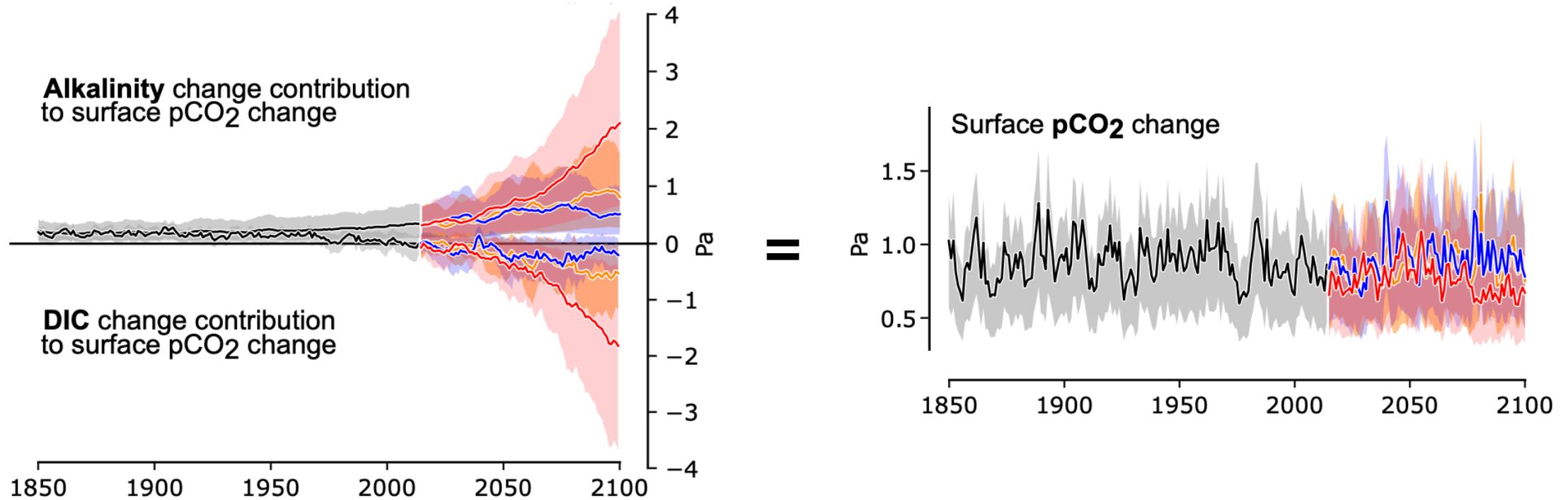


Decomposing pCO₂ in contributions from alkalinity and DIC



- Alkalinity decreases (OM remineralization) and decreases the Ocean's ability to take up CO₂

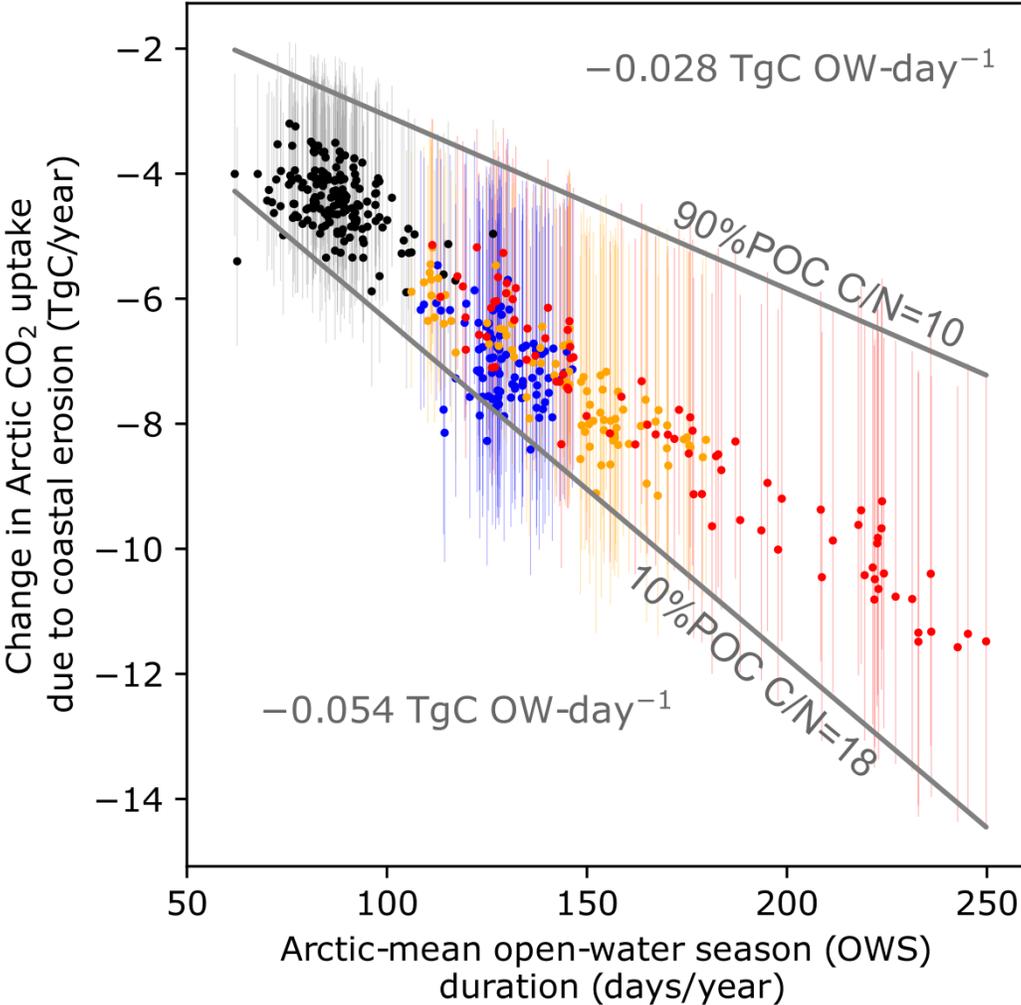
Decomposing pCO₂ in contributions from alkalinity and DIC



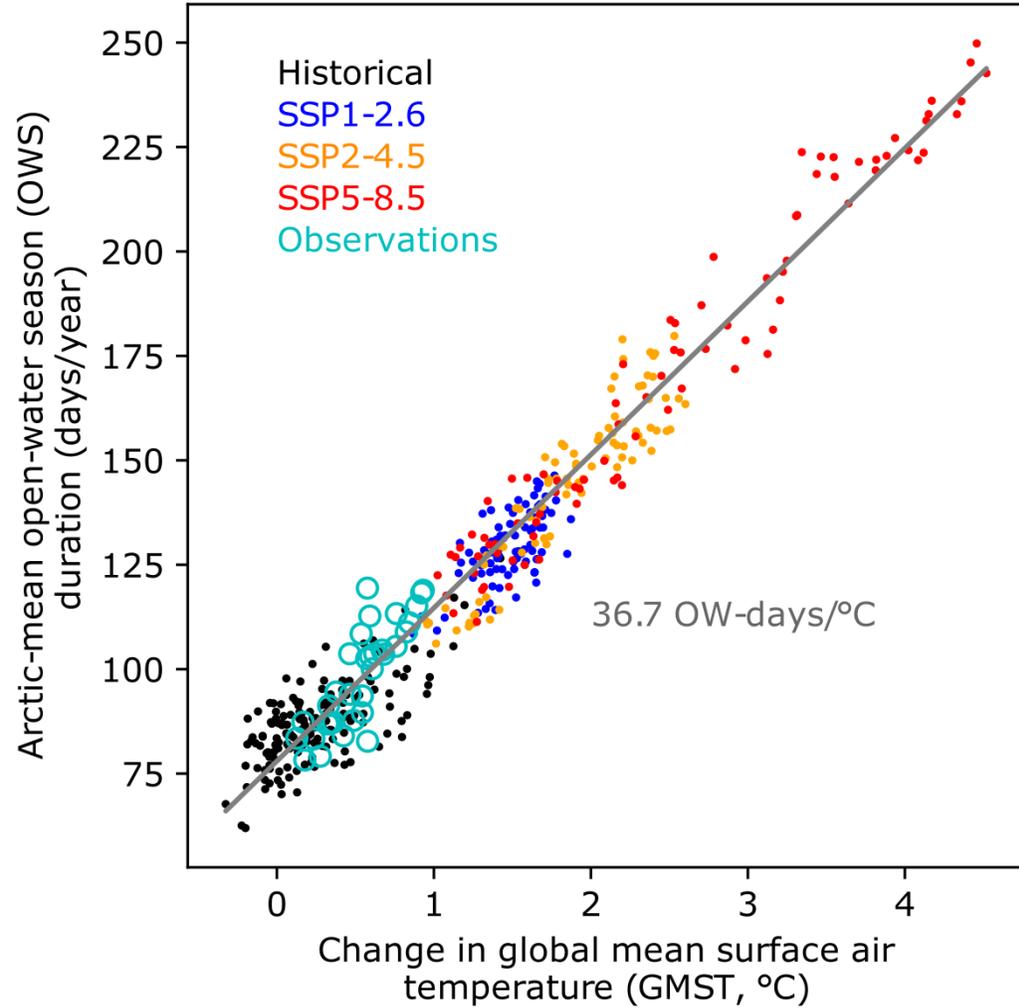
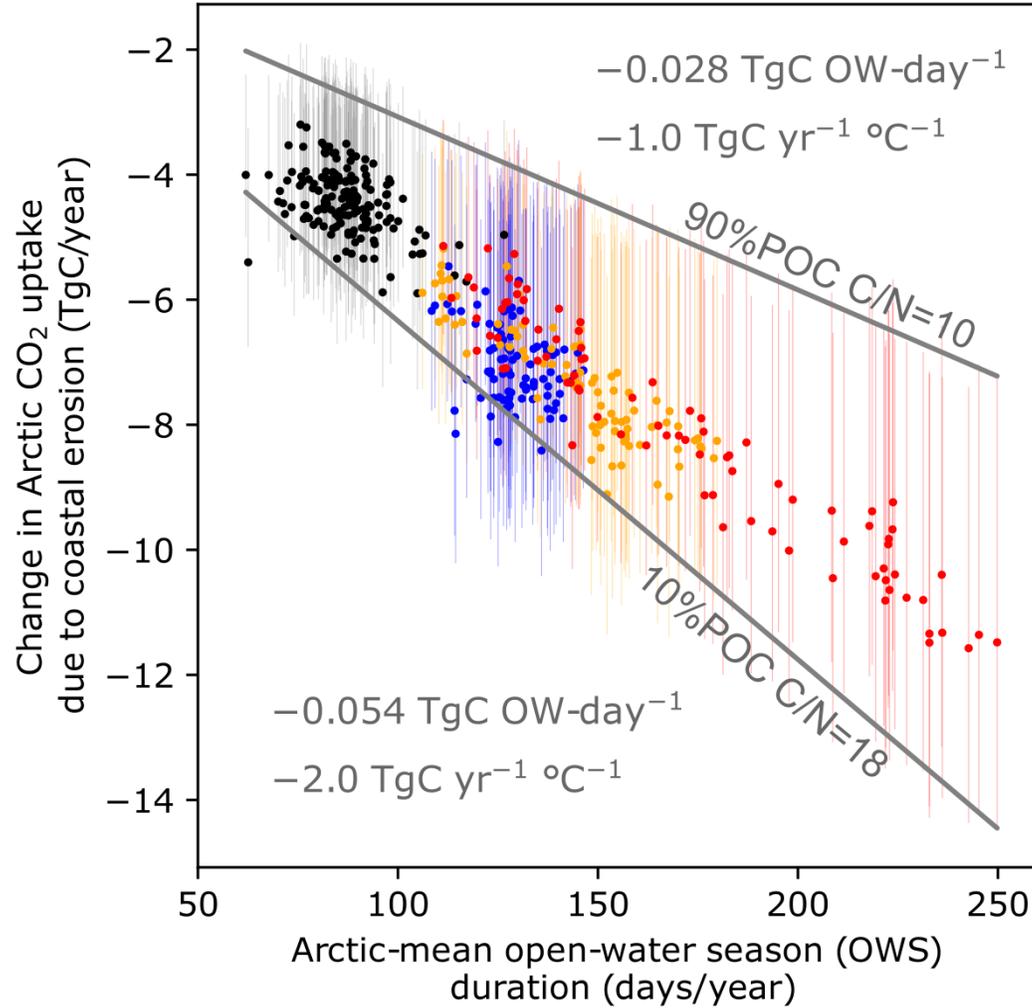
- Alkalinity decreases (OM remineralization) and decreases the Ocean's ability to take up CO₂

- The combined effect of decreasing alkalinity and DIC maintain a constant surface pCO₂ increase

Sea-ice modulates the impact of erosion on CO₂ uptake



Sea-ice modulates the impact of erosion on CO₂ uptake



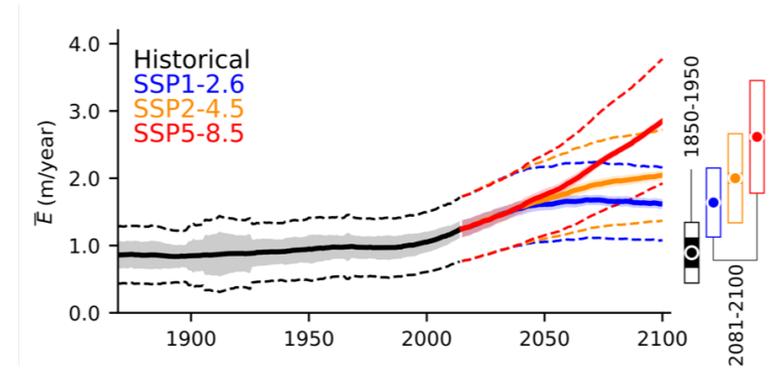
Summary



- Simple, semi-empirical model for Arctic coastal permafrost erosion compatible with ESMs

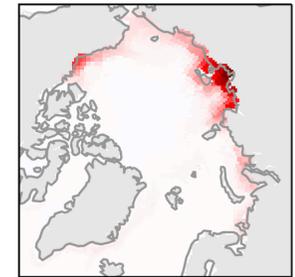
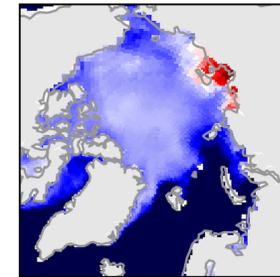
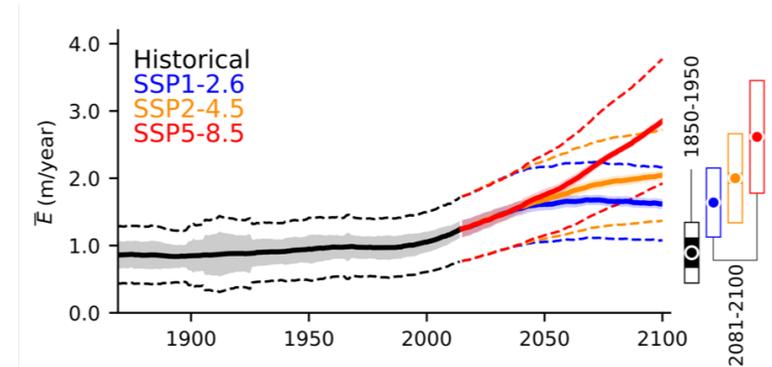
Summary

- Simple, semi-empirical model for Arctic coastal permafrost erosion compatible with ESMs
- **Erosion rates could increase by a factor of 2-to-3 by 2100 and exceed the historical range of variability**



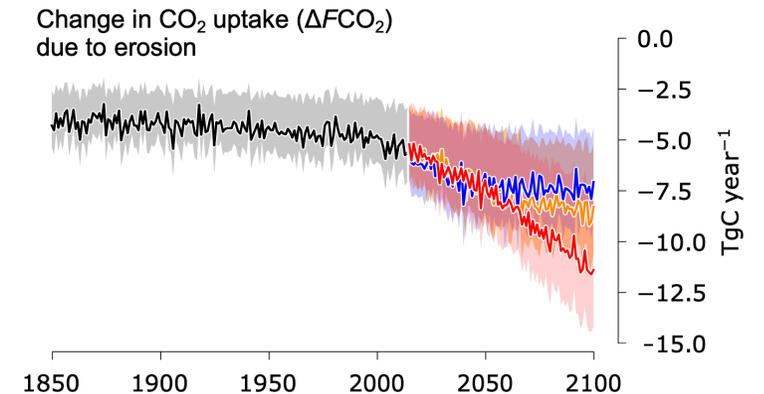
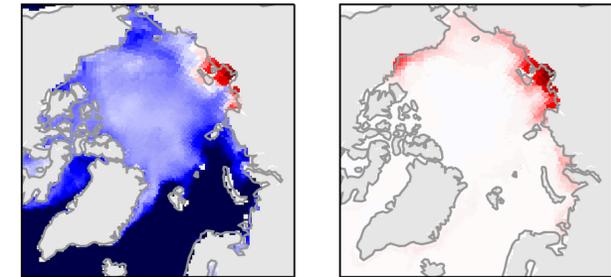
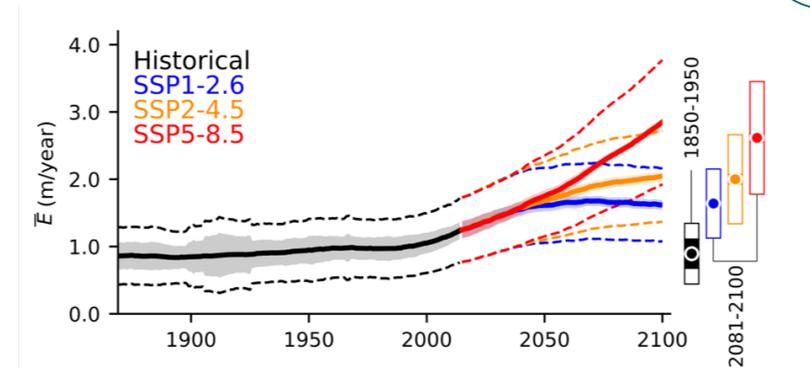
Summary

- Simple, semi-empirical model for Arctic coastal permafrost erosion compatible with ESMs
- **Erosion rates could increase by a factor of 2-to-3 by 2100** and exceed the historical range of variability
- **Coastal erosion reduces the Arctic Ocean's CO₂ uptake** from the atmosphere by 5 to 15 TgC/year, which is equivalent to:
 - **15(±5)%** of Inner Arctic uptake in the historical period
 - **11(±4)%** of Inner Arctic uptake in SSP5-8.5 (2081-2100)
 - **Between 1 and 2 TgC year⁻¹ °C⁻¹ GMST**

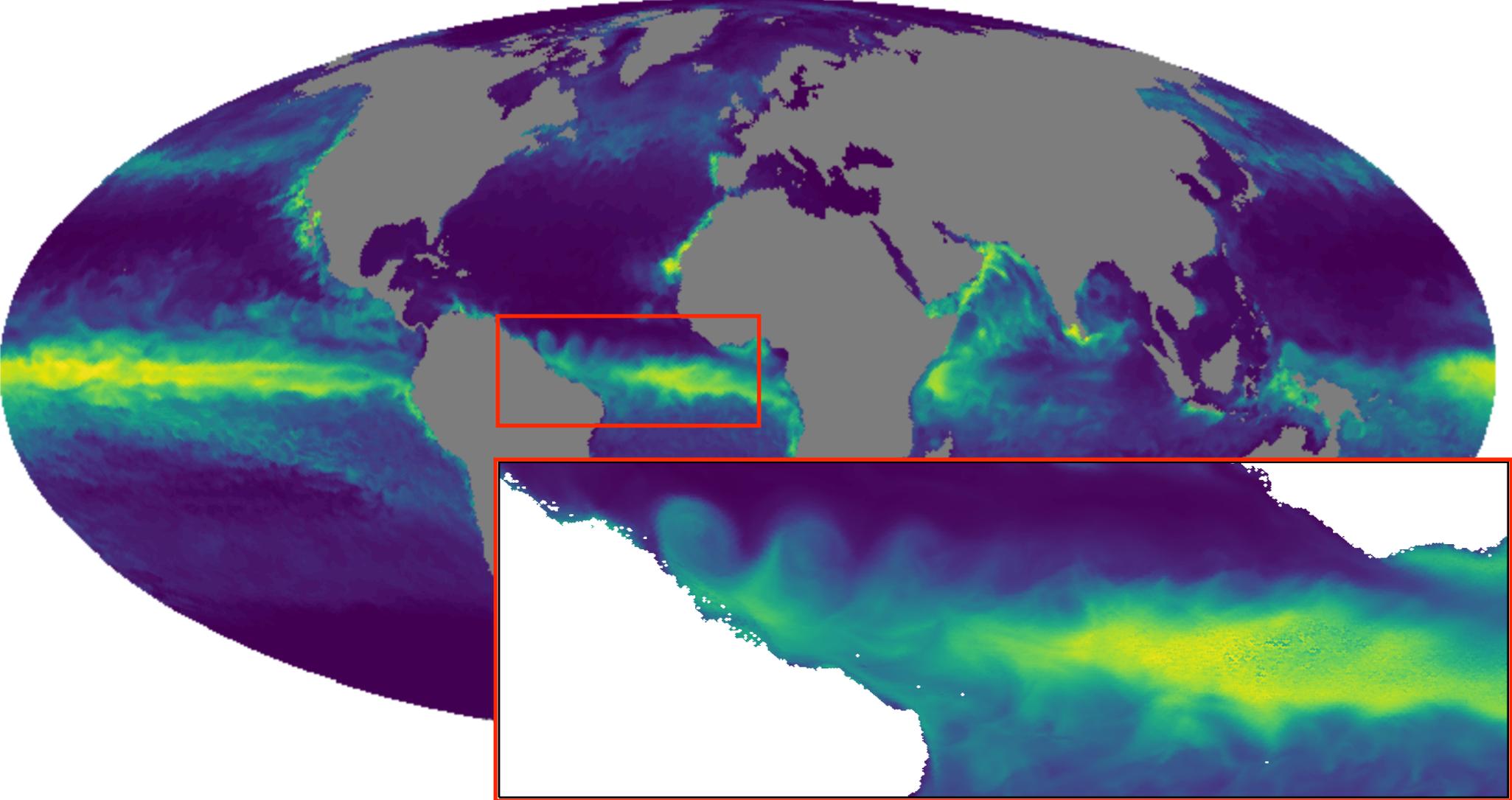


Summary

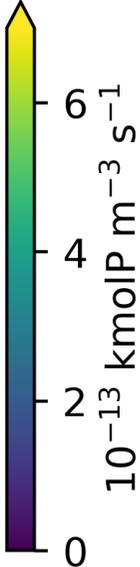
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 - Protecting the coastal permafrost from erosion, and
 - Modulating the effect of erosion on the Ocean's CO₂ uptake



Global, km-scale ICON-ESM with Ocean Biogeochemistry



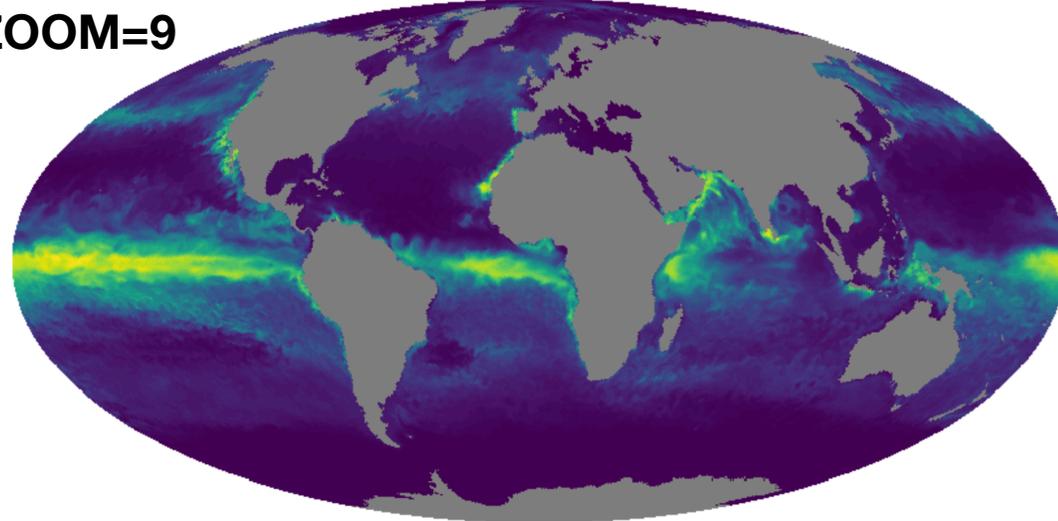
Net Primary production



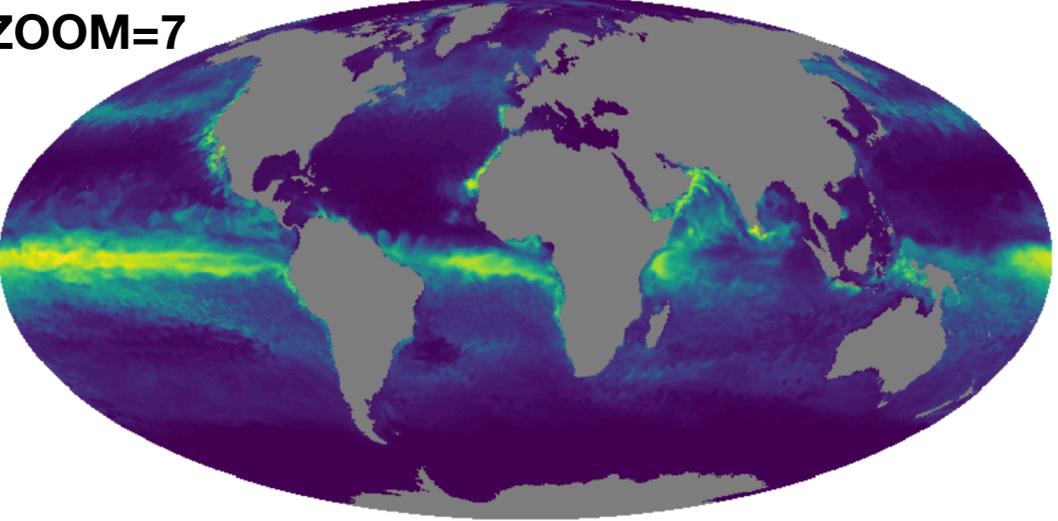
HEALPix (Hierarchical Equal Area isoLatitude Pixelation)



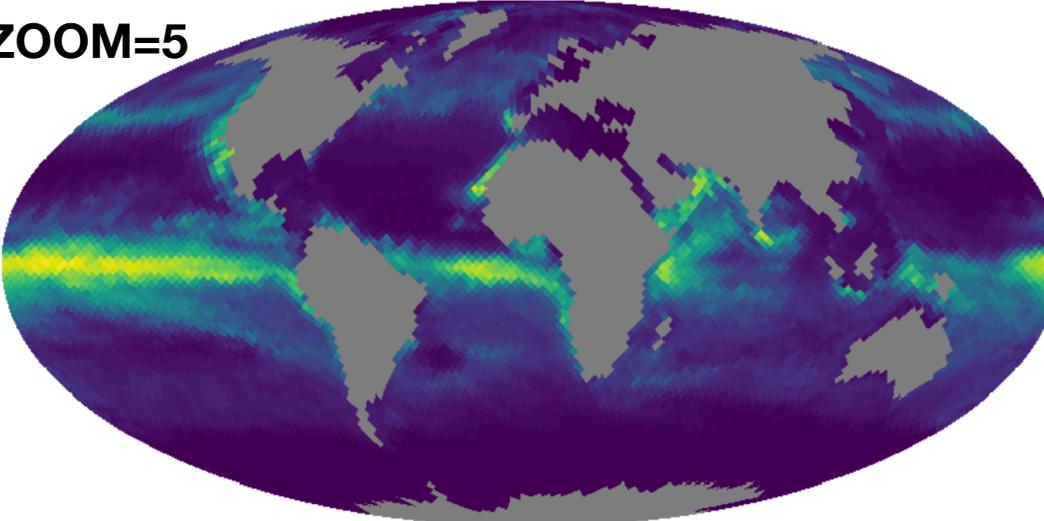
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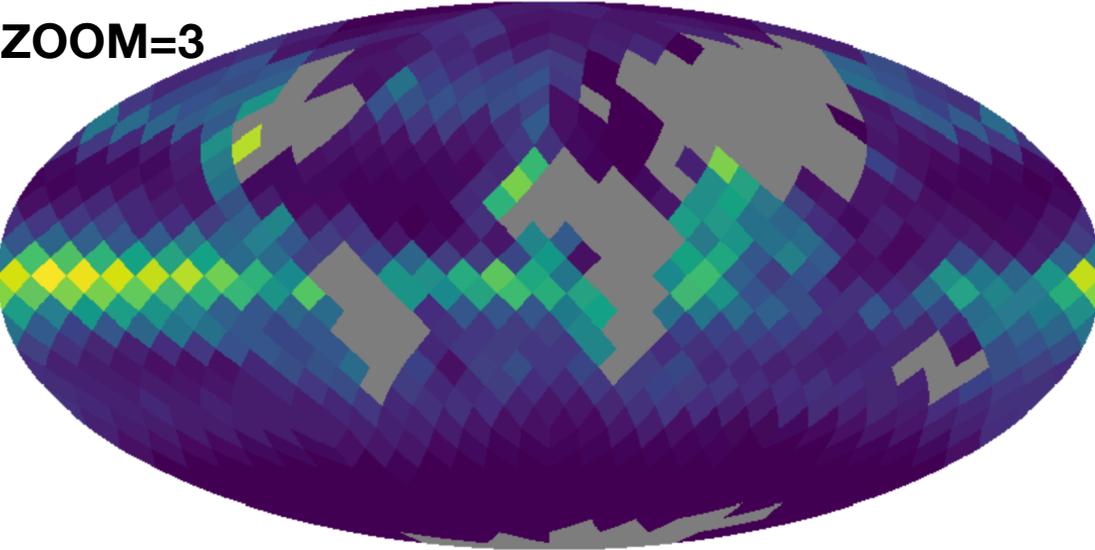
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ZOOM=5

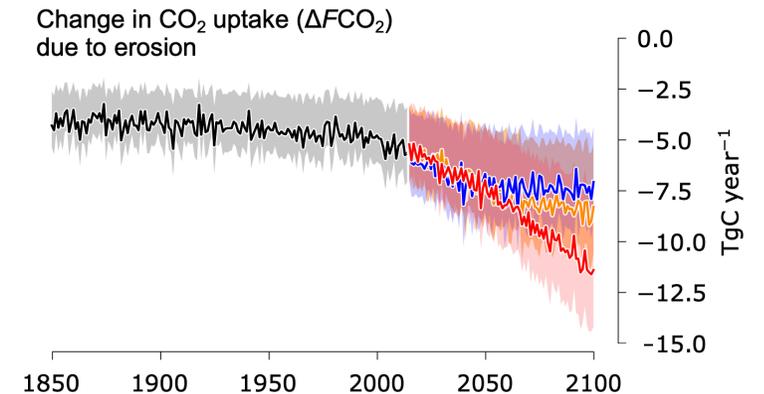
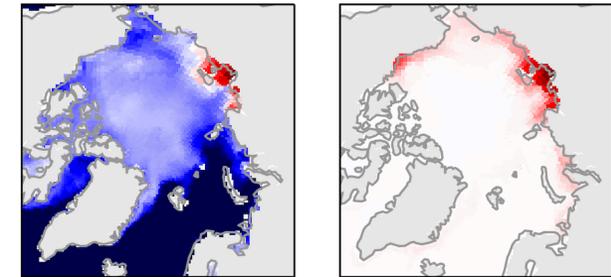
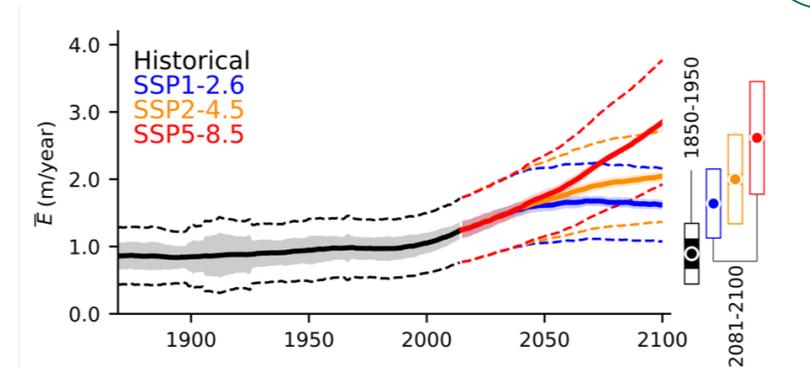


ZOOM=3



Summary

- Simple, semi-empirical model for Arctic coastal permafrost erosion compatible with ESMs
- **Erosion rates could increase by a factor of 2-to-3 by 2100** and exceed the historical range of variability
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- **Sea ice** plays a fundamental role in:
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Extra slides

A simple, semi-empirical model for Arctic coastal erosion

as a linear combination of the thermal and mechanical drivers of erosion,
compatible with ESM simulations

$$E(x, t) = \overline{E}(t) + \Delta E(x, t)$$

Temporal Spatial

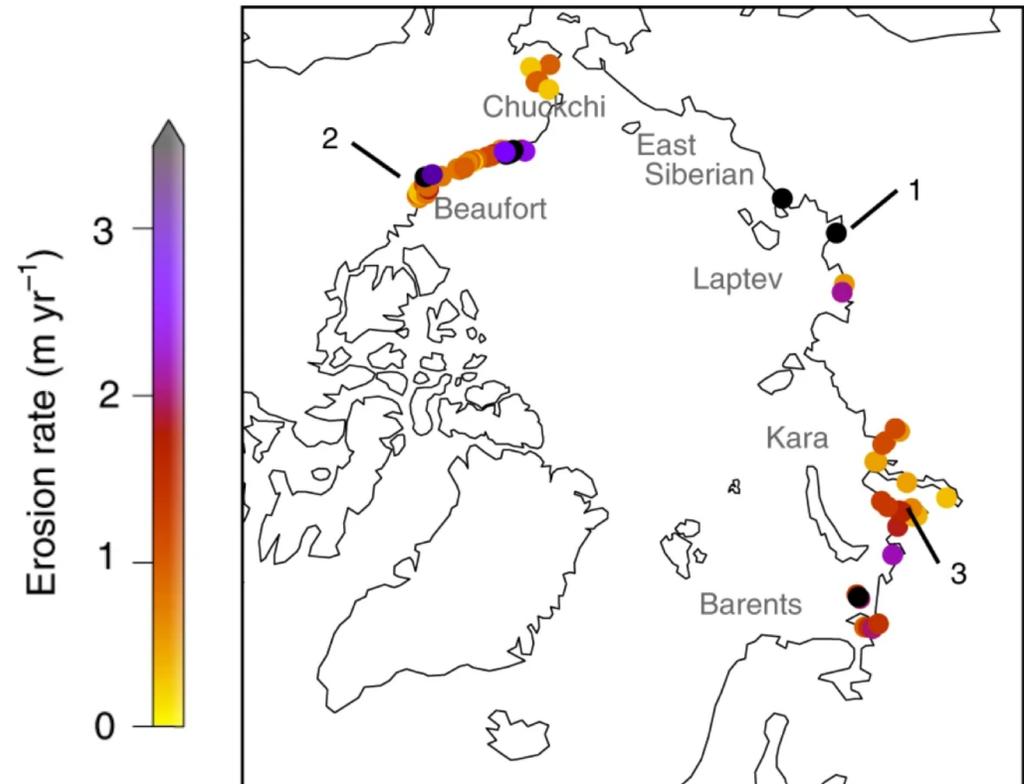
$$E(x, t) = 0 \text{ if SIC}(x, t) \geq 15\%$$

$$\overline{E}(t) = a_{TD} \overline{T}(t) + a_{TA} \overline{H}(t)$$

Thermal Mechanical

$$\Delta E(x, t) = \text{Linear regression of waves, temperature and ground ice}$$

ACD database (Lantuit et al. 2012)
“High-quality” segments

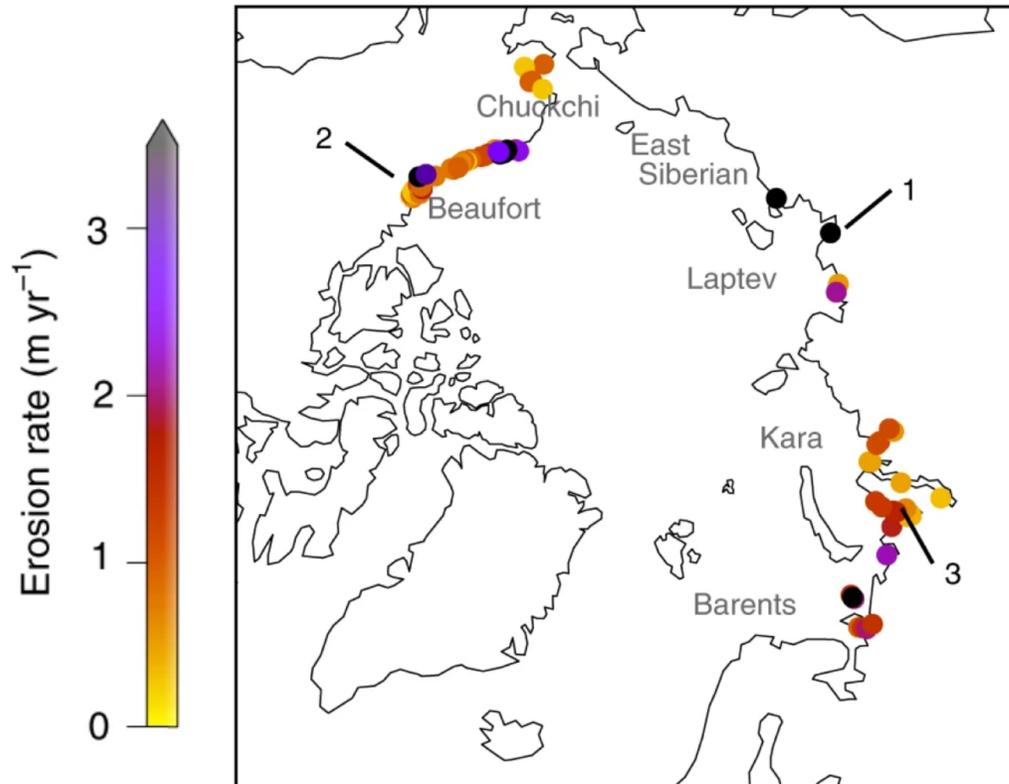


- 1: Bykovsky Peninsula and Muostakh Island
- 2: Mackenzie River Delta region
- 3: Yamal Peninsula

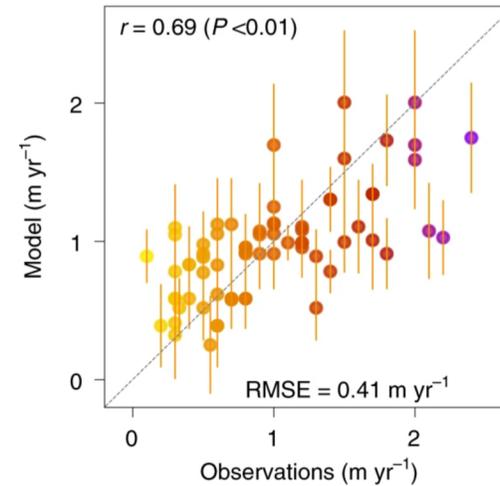
The Spatial Component

$$\Delta E(x, t) = \begin{cases} b_{\theta} \Delta \theta(x) + b_H \Delta H_{\text{day}}(x, t) & \text{if } E_{\text{obs}}(x) < 2.5 \text{ m yr}^{-1} \\ b'_{\theta} \Delta \theta(x) + b_T \Delta T_{\text{day}}(x, t) & \text{otherwise} \end{cases}$$

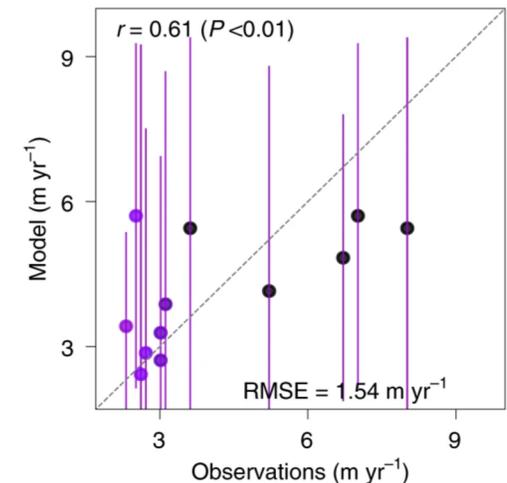
ACD database (Lantuit et al. 2012)
Trained on “High-quality” segments



- 1: Bykovsky Peninsula and Muostakh Island
- 2: Mackenzie River Delta region
- 3: Yamal Peninsula



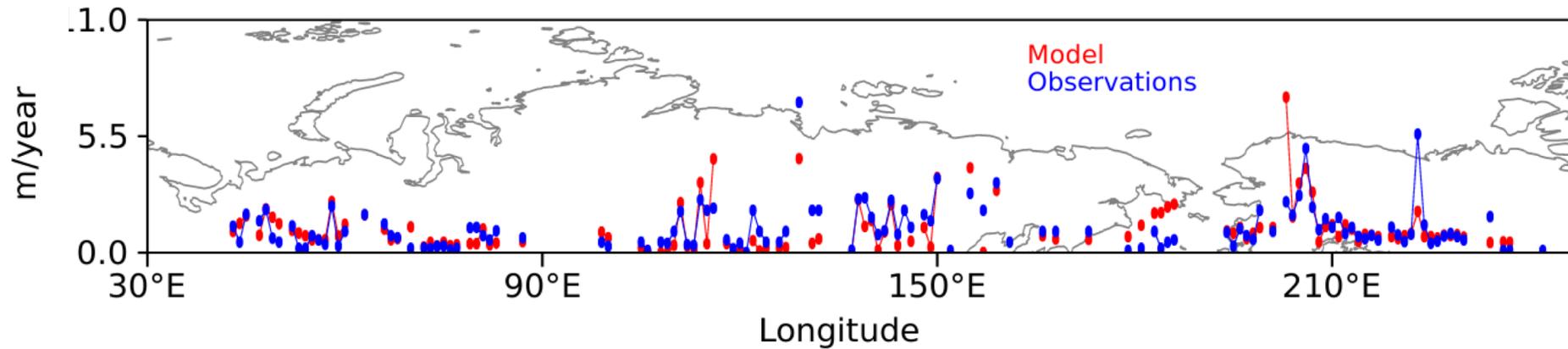
Non-extreme:
waves +
ground ice



Extreme:
temperature +
ground ice

Assumptions

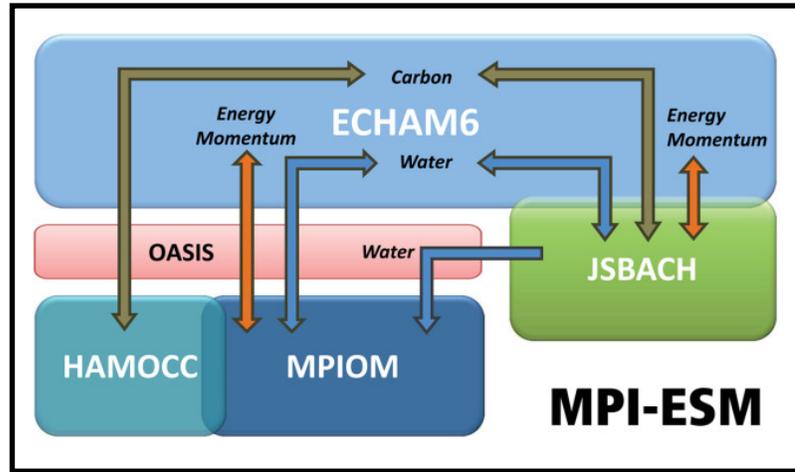
- Only the erosive coastal segments from the ACD (306 out of 1314)
- Only linear additive effect of the thermal and mechanical drivers
 - i.e.: absence of non-linear, synergistic effects
- Ground-ice is constant in the future
- Linear regression model trained with observations is valid in the future



Forcing:

- 10-member Ensemble of MPI-ESM from CMIP6 (Historical and Scenarios)
- Ocean surface waves (WAM)

Earth system model (ESM) simulations including erosion

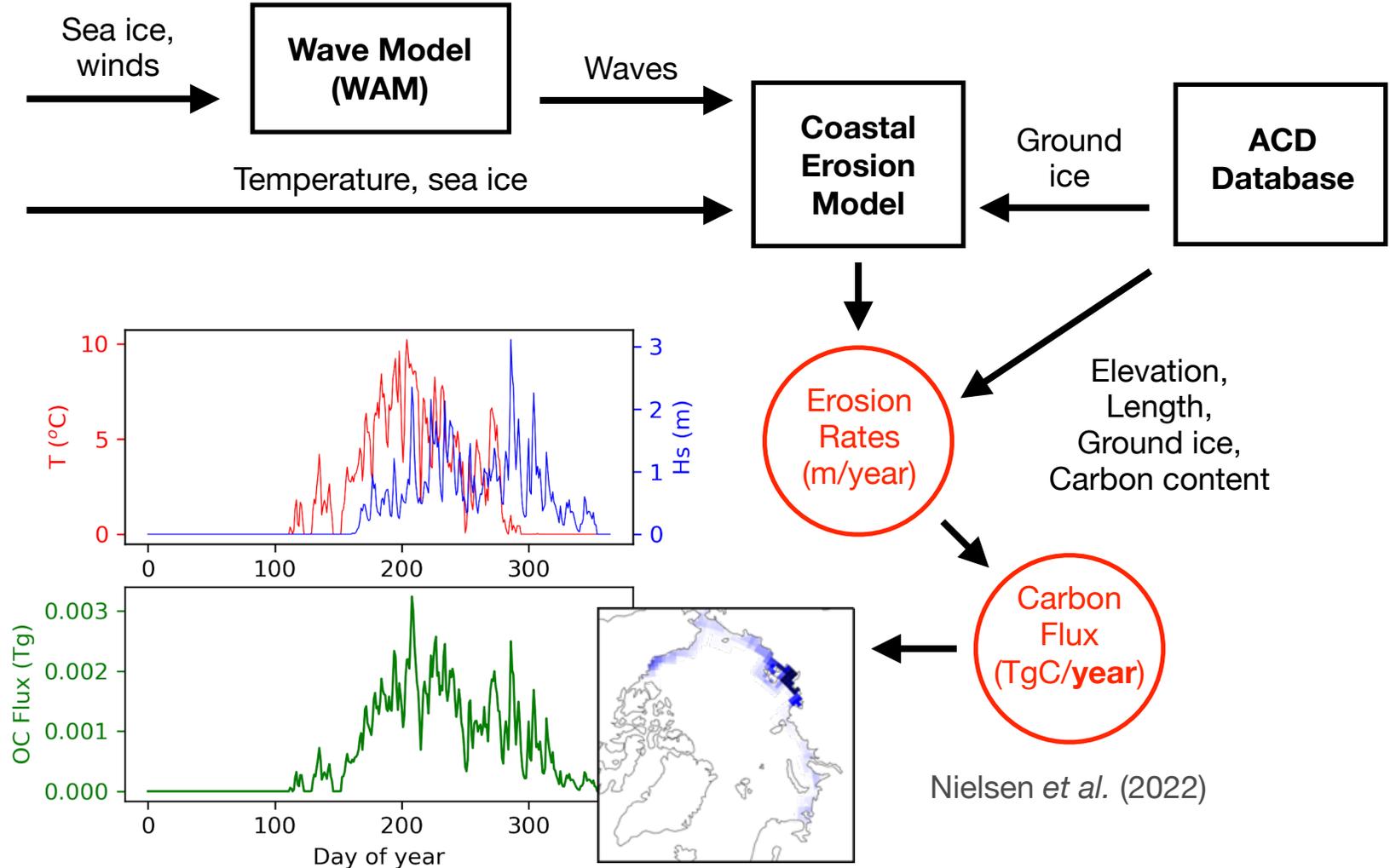
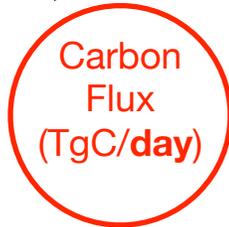


Giorgetta *et al.* (2013)

Changes in ocean biogeochemistry, air-sea CO₂ fluxes

?

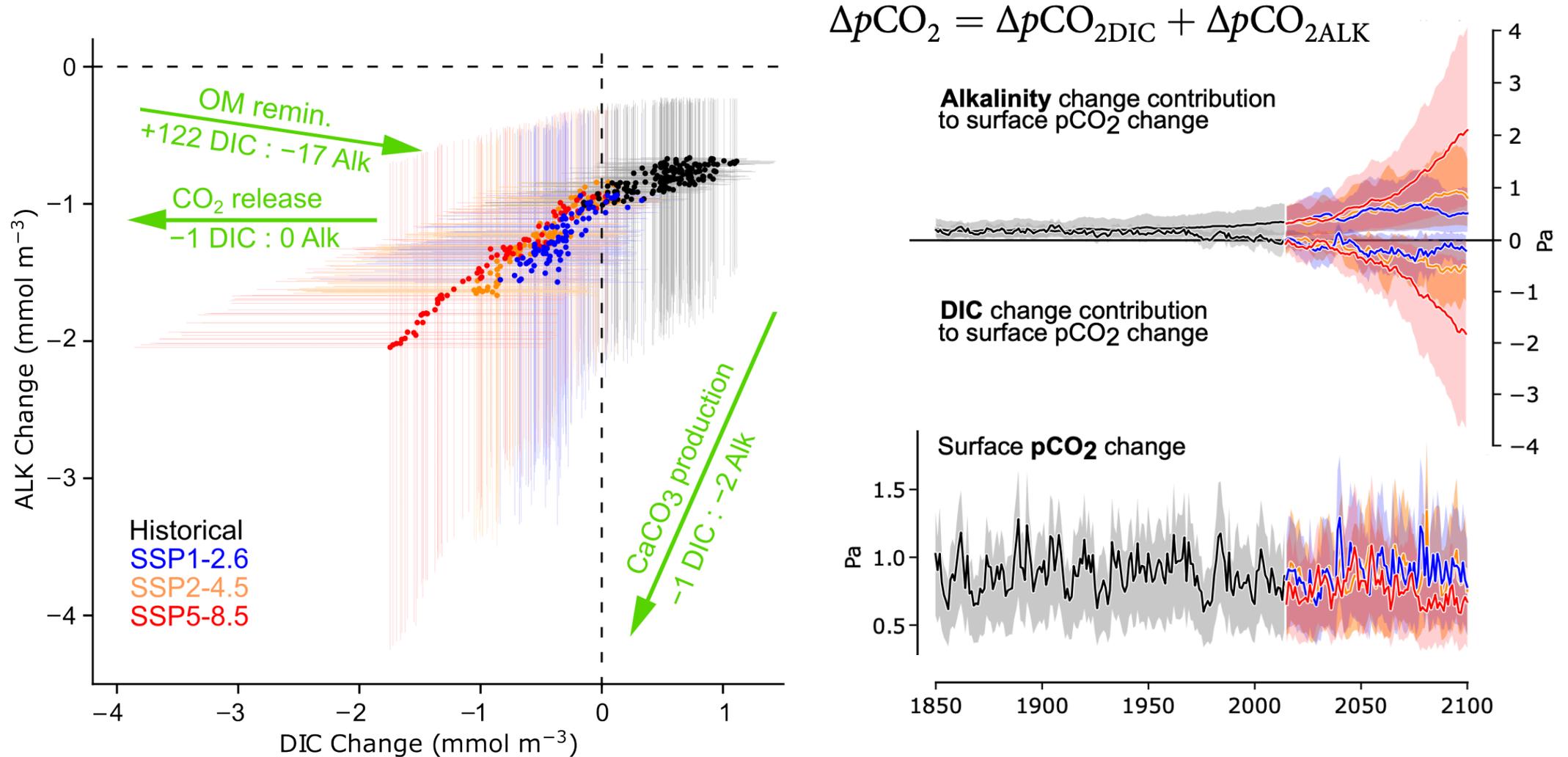
Nielsen *et al.* (in prep.)



Nielsen *et al.* (2022)

Alkalinity and DIC response to coastal erosion

Changes in alkalinity and DIC maintain a relatively constant increase in surface pCO₂

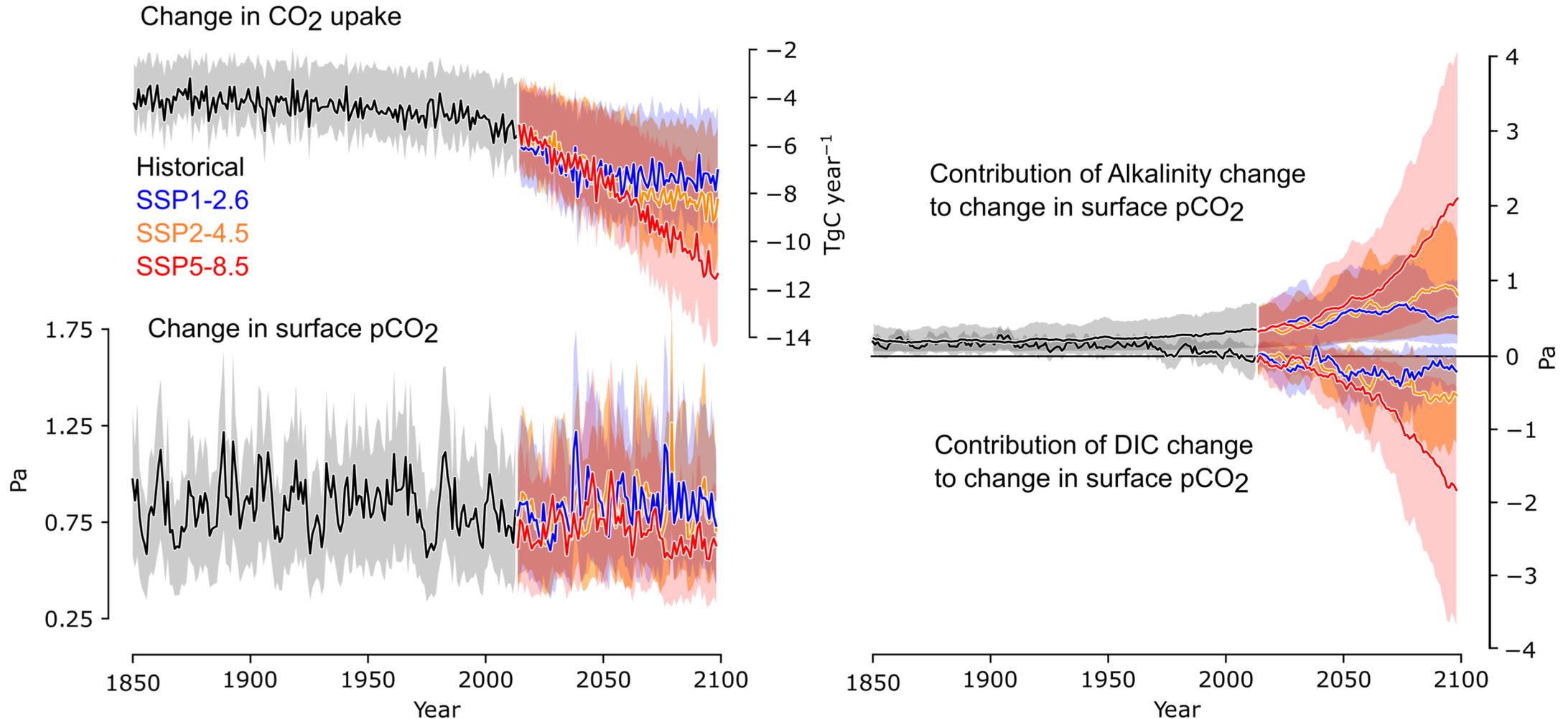


Decomposing $\Delta p\text{CO}_2$

$$\Delta p\text{CO}_2 = \Delta p\text{CO}_{2\text{DIC}} + \Delta p\text{CO}_{2\text{ALK}} + \cancel{\Delta p\text{CO}_{2\text{SST}}} + \cancel{\Delta p\text{CO}_{2\text{SSS}}}$$

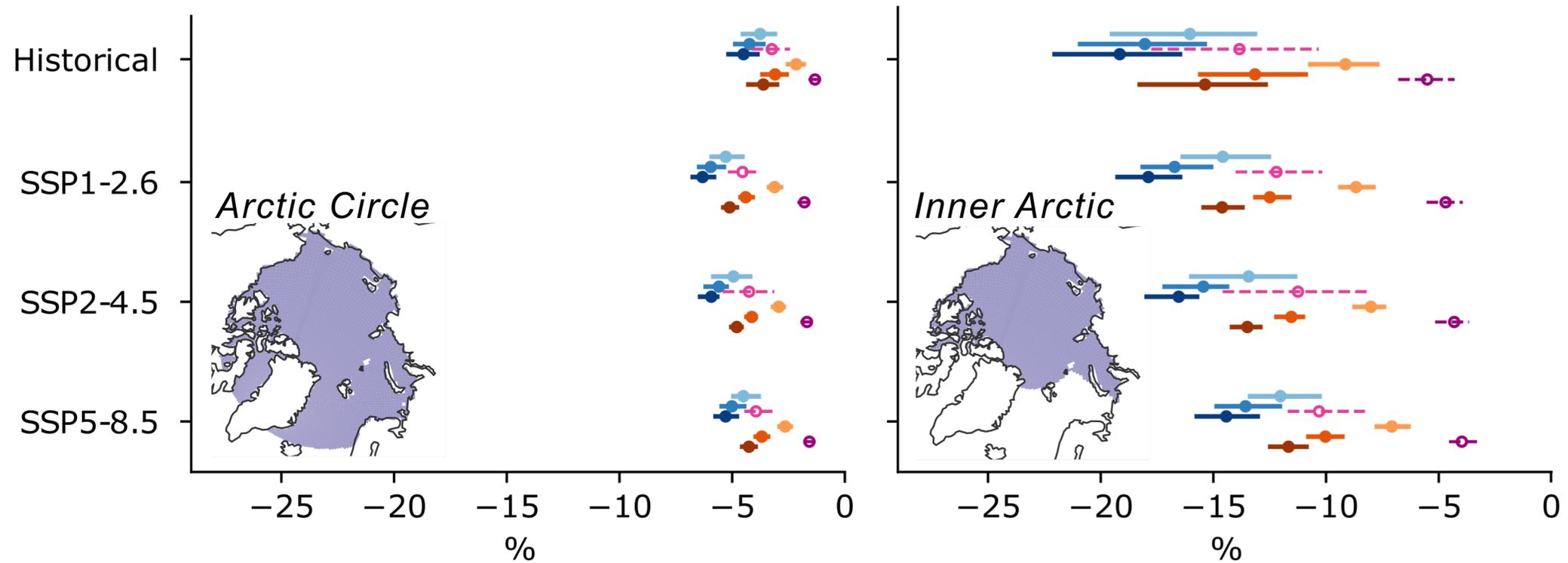
$$\Delta p\text{CO}_{2X} = \Delta X * \gamma_X * p\text{CO}_{2\text{Ref}} / \bar{X}$$

Same physics



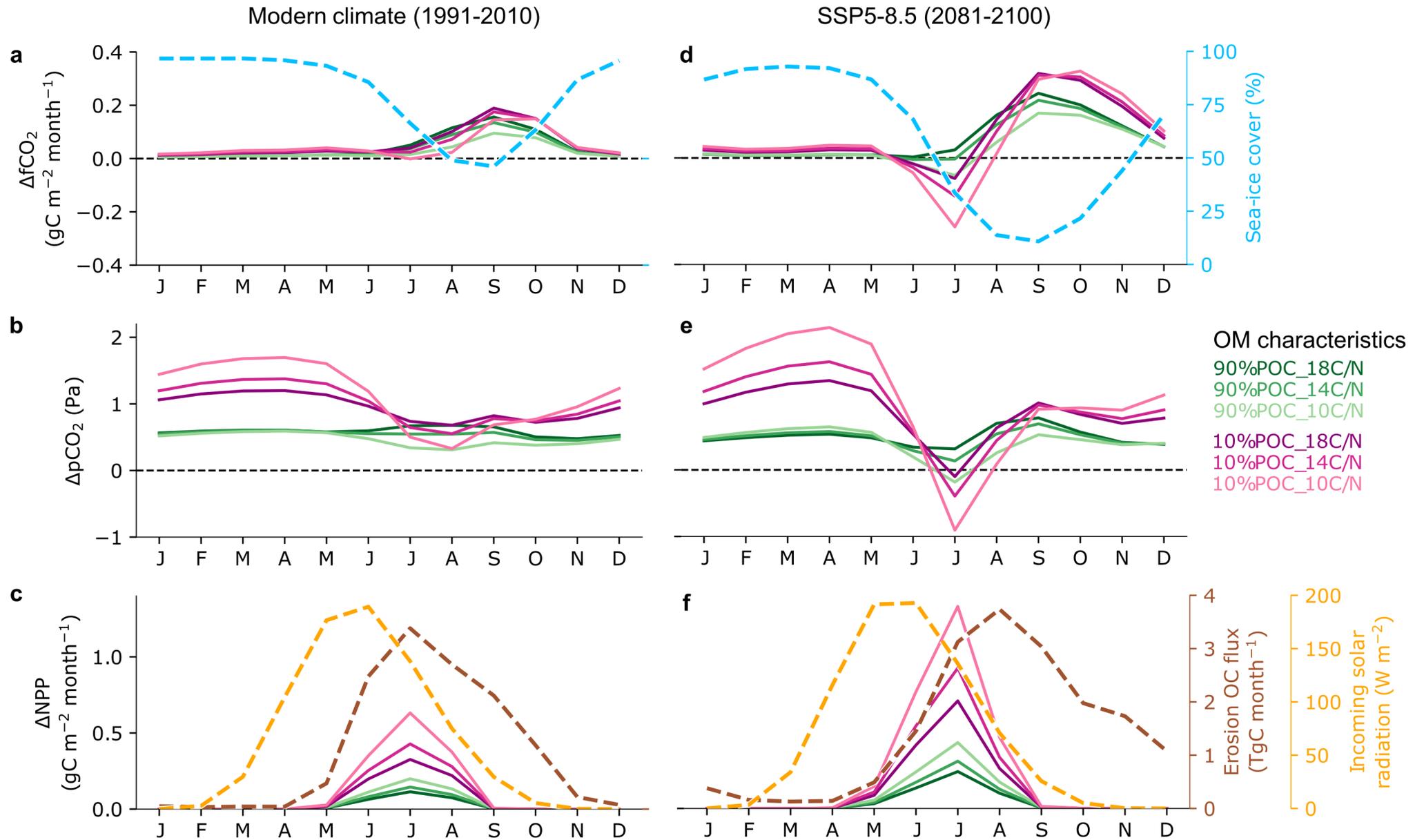
Coastal erosion reduces the Arctic Ocean's CO₂ uptake

Percentage change in surface CO₂ uptake due to coastal erosion

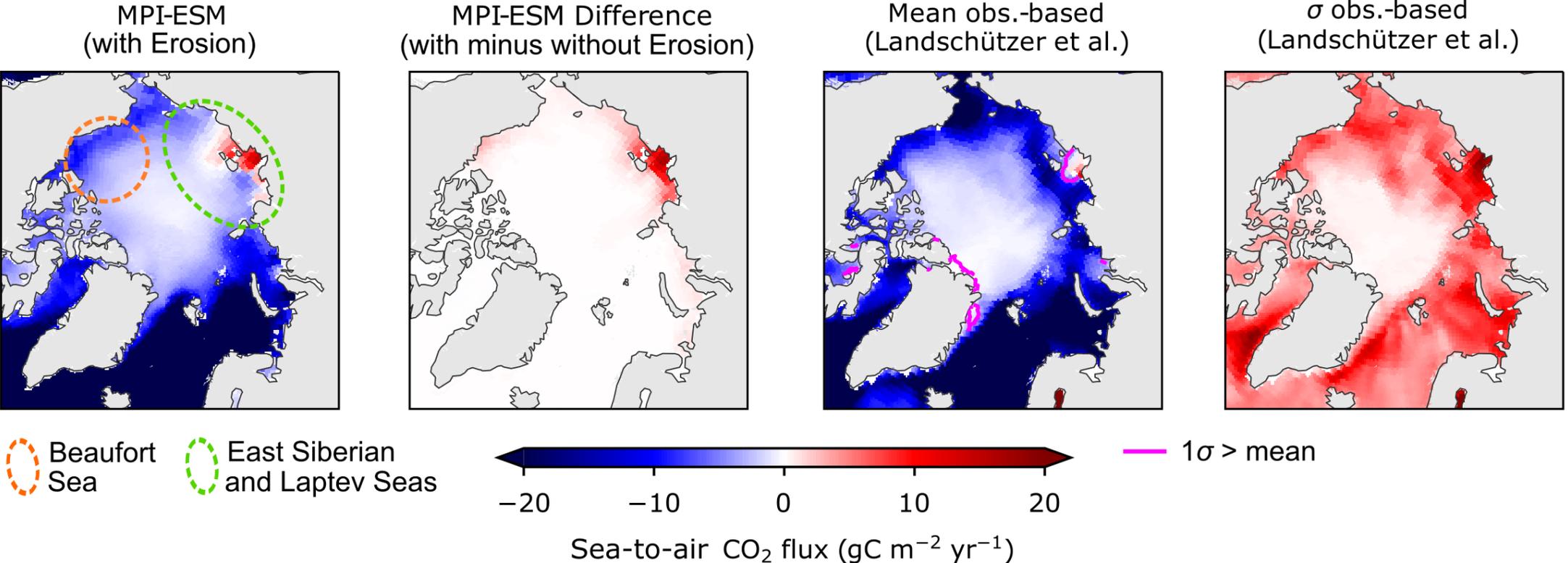


90%POC_18C/N 10%POC_18C/N 90%POC_7.6C/N
 90%POC_14C/N 10%POC_14C/N 10%POC_7.6C/N
 90%POC_10C/N 10%POC_10C/N

Seasonal cycle

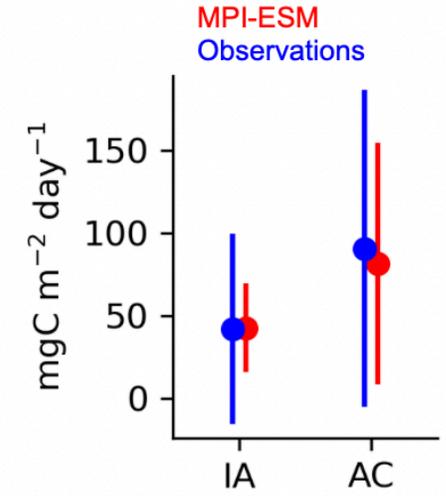
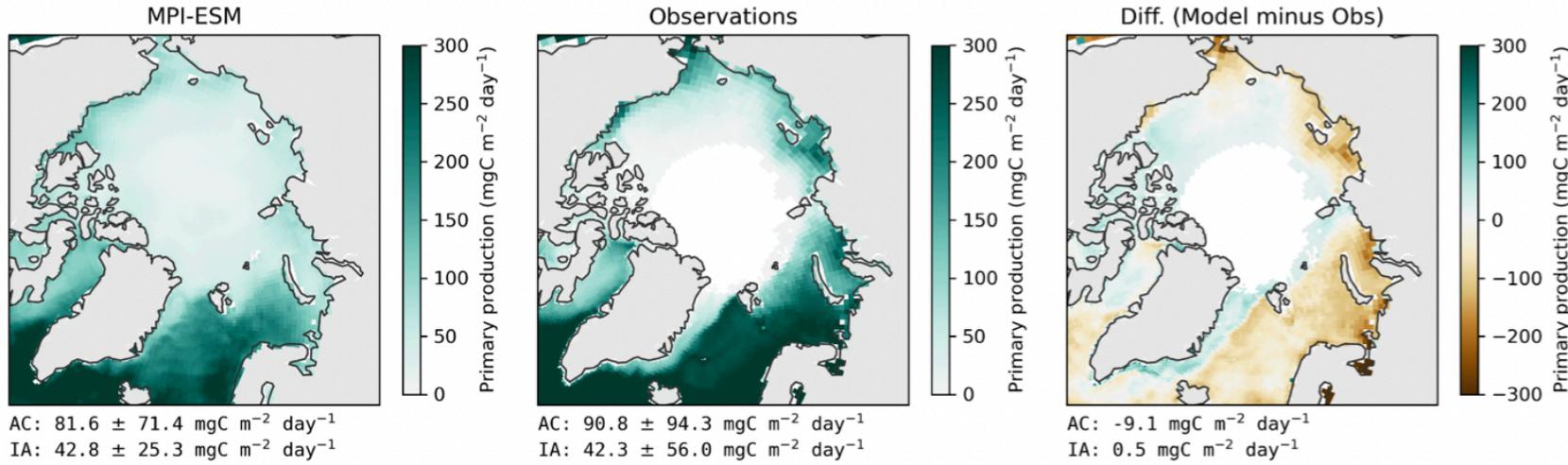


Erosion inverts the air-sea CO₂ flux direction on the Siberian shelf

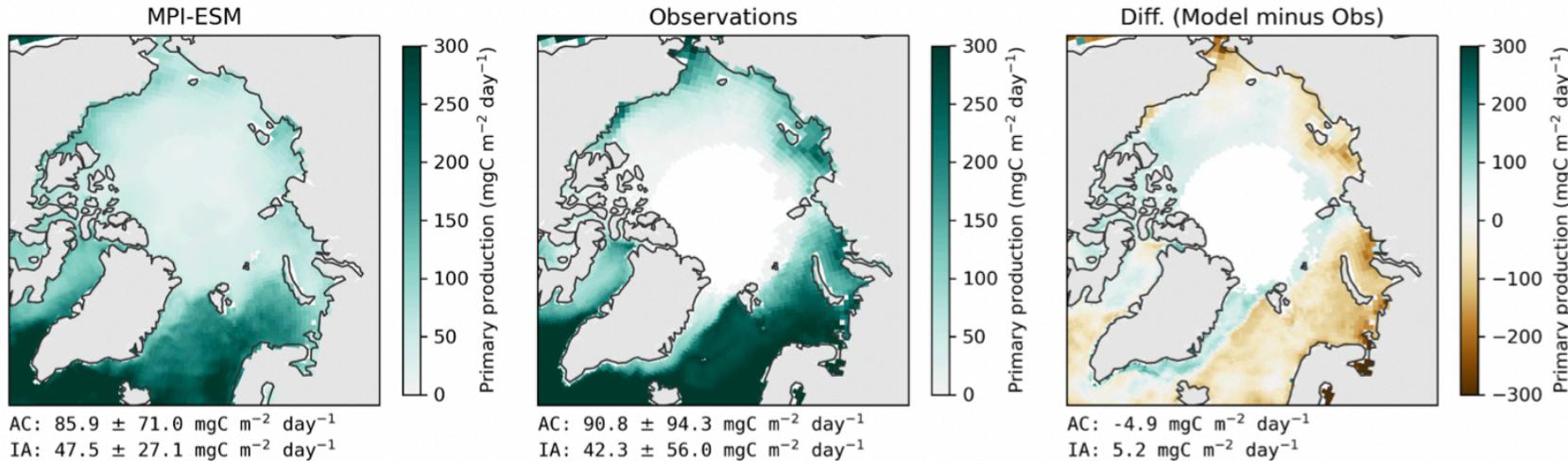


NPP - MODIS CAFE

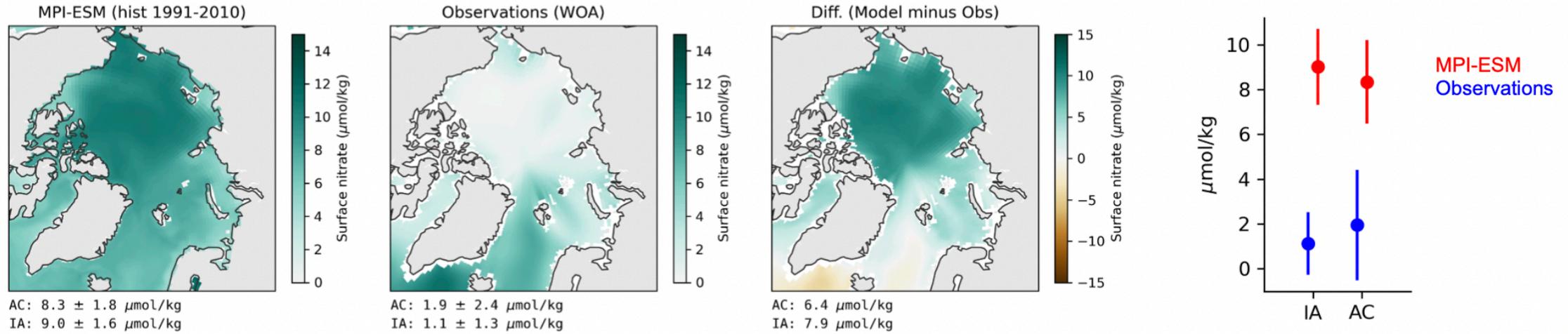
Without Erosion



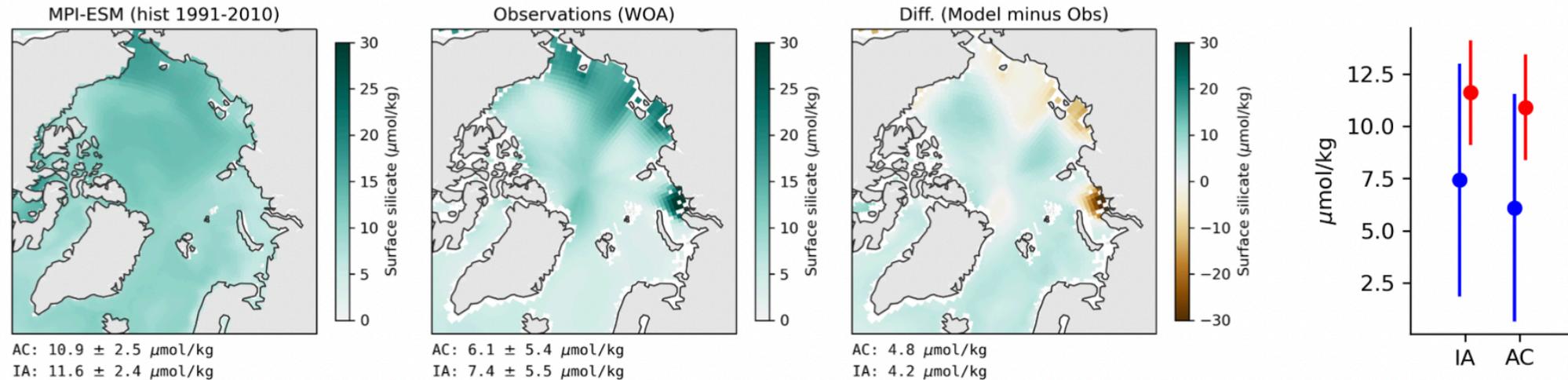
With Erosion



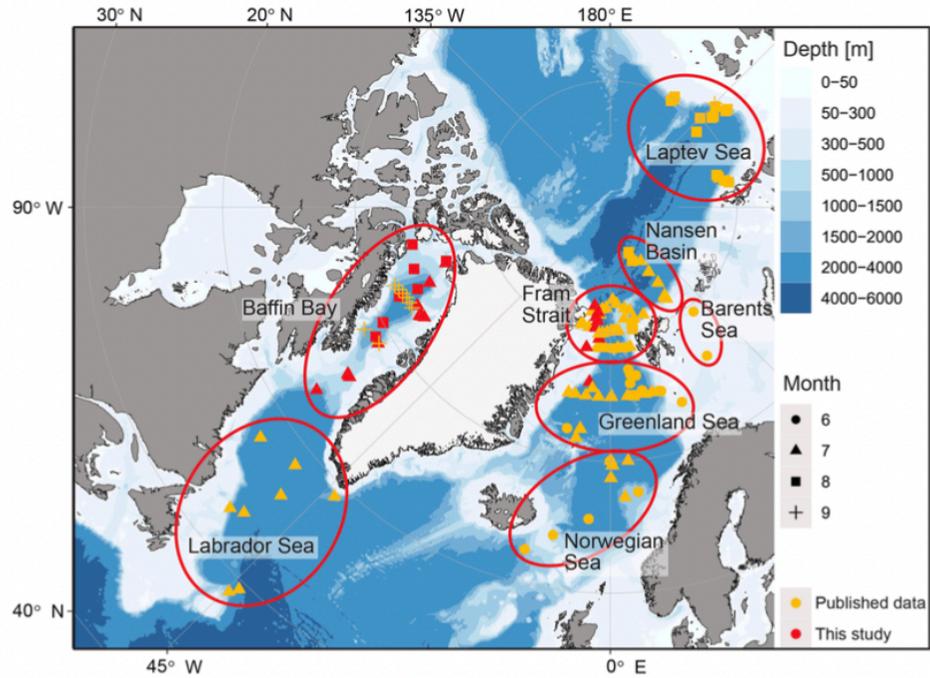
Surf. Nitrate (WOA)



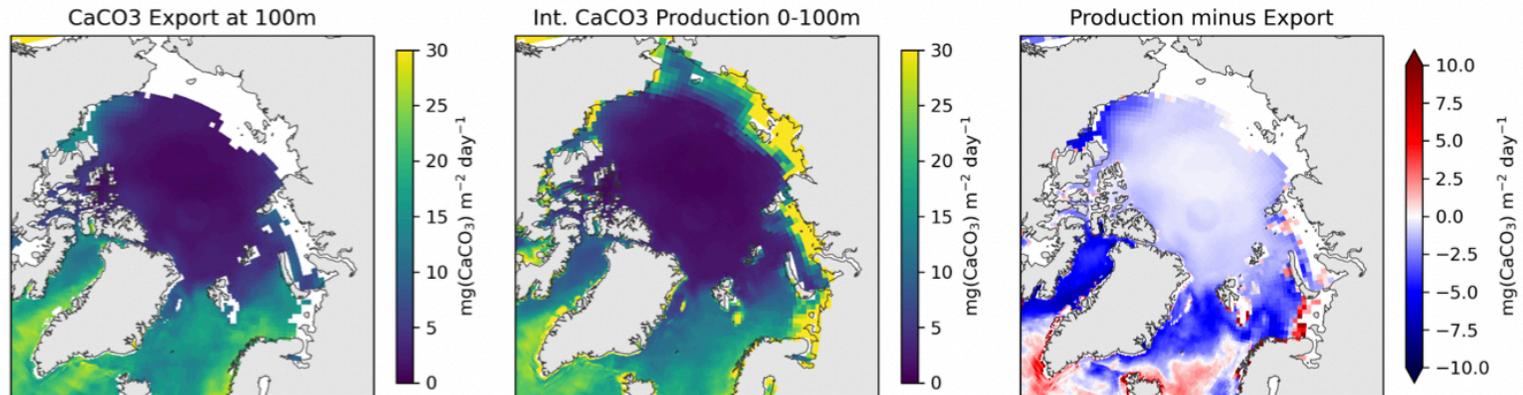
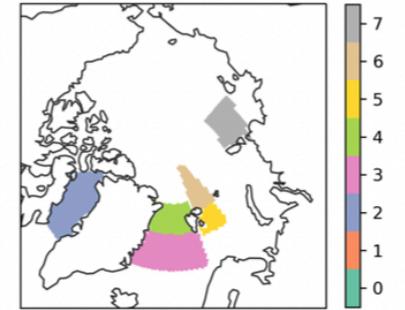
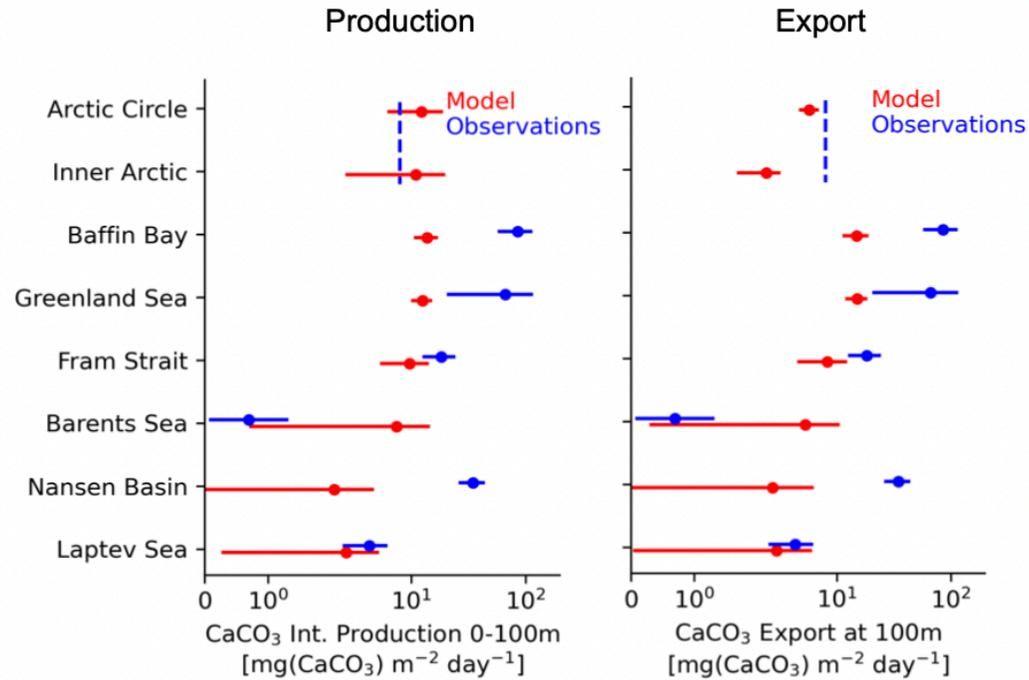
Surf. Silicate (WOA)



CaCO₃ Production

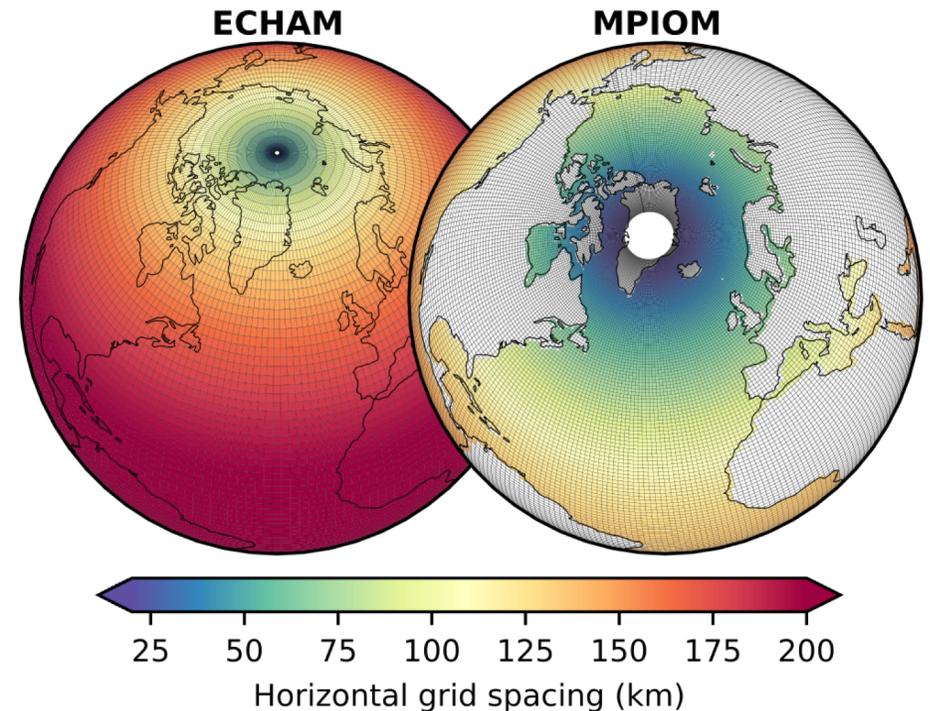
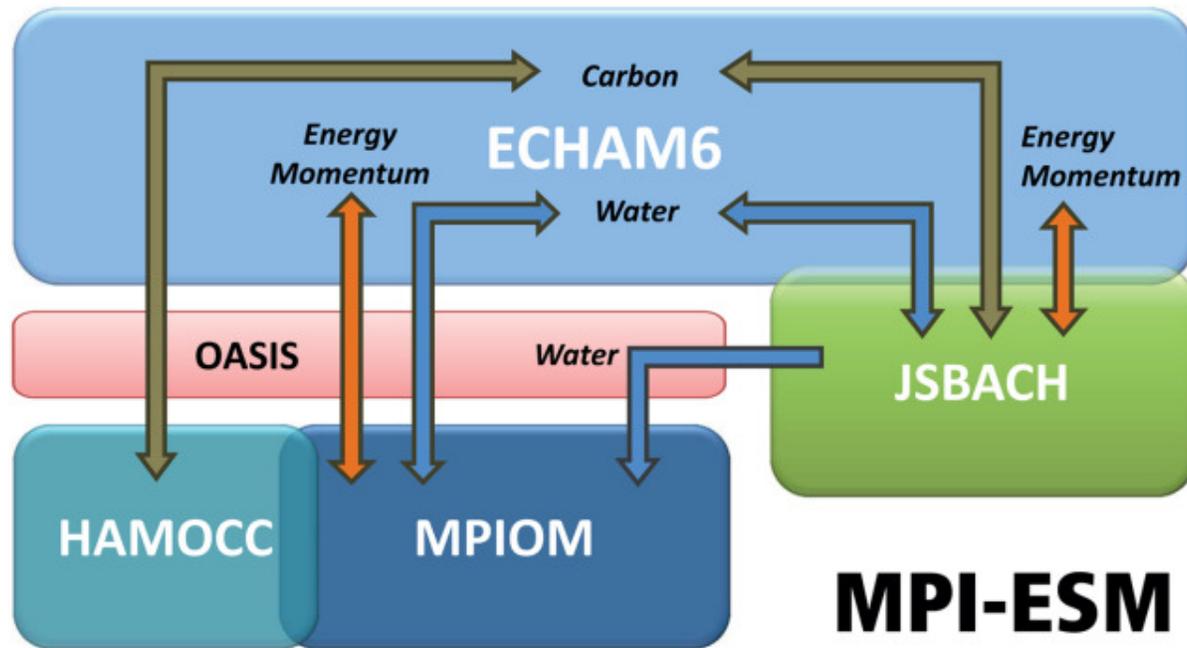


Tell, F., Jonkers, L., Meilland, J., and Kucera, M. (2022) Upper-ocean flux of biogenic calcite produced by the Arctic planktonic foraminifera *Neogloboquadrina pachyderma*, *Biogeosciences*, 19, 4903–4927, <https://doi.org/10.5194/bg-19-4903-2022>



Max Planck Institute Earth System Model (MPI-ESM)

MPI-ESM simulates the global Land, Ocean and Atmosphere, representing the main fluxes of water, energy and carbon. Scenario simulations require spatial resolution on the order of 10-100 kilometers.



Schematics of fluxes between components in MPI-ESM
(Giorgetta et al. 2013, <https://doi.org/10.1002/jame.20038>)